

M.Sc. Thesis

**GROUNDWATER RECHARGE OF LAHORE CITY BY PONDING IN
RIVER RAVI**



Advisor

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ABSTRACT

Groundwater plays an important role for domestic, agriculture, urban and in industrial purposes for the city of Lahore. Due to continuous increase in population and people moving towards the urbanized areas causes the extra groundwater extraction cause depletion in aquifer. To enhance groundwater recharge and control of groundwater depletion a dam is proposed at 3 mile Upstream of Mohlanwall and 5 mile downstream of motorway bridge. The height of the proposed dam is 10 ft against which a back watering of 12.5 mile was simulated using HEC-RAS.

HEC-RAS model was used to check the backwater effect in river Ravi due to dam structure. Model was calibrated using manning roughness values between 0.015-0.025 for the main channel, left over bank (LOB) and right over bank (ROB).

Recharge rate of the river Ravi was calculated by Zhukovsky's method against a maximum discharge of 18000 cusec observed in last two decades and found out that River Ravi has a recharging capacity of about 6978783 ft³/year. The slope of the River was found to line 1.4ft/mile which was from the longitudinal section of River Ravi. Visual MODFLOW was run to check the effect of river stage on the groundwater level. Pumping data of about 3200 tube wells were used for modelling and R² value of about 0.9 was observed during calibration period. A groundwater level of about 20-40ft was observed in the surrounding areas of River Ravi and about 3ft depletion rate in groundwater level was found under normal circumstances. After the proposed dam the extra recharge capacity 24666069 ft³/year of River Ravi was calculated.

A rise of 3-6 ft. was simulated in the areas close to the proposed site with a coverage area of 36.33 mile² on the left side of the river towards Lahore site. It is recommended to install more piezometric tube wells to check the groundwater levels situation and to maintain water levels more recharging options should also be considered.

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CHAPTER I

INTRODUCTION

1.1 General

Lahore is the capital of Punjab province and the second largest and metropolitan city in Pakistan. It is located on the world map with geographical coordinates as 31° 32' 58" N and 74° 20' 36" E. With its proximity to Kasur district in the South and Sheikhpura in the North-west, the district area is 1,772 km² and its urban area comprises largely of converted agricultural lands (Figure 1.1). Urban area of Lahore has expanded almost double in the last 12 to 15 years. According to the 1998 census, there were nearly 6.32 million people; a population density of 3,566 persons/Km². Urban population consisted of 82.4% of the total. Lahore is the second largest financial hub of Pakistan and has industrial areas including Kot Lakhpat and the new Sundar Industrial Estate (near Raiwind). Thus, rapid urbanization of the area, including the industrial establishments, is posing many key challenges, and at the same time, opportunities in the new millennium. Out of these challenges, sustaining the water and sanitation demands of the residents is an emerging issue. (Lahore master planning, 2016)

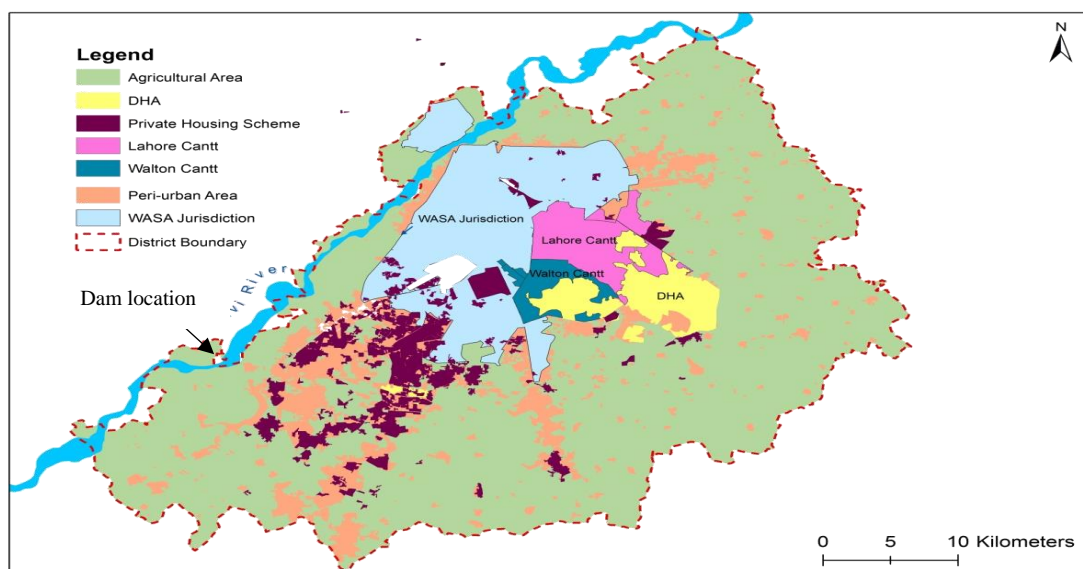


Figure 1.1: Land use map of Lahore District showing River Ravi and proposed dam site

1.2 Climate and Geography

The climate of the region is characterized by large seasonal variations in temperature and rainfall. Mean annual rainfall is 712 mm and reference crop evapotranspiration (ETO) is 1,649 mm. Mean annual temperature is about 24.3 °C ranging from 33.9 °C in June to 12.8 °C in January (Pakistan Meteorological Department, 2006). The hottest month is June, where average highs routinely exceed 40 °C, highest temperature observed is 48.3°C, recorded on May 30, 1944, and the lowest recorded is –1°C, observed on January 13, 1967. The wettest month is July, with heavy rainfalls and evening thunderstorms with the possibility of cloudbursts. (Lahore master planning, 2016)

The Lahore District forms the part of the Bari Doab – inter-fluvial lands between the Ravi in the North West and the Sutlej River to the south. The area is underlain by a significant thickness of alluvial deposits, up to 1200 ft in depth, proven by the investigations carried out by Water and Soils Investigation Division (WASID), during the period 1961-62 (WAPDA, 1980). The recharge sources are Ravi River (occasional flows, nowadays) on the North West and BRBD Canal on the East. Discharging boundaries are Sukh-Beas drainage channel on the south and again the Ravi River in the west. However, these discharging boundaries only function in the form of surface runoff during heavy rainfalls; no groundwater discharge is possible due to excessive groundwater depletion in the area. The general altitude of the area is 682 to 699 ft, above mean sea level (amsl). The area slopes towards the south and south-west, at an average gradient of one in 3000. The upper most part of the upper Bari Doab is now Indian Territory, after partition, but being the integral part to the Lower and Central Bari Doab provides natural route for the regional groundwater flow.

1.3 Problem Statement

Due to Rapid population growth, intensive migration of people to Lahore and the establishment of several industries has increased water demand by manifolds. Mushrooming private housing societies, particularly on the periphery of cities, along with industrial activities has caused two-fold stress on the underlying groundwater reservoir, i.e. increased groundwater pumping for domestic and industrial activities and reduced groundwater recharge from agricultural fields and rainfall due to most urban areas being covered by buildings and roads.

Despite a reduction in Ravi flows due to upstream water use by India the main recharge (82 per cent) to groundwater is contributed by the river. The rainfall and canal system contribute only 12 percent whereas the return flow from irrigation fields is about 6 percent. This shows the importance of Ravi flows in sustaining the Lahore aquifer

WASA supplies drinking water to more than 6 million people by means of 484 tube wells. These tube wells are located in different area and their depth varies between 490 to 650 ft. Over time water, demand has increased from 180 litres per capita per day (lpcd) in 1967 to 274 lpcd in 2013. (WASA Report, 2013)

Due to these reasons, it is necessary that such methods and techniques are to be adopted that could minimize the ground water depletion and enhance groundwater recharge. The Ravi river is a major source of groundwater recharge, which ultimately increase the groundwater recharge capacity of the Lahore city

1.4 Study Area

The study area of the present research was from Ravi siphon to katcha kahna near Chung Lahore. The length of the reach is nearly 38 mile, and cross section varies at different locations. The area adjacent to the Ravi River is urban and mostly populated. Therefore, it considered that if the dam was constructed in that area and if the high flow

comes in moon soon season what will happen to the adjacent area and how the effect of dam will be spread in that area. Study area with proposed dam location is showed in Figure 1.1

1.5 Study Objectives

- To evaluate the availability of the groundwater recharge capacity from river Ravi without and with dam.
- To determine the water surface profile in case of high flow floods in River Ravi.
- To determine the groundwater recharge capacity and variability in water level of Lahore city by the river by using VISUAL MODFLOW model.

1.6 Limitations of Study

The research only deals with the ground water recharge of the Lahore city and the effect of dam construction in Lahore city in case of the high flow flood. The source of water in river Ravi may be discussed, so that it can be understood that how much discharge is flowing in river Ravi and what contribution is from rainfall, and drainage system of Lahore. It will only deal with the ground water recharge through the river Ravi and not from any other source. The study did not include the effect of recharge on the adjacent cities like Sheikhpura because proposed dam and its backwater effect is not recharging to the Sheikhpura city. Water quality will not be discussed in this study.

1.7 Utilization of Research

This research will help to identify the groundwater potential of Lahore city from the River Ravi and what is the current condition of groundwater levels, and then from this study methods can be devised that help to reduce the groundwater depletion in Lahore city. This research would be helpful to check the current situation of Lahore district and how the groundwater levels is going to vary in the future, so the Irrigation & Agriculture departments can work on the planning to limit the groundwater use for future.

CHAPTER II

LITERATURE REVIEW

2.1 Sources of Water

2.1.1 Surface Water Resources

The main source of surface water in Pakistan are Rivers. In the Indus Water Treaty 1960, water rights of three eastern rivers (Ravi, Beas and Sutlej) were given to the India and three Western Rivers (Indus, Jehlum and Chenab) were given to Pakistan. Indus River and its Tributaries on an average brings 190 Billion cubic meter (BCM) of water annually. This include 179 BCM from the western River and 11 BCM from Eastern Rivers whose water rights is to India.

Currently, 93 percent of total water withdrawal (176.7 BCM) was distributed for the Agricultural use, 4 percent (7.6 BCM) is for domestic use and the remaining 3 percent (5.7 BCM) is for industrial use (Bakshi and Trivedi, 2011)

The flow In River Ravi mainly comes from Marala-Ravi link Canal and five flash streams, which flow, into Rive Ravi within Pakistan. In recent years, the flow of River Ravi declined due to climate change and construction of the irrigation and Hydropower diversion structure by the India. The Thein Dam constructed by India in 2000 impact the flow hydrology of the river Ravi that has reduced groundwater recharge significantly in and around the Lahore city.

The average flow in Ravi during 1922-1961 was 1300 million cubic meter per day MCM/day which was reduced to 800 MCM/day in 1985-1995, and further declined to its lowest level of 175 MCM/day in 2000-2009. However, peak flows of more than 260 MCM/day was also observed during this period (Basharat and Rizvi, 2011).

During the monsoon of 2010, a highest discharge of only 75-100 MCM/day was reported in River Ravi in spite of heavy monsoon on the regional scale (Khalid et al., 2013). This means that even in future the floods passing along the Lahore will be only from Ravi tributaries downstream of the Thein Dam or the flood that is more than the reservoir capacity of the Thein Dam. Therefore, except flood events from during an extraordinary rainy season, no regular flow of considerable amounts are expected except for M-R link canal releases from Marala Barrage.

2.1.2 Groundwater Resources

Most of the groundwater resources of Pakistan exist in the Indus Plain, extending from Himalayan foothills to Arabian Sea, and are stored in alluvial deposits. The Plain is about 1,600 Km long and covers an area of 21 Mha and is blessed with extensive unconfined aquifer, which is fast becoming the supplemental source of water for irrigation. The aquifer has been built due to direct recharge from natural precipitation, river flow, and the continued seepage from the conveyance-system of canals, distributaries, watercourses and application-losses in the irrigated lands during the last 90 years (Khalid et al., 2013).

The Lahore aquifer was composed of unconsolidated alluvial soil of up to 400 m in thickness. At present, the maximum observed water Table depth in Lahore city was 40 m and in the area of Raiwind was about 12 m (Basharat and Rizvi, 2011).

In 1960, the groundwater Table depth of Lahore city was 4.6 m. More use of the groundwater Table depth has headed to lowering of the water Table about 0.5 m for the last 30 years. In 1987 the depth of water Table ranging from 8-20 m, which has reduced further to 51 m in 2011. (Khalid et al., 2013)

WASA supplies water to the 6 million people by mean of 484 tube wells in Lahore city with a groundwater depth of 200-400m (WASA 2013).

2.1.3 Groundwater Recharge Techniques

Groundwater recharge techniques can be classified into Direct Methods and Indirect Method. Direct methods include Surface methods and Subsurface methods. In surface methods flooding, recharging basins, stream modification, and surface irrigation are included while in subsurface groundwater, recharging techniques include injection wells (recharging wells), gravity head recharging wells, and recharging pits. (Debu 2016)

Indirect method for groundwater recharge includes collector wells, infiltration gallery, and groundwater conservation structure development for the indirect groundwater recharge. (Debu 2016)

For the current study, the selected method for the groundwater recharge was ponding in river and a groundwater model in Visual MODFLOW was developed to evaluate the effect on groundwater recharge with different scenarios.

2.2 Model Background

2.2.1 HEC-RAS

HEC-RAS is a model developed by the hydraulic engineering centre of US Army Corps of Engineering. HEC-RAS is an integrated system of software. It is comprised of the graphical user interface (GUI), separate analysis component, data management capabilities and graphics. It is used for the analysis of the river system.

HEC-RAS contains the following different components of river analysis system.

- Steady flow water surface profile computation;
- One dimensional and two-dimensional unsteady flow simulation;

- Water quality analysis; and
- Quasi unsteady or fully unsteady sedimentation analysis.

HEC-RAS is designed to perform one-dimensional and two-dimensional hydraulic calculations for a full network of natural and constructed channels, overbank, levee protected areas etc.

For two-dimensional unsteady flow analysis, HEC-RAS uses Navies-Stokes equation, which describe the motion of fluids in two dimensions. While in the context of channel and flood modelling, further simplification could be done.

Now in HEC RAS model RAS Mapper is also available which help us to georeferenced our cross section more easily because we can add a terrain of the study area in the back of our model (HEC-RAS User Manual, 2016).

2.2.2 Groundwater Visual MODFLOW

The Visual MODFLOW software (2011.1 ver.6.0.159), was developed by Schlumberger Water Services (SWS), Canada, has been used for this study. The Visual MODFLOW is the most complete and easy-to-use modelling environment for practical applications in three-dimensional groundwater flow and contaminant transport simulations. This fully-integrated package combines MODFLOW, MODFLOWSURFACT, MODPATH, Zone Budget, MT3Dxx/RT3D, MGO, and WinPEST with the most intuitive and powerful graphical interface available.

The model input parameters and results can be visualized in 2D (cross-section and plan view) or 3D at any time during the development of the model, or the displaying of the results. For complete three-dimensional groundwater flow and contaminant transport modelling, Visual MODFLOW is one of the best software packages. The main interface for Visual MODFLOW has much of the same user-friendly look and feel as the

previous versions of Visual MODFLOW. Some of the more significant upgrade features in the latest version of Visual MODFLOW are described below.

The logical menu structure and easy-to-use graphical tools allow users to:

- easily dimension the model domain and select units;
- conveniently assign model properties and boundary conditions;
- run the model simulations;
- calibrate the model using manual or automated techniques;
- optimize the pumping well rates and locations; and
- visualize the results using 2D or 3D graphics.

The main interface for Visual MODFLOW has much of the same user-friendly look and feel as the previous versions of Visual MODFLOW, but what's "under the hood" has been dramatically improved to give you more powerful tools for entering, modifying, analyzing, and presenting your groundwater modeling data (VMOD help)

A number of scientists and scholars used mathematical modelling in different area of groundwater like planning, development and managing groundwater studies. In this chapter, the studies done by the different researchers were reviewed and main findings were summarized.

2.3 HEC-RAS Modelling

Different Researcher used HEC-RAS model for the water surface profile computation and flood inundation for different scenario. Some of these researches are reviewed below for understanding the HEC-RAS modelling.

Bashir et al, (2010) studied hydrological modelling and flood hazard mapping for Nullah lai using HEC RAS and HEC GeoRAS. GIS model with hydrologic modelling was used to estimate of flood zone of Nullah lai, Rawalpindi. Topographic survey of fine resolution was used to create DEM. Interpolation method was used to obtain

elevation data at unknown locations. Inundation area estimated at the discharge value of $3000 \text{ m}^3/\text{sec}$ is 3.4 km^2 out of which 2.96 km^2 is occupied under the inundation depth from 1 to 5 meters. Maximum flood depth of 20 m was observed at this discharge for the study area. Output values of flood depth and flood area obtained by HEC-RAS was closed to the Survey done by Japan International Cooperation Agency (JICA).

Masood et al, (2011) developed a HEC-RAS model for the flood inundation of eastern Dhaka by using DEM from Shuttle Radar Topography Mission and hydrologic data for 30 years (1972-2004) is used for the study. The DEM was modified to cover the topography of the area because of high land filling by the developer. The model was run for the flood of a 100-year return period for the flood inundation of the eastern Dhaka. Simulation results shows a maximum depth of 7.55 m and south-eastern part of the study is flooded more than 50%. A flood Hazard map and Risk analysis map of the area was developed using ArcGIS considering depth of flow and vulnerability to people and property. These maps are helpful for the future management of floods.

Hagizadeh et al, (2012) studied the Neka's River Basin, Malaysia for the flood analysis using HEC RAS and HEC GeoRAS and water surface profile was developed for the Downstream of Neka's River. Hydraulic studies during the flood season of 1986-1999 and 2002-2004 was used to calibrate the model. 200 X-Section of the rivers was used for the model input. Model was calibrated by adjusting the roughness coefficient. The peak flood was well simulated at Nika's River Basin with mean square error of 10.7 cm compared to observation. Simulations of peak flood stages and hydrographs' evaluations are congruent with studies and observations, with the former showing mean square errors of 4.8 – 10 cm. The coupled HEC-RAS and HEC GeoRAS shows good results.

Siddique et al, (2011) compared the water level estimates from hydrodynamic hydrologic modelling and satellite altimetry for a complex environment. For hydrodynamic modelling he used 1D calibrated, HEC-RAS model for the rivers in Bangladesh, for the model input gaged stream flow; water level data and river geometry were input. On the other hand, Envisat mission was selected for satellite-based water level elevation data. Results showed that the valued obtained from satellite data is 0.2 m and 1.9 m higher than the values estimated by using HEC RAS in monsoon and Dry conditions respectively. Envisat satellite was found to be disagreed with most of the small and medium river in mountainous and flashy flood basins.

Sarhadi et al, (2012) studied flood inundation mapping of ungauged rivers at Halilrud basin and Jiroft city in south-eastern Iran. Land used and the channel morphology was extracted using GIS techniques. HEC GeoRAS pre-processing files was created in GIS, which is then imported to HEC RAS. Discharge at different return periods was calculated using frequency analysis and used in the flow inundation computation in HEC RAS which was then exported to GIS for better visualization satellite imagery was used to find out the extent of the model in real time. Previous floods data was used for the model performance evaluation in predicting flood inundation. Then map was prepared to check the vulnerable area form the extracted images from the satellite and flood inundation extent to find out the critical areas.

Baky et al, (2012) studied flood flows for the crop production and risk management with HEC RAS and GIS in the Agricultural area of Shariatpur, Dhaka. He prepared inundation maps of the area by using hydrodynamic models of HEC-RAS and HEC GeoRAS in Arc GIS. Optimal flood depth for the crop was estimated out using statistical analysis and finally water above the optimal depth was calculated using GIS techniques. This study also proposed two equation for the estimation of MAWA

discharge depending on 2 upstream stations. This equation helped to find out the amount of water required to release from MAWA to the study Area for obtaining the optimal flood depth for crop production.

Francisco et al, (2015) studied water surface profiles developed in river flows. The HEC-RAS model usually computes the water surface profile. They studied 13 different discharge prediction methods for the computation of water surface profile for compound channels where both momentum and energy equation was applied, based on his results they developed a new theory for the computation of water surface profile in rapidly varied flow. It was found that the interaction of main channel and flood plain is compulsory for the discharge prediction as well as velocity correction coefficient and water surface profile in case of gradually varied flow. It was also found that the results of momentum equation are better than the energy equation.

Vladimir et al, (2015) studied the flood Hazard mapping for Mert River, Samsun, Turkey using GIS and HEC-RAS model. During the 2012 flood a large loss was observed so flood risks maps for the study area was developed using the model for future management of floods. The x- section is digitized and extracted using GIS and then HEC-RAS model was run on peak flow of 10, 25, 50 and 100 year to determine the effect of peak flow in different return period. Then flood maps were developed in GIS considering the Vulnerability and Risks at flood peaks of 10, 25, 50 and 100 year return period.

Weicia et al, (2015) studied threat assessment of Cirenmaco glacial lake in Zhangzangbo valley, Central Himalaya's. He used HEC RAS 2D modelling for the failure of moraine dam Cirenmaco and estimate' glacial lake outburst flood in downstream due to dam failure. He finds out that HEC-RAS produce reasonable depth and extent of flood and can use for the modelling of complex glacial lake outburst

floods effectively. He finds out that HEC-RAS glacial Lake Outburst floods results can be used for the implantation of disaster management techniques.

Punys et al, (2015) studied hydro kinematic resource assessment in a lowland river in lithuania. For this study, hydrologic and hydraulic modelling methods were applied. HEC-RAS with HEC GeoRAS was used for the hydraulic computation. High Resolution DTM was embedded in study for extraction of cross section. Model was calibrated by adjusting the manning n values after computation of slope and stream discharge was measured. After model calibration model was simulated at mean annual, bankful, at low discharge is run and velocity, depth, and cross-sectional area was computed for the river cross sections.

Wicktoria et al, (2016) used HEC-RAS for modelling the flood capacity of a Polish Carpathian river. He compared the results with constrained channel and a free channel condition. For model input, he used 10 pair of river x-section adjacent to constrained and freely developed channel with channel slope and roughness for the Hydraulic modelling of the river using HEC-RAS 1D steady state analysis. The sections developed with freely cross section usually shows the lower water depth and larger width at a particular discharge. They also showed low velocity and bed shear stress. While the cross sections with constrained channel show different results for low frequency high flood. He found that only low frequency high magnitude can be carried out with low shear stress on wide channel x-section.

2.4 Groundwater Modelling

Fouepe et al, (2009) conducted a study for the groundwater modelling in the upper Anga'a River Watershed, Yaonde, Cameroon. MODFLOW software was used for modelling the flow and particle-tracking model for shallow unconfined aquifer for the Anga'a River watershed. The model grid was created of 106 x 68 columns and rows

with two layers each of 50 m depth consisting unconfined and semi-confined aquifer. Lateral inflow and outflows were simulated by constant head boundary condition. Riverhead and Lake Levels in the model were simulated using River Package in MODFLOW. A steady state groundwater model was calibrated using 18 observation wells record for the year 2008. Computed contour show that the River Base flow was a part of the groundwater regime in the area. Results showed that the base flow is an important factor for modelling the groundwater in Anga'a River Watershed, and groundwater flows was controlled by the topography of the area.

Ahmed et al, (2009) used Visual MODFLOW, Pro 4.1 to evaluate water balance and study the flow system Yamuna–Krishni interfluve. The horizontal flows, seepage losses from unlined canals, recharge from rainfall and irrigation return flows were applied using different boundary packages available in the VISUAL MODFLOW. The river aquifer interaction was simulated using the river boundary package. Hydraulic conductivity values were applied to specific zones and these ranged from 9.8 to 26.6 m per day. Pumping rates of 500, 1000, 1500, 2000 and 2500 m³/day were applied to appropriate areas of the model to simulate areas of stress. The zone budget of the area shows a water balance deficit for the period from 2006 to 2007. The total recharge to the study area was 160.21 million m³ (MCM). The sensitivity analysis showed that the model was the most sensitive to hydraulic conductivity and recharge parameters.

Vijai et al, (2011) developed a conceptual model of groundwater for the Pali Area, India using VISULA MODFLOW. GIS was used for preparing out the hydrological, hydrogeological and geological data. The model was developed on spatial bases using GIS tool and not taking the average values for the whole area. Boundary conditions for the model was analysed and studied using GIS and a map created to import for the conceptual modelling showing various boundary condition in the area. GMS

(groundwater modelling system) is used with the GIS for the preparing the 3D visualization of the study area. Groundwater recharge values were used on spatial bases for the whole area for the better results.

Senthilkumar et al, (2011) developed a model using MODFLOW and GMS version 3.1 for finding the impact of square on the groundwater stream for the Palar River catchment, Tamilnadu, Southern India. Groundwater requirements of the atomic center were increasing and in order to fulfil the requirements of the atomic plant an underground barrier across the Palar river was proposed to increase the recharge to groundwater. The groundwater model utilized as a part of his study to check that groundwater levels might increase by about 0.1 to 0.3 m broadening out a separation of about 1.5 to 2 km from the upstream side of the control, while on the downstream side, the groundwater head was lower by about 0.1 to 0.2 m. He finds out that the model also required that with the subsurface control set up the supplemental groundwater crucial component of essentially 13,600-m³/day may be met with least lessen in adjacent groundwater head.

Fethi et al, (2012) developed a groundwater flow model in Zeramdine- Beni Hassen Miociene Aquifer system (east central Tunisia). The model was developed to validate the groundwater properties deduced from geological, geophysics and hydrodynamic studies were done in the region using the MODFLOW 2000 and GIS. A 3D model was developed for ZBH aquifer using previous data. The model was calibrated and validated for the year 1980-2007. The results showed that aquifer shows susceptibility to change in infiltration and hydraulic conductivity values. The model results showed a good agreement with the observed data. The model could be used further for the area similar to ZBH aquifer characteristics and showed a good result in arid to semi-arid region.

Hussein et al, (2013) developed a numerical model using VISUAL MODFLOW for the Baoyang irrigation area and aquifer is considered a single layer. Model was run to check the sustainability of groundwater and to develop a suitable plan for the future groundwater use. Conceptual model was built by using Aquifer properties, pumping and observation data and model was run for 2006-2009 for calibration and validation, and then model was extrapolated to check the effect of groundwater level for future prediction up to year 2048. Model was calibrated using Parameter estimation Package (PEST) and water balance for the area is checked to find out the difference between the water entering and leaving the system. It was found that model results are very close to that of observed values of the observation wells and the model results showed that there is a slight fall in groundwater level for the Baoyang irrigation area.

Mohanty et al, (2013) developed groundwater flow model for khatajudi surua inter basin of Odisha, India. MODFLOW was used for the modelling; model was calibrated using weekly data of water level for (Feb 2004 to May 2006) and validated for the 1 year (June 2006- May 2007). 18 Observation wells were used for the calibration purpose. Calibration of the groundwater model was done by hit and trial method. Combined with automatic PEST with a mean RMSE of 0.62 and NSE value of 0.915. The inputs for the Artificial neural network were rainfall, ET, River stage, pumping rates and groundwater levels of these wells in previous time. It was found that ANN provide better results in the study area than the numerical model of short time period over the area.

Mirudhula (2014) developed a groundwater model using VISUAL MODFLOW to check the impact of Lined / Unlined Canal on groundwater recharge for the Bhavani Basin. Groundwater recharge was calculated using groundwater flow modelling using visual mudflow. Three major natural sources of groundwater recharge were considered

for the study consisting of rainfall infiltration, seepage from canals and return flow from irrigation. Zone budget of the area was carried out to study the effect on groundwater recharge from lined/unlined canals. Results show that unlined canal recharge 20% more to groundwater than the lined canal.

Pradeep et al, (2014) developed a steady state finite model using VISUAL MODFLOW to quantify groundwater in Choutuppal Mandal, Nalgonda (Dt) AP. Used groundwater well data from 19 wells and found that the groundwater model gave good results because computed values have a good fitness. The groundwater budget was calculated using zone budget for the entire watershed using the groundwater flow model. It was determined that groundwater enters from Borrollagudem and leaves at Aregudem and Katrevu through Choutuppal. The output indicated that there was a very little amount of storage. This necessitates immediate arrangement for recharge of groundwater by all possible means to save the groundwater for future usage.

Vladimir et al, (2015) developed a groundwater model for the Kratal Agriculture Area to assess the artificial recharge through the infiltration pools near the Kratal river. The aim was to reduce the groundwater shortage. Outcomes showed that maximum recharge from the infiltration pool can be exceeded to 1000 m³/day, in response to an infiltration rate of 0.2 m/day. The radius of the infiltration pool should not exceed from 500 m for better efficiency. A spatial-distribution groundwater-flow model such as MODFLOW could provide reliable information for the planning of artificial groundwater recharge from infiltration pools under complex hydrogeological and intensive farming conditions.

Hong et al, (2015) studied the effect of rainfall intensity on groundwater recharge by using a 3D groundwater finite difference model in MODFLOW. A 3D finite difference groundwater model was calibrated and validated using MODFLOW and simulated

rainfall was input to check the effect of the rainfall. Groundwater recharge coefficient was calculated using PES-ASP Package of MODFLOW. The value of groundwater recharge coefficient was reduced from 0.439 to 0.076 for increasing the rainfall intensity from 45 mm/hr to 120 mm/hr respectively. He concluded that rainfall intensity has great effect on groundwater regime, water shed management and river flows.

Maheswaran et al, (2016) developed groundwater model for Ganga River Basin. He used Visual MODFLOW covering an area of 1 million km² and run as transient model. The model domain was selected of 2500x2500m consisting of 500x500 rows and columns respectively. The study includes historical groundwater data from the concerned departments of the India and incorporate major tributaries of River Ganga as geo hydrological data was obtained from central groundwater board India. Model was run for 4.5 year and was calibrated using parameter Estimation (PEST) simulations. Analysis of stream network with zone budget helped to decide that the stream is gaining or losing in the study area. The study was helpful to facilitate development of detailed impact scenarios for groundwater contribution in the area.

Ambe et al, (2017) studied recharge through rainfall in north region of Cameron. He used water balance model GROWA for the assessment of recharge in north Cameron. For the input parameter of the GROWA, a world soil map of a scale 1: 3000,000 was used. The model results showed a good agreement between the computed values in model and observed values of recharge in the study area. It was found that GROWA model present good results for the computation of reliable recharge values for rainfall even the input parameters were extracted from a small-scale international soil map. The model would give more accurate results if high-resolution soil data is available for the calculation of groundwater recharge.

Alemu et al, (2017) used wetspa model to study the temporal and spatial variation in recharge of Geba Basin, Northern Ethiopia. The model was developed with covering an area of 4249 Km² and incorporate elevation, soil and land use data, hydro meteorological and river discharge data. He simulated wetspa for 4 years on daily basis (Jan 1, 2000- Dec 18, 2003). The model showed a good agreement between observed and simulated stream flow hydrographs. The water balance results showed a good value for both calibrated and validated period. He found that mean annual recharge is mostly dependent on rainfall amount followed by soil and land cover of the area. He computed that 21% of the geba basin has recharge more than 120 mm and 1% of the area has recharge less than 5 mm.

In previous studies different researches used different models for the groundwater recharge computation. The best model that was recommended by most researchers was Visual MODFLOW for the groundwater computation. In my study a dam is proposed at 3mile U/S of the Mohlanwall in River Ravi and then the effect of the proposed dam is studied in HEC-RAS & Visual MODFLOW for backwater effect and recharge computation for the Lahore city.

CHAPTER III

DATA COLLECTION AND ANALYSIS

3.1 Data Collection

- River Ravi X-section from Ravi Siphon to Head Balloki was collected from River Survey Division of Irrigation Department Lahore.
- Flow data for the river Ravi at Ravi Siphon and Shahdra gauge station was collected from the Chakbandi Division, Irrigation Department Lahore.
- Tube wells discharge and location data was obtained from WASA Lahore and Mott McDonald Pakistan (MMP) Consulting firm.
- Terrain Data of Lahore was obtained from Mott McDonald Pakistan; and
- Hydraulic conductivity data was extracted from the map developed by NESPAK for the past Lahore groundwater study.

3.2 Makesens Analysis

Mann Kendall test and Sens slope method were used to carry out the trend analysis and slope of the given data on yearly basis. Mann Kendal method gives the trend line for the data that how much the data deviate from its mean value and Sens method was used to check the slope of data.

Mann Kendall test requires minimum four values to study / analyse the non-parametric trend of the data. The confidence interval for Sens slope method was used to check the linear trend between the data. it requires a minimum of 10 values for the trend analysis.

3.3 Trend Analysis

3.3.1 Discharge Trend

Discharge trend was analysed using Makesens analysis for a period of 15-year data of discharge 2000-2014, Makesens gives trend of annual based daily data. Graphs were

drawn for different time periods such as January to December, summer, winter season separately to determine that how the flow is available in river Ravi. The discharge yearly trend is shown in **Figure 3.1** and remaining are shown in Annexure B.

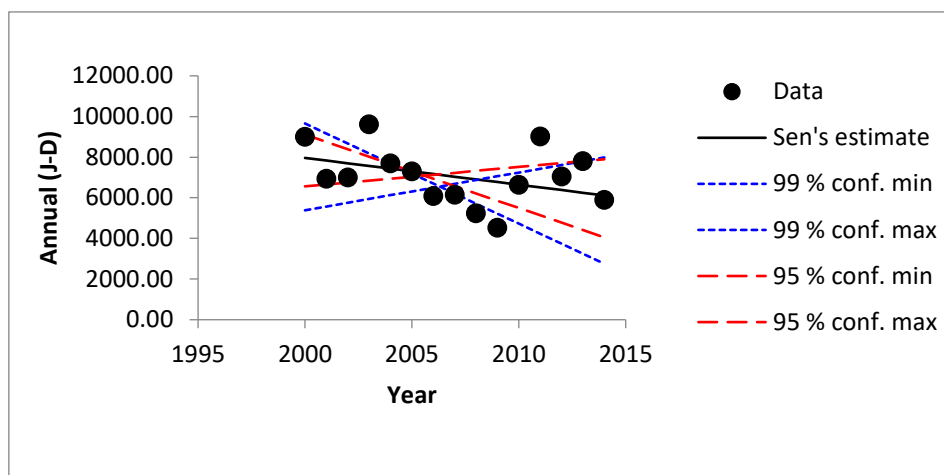


Figure 3.1: Annual Average Daily flow for the year 2000-2014 of River Ravi

The Sen's estimate for discharge values and for 99% confidence interval was observed by using Sen's estimate and a discharge data for the 15 year was used for this purpose the average difference between the actual daily discharge and Sen's estimate is almost 1200 cusec as shown in **Table 3.1**.

Table 3.1: Daily average discharge of river Ravi estimated by Sen's method

| Year | Data Cusic | Sen's estimate | 99 % confidence min | 99 % confidence max | 95 % confidence min | 95 % confidence max | Residual |
|------|------------|----------------|---------------------|---------------------|---------------------|---------------------|----------|
| 2000 | 9006.81 | 7960.82 | 9657.37 | 5384.67 | 9112.59 | 6562.68 | 1045.99 |
| 2001 | 6934.76 | 7829.06 | 9164.43 | 5570.36 | 8750.94 | 6657.76 | -894.30 |
| 2002 | 7000.61 | 7697.30 | 8671.48 | 5756.06 | 8389.29 | 6752.83 | -696.69 |
| 2003 | 9620.11 | 7565.55 | 8178.53 | 5941.75 | 8027.64 | 6847.91 | 2054.56 |
| 2004 | 7685.59 | 7433.79 | 7685.59 | 6127.44 | 7665.99 | 6942.99 | 251.80 |
| 2005 | 7304.34 | 7302.03 | 7192.64 | 6313.13 | 7304.34 | 7038.06 | 2.30 |
| 2006 | 6079.99 | 7170.27 | 6699.69 | 6498.83 | 6942.69 | 7133.14 | 1090.28 |
| 2007 | 6144.21 | 7038.52 | 6206.75 | 6684.52 | 6581.04 | 7228.21 | -894.30 |
| 2008 | 5235.48 | 6906.76 | 5713.80 | 6870.21 | 6219.38 | 7323.29 | 1671.28 |
| 2009 | 4528.73 | 6775.00 | 5220.85 | 7055.90 | 5857.73 | 7418.37 | 2246.27 |
| 2010 | 6643.24 | 6643.24 | 4727.91 | 7241.60 | 5496.08 | 7513.44 | 0.00 |
| 2011 | 9019.40 | 6511.49 | 4234.96 | 7427.29 | 5134.43 | 7608.52 | 2507.91 |
| 2012 | 7045.65 | 6379.73 | 3742.01 | 7612.98 | 4772.78 | 7703.60 | 665.92 |
| 2013 | 7798.67 | 6247.97 | 3249.07 | 7798.67 | 4411.13 | 7798.67 | 1550.70 |
| 2014 | 5900.33 | 6116.21 | 2756.12 | 7984.37 | 4049.48 | 7893.75 | -215.88 |

3.3.2 Rainfall Analysis

The rainfall analysis by Sen's estimated was done and the average daily rainfall on annual basis is between 1-3mm for the year between 2000-2014, and the graphs for 99% confidence interval as well as 95% confidence interval is shown in **Figure 3.2**

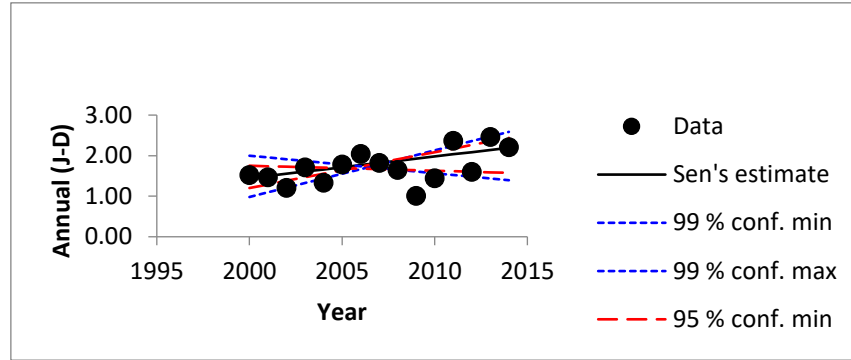


Figure 3.2: Average annual dialy rainfall for year 2000-2014

The detailed values for rainfall calculated using sen's analysis is shown in **Table 3.2** which shows that sen's estimated rainfall value is almost near to the average daily rainfall of the Lahore. 95% confidence interval was calculated by sen's method also which is accurate 95% for the given data.

Table 3.2: Average daily rainfall of Lahore city estimated by Sen's method

| Year | Data mm | Sen's estimate | 99 % confidence min | 99 % confidence max | 95 % confidence min | 95 % confidence max | Residual |
|------|---------|----------------|---------------------|---------------------|---------------------|---------------------|----------|
| 2000 | 1.52 | 1.45 | 2.00 | 0.98 | 1.75 | 1.20 | 0.07 |
| 2001 | 1.47 | 1.50 | 1.95 | 1.09 | 1.74 | 1.29 | -0.03 |
| 2002 | 1.21 | 1.55 | 1.91 | 1.21 | 1.73 | 1.38 | -0.34 |
| 2003 | 1.72 | 1.61 | 1.87 | 1.32 | 1.72 | 1.47 | 0.11 |
| 2004 | 1.34 | 1.66 | 1.82 | 1.44 | 1.70 | 1.56 | -0.32 |
| 2005 | 1.79 | 1.71 | 1.78 | 1.55 | 1.69 | 1.64 | 0.07 |
| 2006 | 2.04 | 1.77 | 1.74 | 1.67 | 1.68 | 1.73 | 0.28 |
| 2007 | 1.82 | 1.82 | 1.70 | 1.78 | 1.67 | 1.82 | 0.00 |
| 2008 | 1.65 | 1.87 | 1.65 | 1.90 | 1.65 | 1.91 | -0.22 |
| 2009 | 1.01 | 1.93 | 1.61 | 2.01 | 1.64 | 2.00 | -0.91 |
| 2010 | 1.44 | 1.98 | 1.57 | 2.13 | 1.63 | 2.08 | -0.54 |
| 2011 | 2.37 | 2.03 | 1.52 | 2.24 | 1.62 | 2.17 | 0.34 |
| 2012 | 1.60 | 2.09 | 1.48 | 2.36 | 1.60 | 2.26 | -0.48 |
| 2013 | 2.46 | 2.14 | 1.44 | 2.47 | 1.59 | 2.35 | 0.32 |
| 2014 | 2.21 | 2.19 | 1.39 | 2.59 | 1.58 | 2.44 | 0.02 |

Significancy of the rainfall data was checked by mann kendall test. Mann kendall uses z test for significacy for different scenario, for annual daily trend and summer season the rainfall trend significantly increase while for the autom the trend was also significant change but for the months of winter and summer the trend is not significant for the daily rainfall of the Lahore ciyt as shown in **Table 3.3**.

Table 3.3: Mann Kendall trend of Rainfall for Lahore city

| TREND STATISTICS | | | | | |
|----------------------------------|-------------------|------------------|----------|---------------------------|----------------------|
| FI17 Virolahti 1987- 2000 | | | | | |
| | | | | Mann-Kendall trend | |
| Time series | First year | Last Year | n | Test Z | Significance. |
| Annual (J-D) | 2000 | 2014 | 15 | 1.68 | + |
| Winter (O-M) | 2000 | 2014 | 15 | 0.20 | |
| Summer (A-S) | 2000 | 2014 | 15 | 1.78 | + |
| Winter(D-F) | 2000 | 2014 | 15 | -0.10 | |
| Spring (M-M) | 2000 | 2014 | 15 | -0.20 | |
| Summer (J-A) | 2000 | 2014 | 15 | -0.10 | |
| Autumn (S-N) | 2000 | 2014 | 15 | 2.18 | * |

CHAPTER IV

METHODOLOGY

4.1 HEC-RAS Model

HEC-RAS model was run for the four Scenarios mentioned below on the selected reach of the River Ravi for the computation of water surface profile.

1. Steady state condition;
2. Unsteady state condition using average flow hydrograph for (2000-2014);
3. Unsteady state condition with dam and 90% flow within reach throughout the time (2000-2014); and
4. Unsteady state condition with dam and low flow condition (2000-2014).

4.1.1 HEC-RAS Geometry Input

Geometry data is the main input for the hydraulic model for water surface profile computation. HEC-RAS geometry data was obtained from Irrigation Department and was input manually in HEC-RAS model. X-sections obtained from River Survey Division, Irrigation Department was at the interval of 2000-3000 ft interval depending on the river pattern. A River Reach of 38 mile was drawn on the terrain for this study manually and then x-section is input manually as shown in **Figure 4.1**.

Terrain was plotted using HEC-RAS mapper and use a DEM of 1m obtained from Mott McDonald Pakistan (MMP). The value of manning's n was obtained from the literature and used as input in the model for its computation.

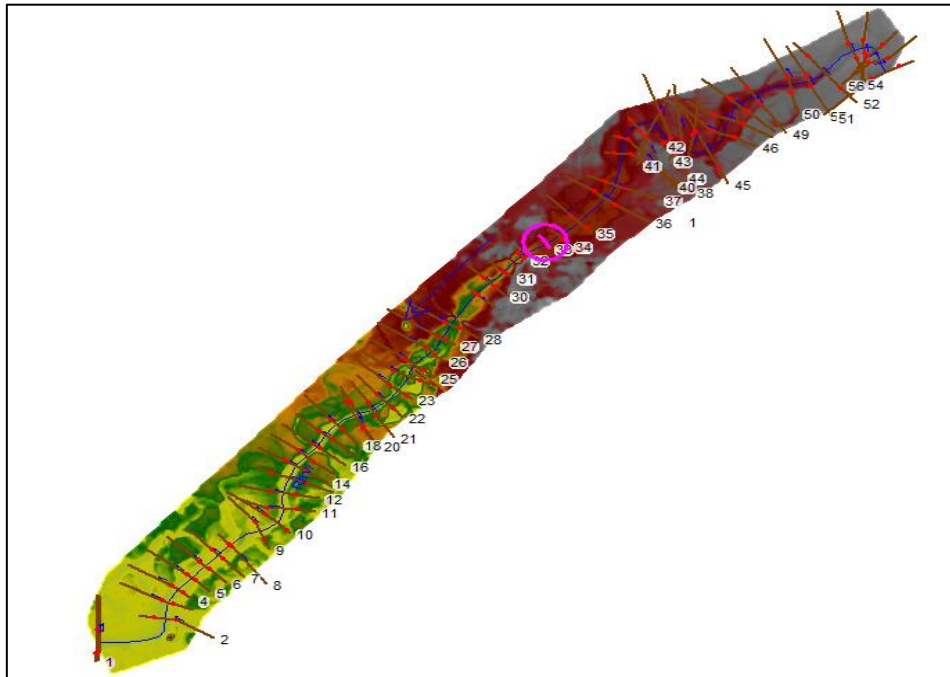


Figure 4.1: X-Section of the river Ravi with terrain for river Ravi study area

River Ravi x-sections were obtained from the Irrigation department and digitized in Excel. River Section at 4500 ft D/S of the Indian boarder is shown in Figure 4.2. River Ravi X- Section of reach from the Shahdra to head Balloki are given in the Annexures C & D respectively.

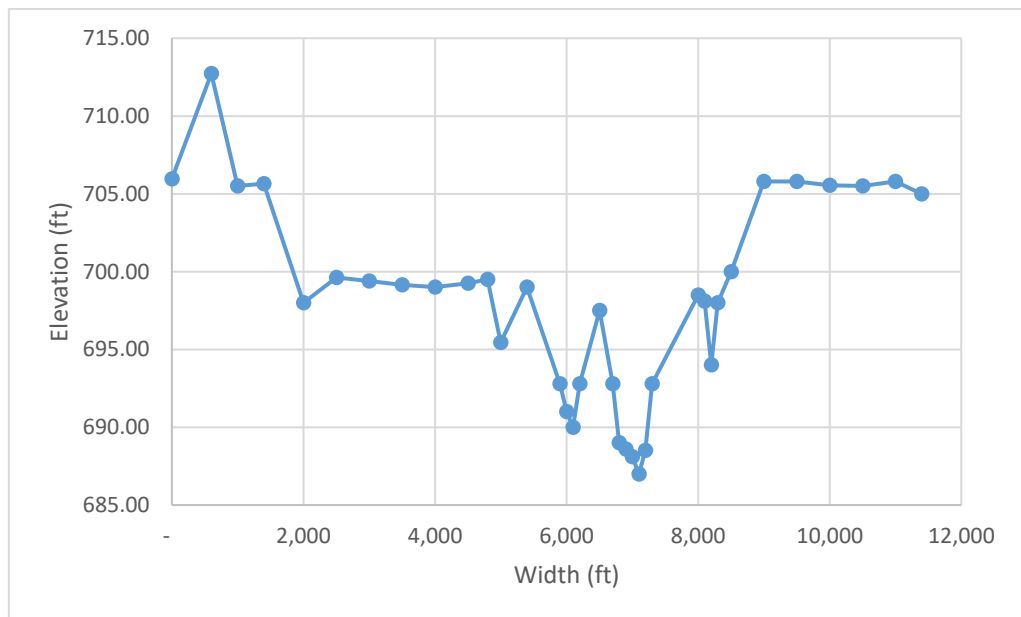


Figure 4.2: River Ravi X-Section at 4500ft D/S of Indian Boarder

4.1.2 Flow Data Input

Flow Data for the River Ravi from 1991-2014 was obtained from Irrigation Department.

The average daily flow of the River Ravi was 6967 cusecs for the year 2000-2014 at shahdra gauge station. The max flow of 314321 cusec was observed for the year 2010 due to peak flood.

Frequency analysis was done on the flow data to calculate the flow at different return period in River Ravi in case when the dam is to be installed on River Ravi for the Pondage that can help for the groundwater recharge.

4.1.3 Boundary Condition

Normal depth was used as boundary condition for steady flow analysis by inputting the slope of channel as input in the hydraulic model.

The model was simulated for the Steady state flow analysis for the calculation of the water surface srofile for the River Ravi. For steady state, boundary condition normal depth was used as a reference in downstream of the channel as shown in **Figure 4.3**.

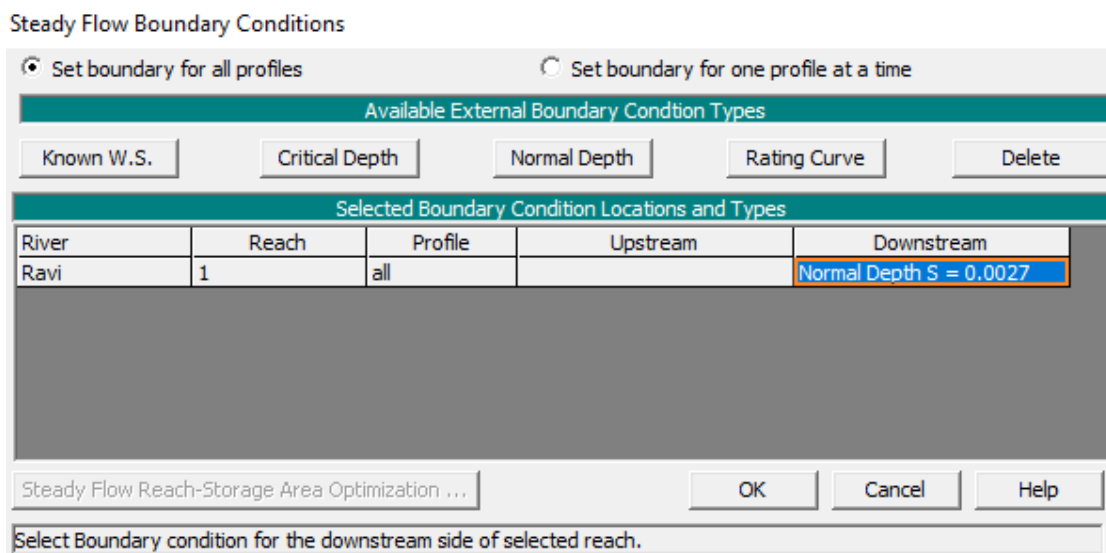


Figure 4.3: Steady flow boundary condition for River Ravi

4.2 HEC-RAS Unsteady Flow

HEC-RAS was run on the unsteady flow for the analysis of dam, for the unsteady flow the boundary conditions are changed. In upstream boundary condition flow Hydrograph and for downstream boundary condition Normal depth was used for the computation.

4.2.1 Boundary Condition

The U/s boundary condition for unsteady was used as flow hydrograph. The model was run for the unsteady flow computation for the 90% available flow in the river as well as on the lowest flow in the river for computation.

Monthly flow hydrograph was developed for the easy computation of the model and then input in the Boundary condition for the easy computation. Flow hydrograph is shown in **Figure 4.4**.

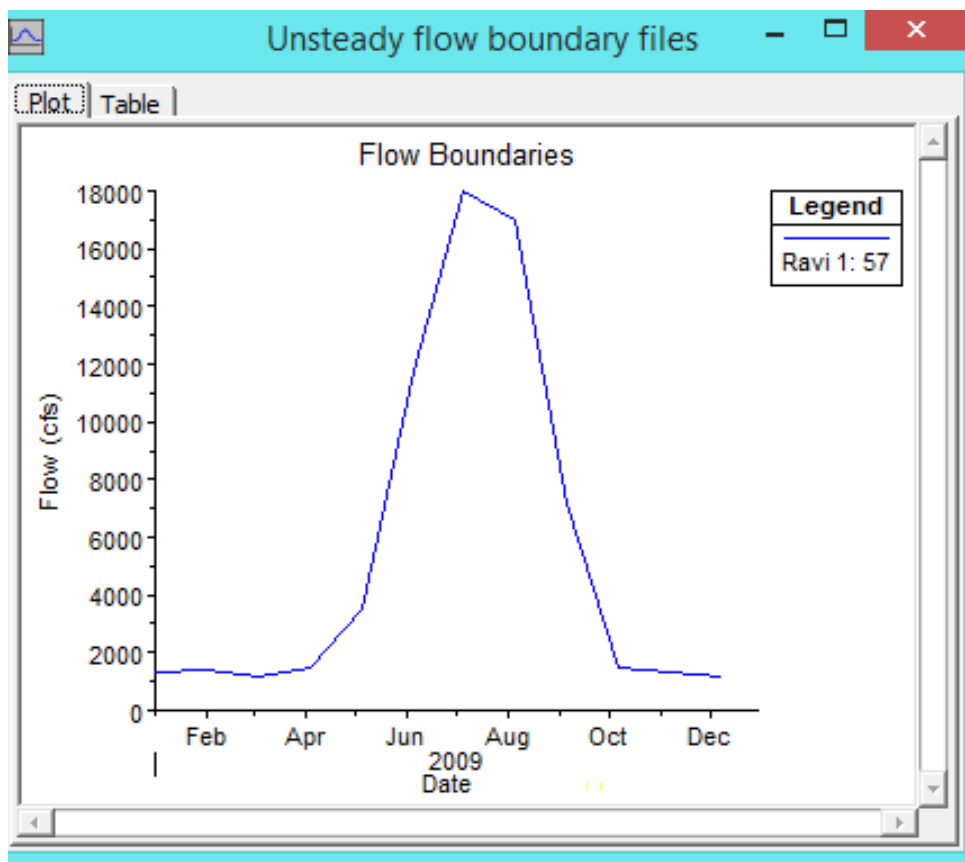


Figure 4.4: Avg. flow hydrograph for river Ravi

For unsteady flow, boundary condition was selected at U/S and D/S of the river channel. At U/S monthly flow hydrograph was used as input, while on D/S normal depth was used for which the slope of the channel was used to calculate normal depth.

Boundary conditions for the unsteady flow are shown in **Figure 4.5**.

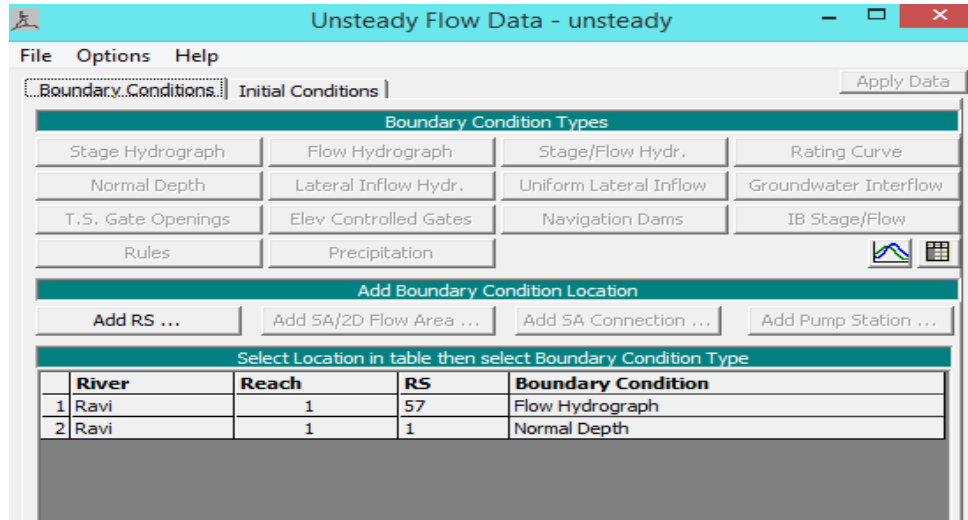


Figure 4.5: Unsteady flow boundary condition for river Ravi

4.3 Unsteady Flow with dam

HEC RAS model was run with dam installed in U/S of Mohlanwal site to check the effect of dam installed. The dam head was selected as 10ft and slope of the river was 0.00027 (1.4ft / mile). The water surface profile was computed on the basis of unsteady flow. Broad crested dam was designed with a dam coefficient of 2.6. as shown in **Figures 4.6 & Figure 4.7** respectively.

Inline Structure Weir Station Elevation Editor

| Distance | Width | Weir Coef |
|----------|-------|-----------|
| 1000 | 500 | 2.6 |

Edit Station and Elevation coordinates

| # | Station | Elevation |
|---|---------|-----------|
| 1 | 0. | 665. |
| 2 | 1000. | 665. |
| 3 | 11000. | 665. |
| 4 | 12500. | 665. |
| 5 | | |
| 6 | | |
| 7 | | |
| 8 | | |

U.S Embankment SS

 D.S Embankment SS

Weir Data

Weir Crest Shape

Broad Crested
 Ogee

Enter distance between upstream cross section and deck/roadway. (ft)

Figure 4.6: Proposed broad crested dam design and Elevation:

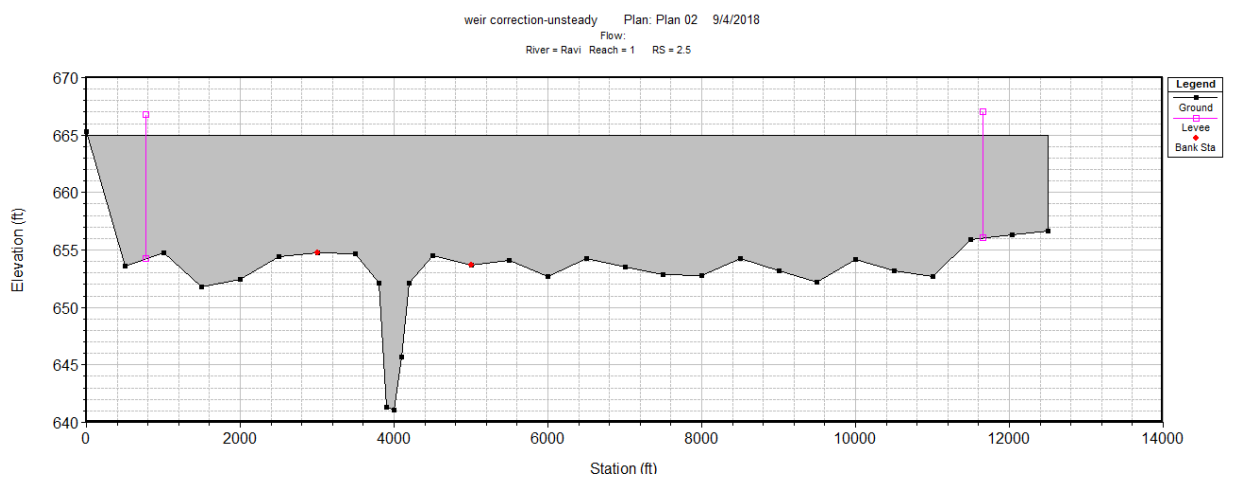


Figure 4.7: Proposed dam section in HEC-RAS

4.4 VISUAL MODFLOW

The Visual MODFLOW software (2011.1 ver.6.0.159), developed by Schlumberger Water Services (SWS), Canada, was used for the present study. The Visual MODFLOW is the most complete and easy-to-use modelling environment for practical applications in three-dimensional groundwater flow and contaminant transport simulations.

Visual MODFLOW was used to groundwater model for the Lahore Aquifer. The packages used for the groundwater model of Lahore city are shown in **Table 4.1**.

Table 4.1: Visual MODFLOW Packages

| Name of Package | Abbreviations | Function |
|----------------------------------|----------------------|--|
| Basic | BAS | Specifies the grid, time step, etc. |
| Block-centered flow | BCF | Calculates the terms in the finite-difference equations, which determine the flow between adjacent cells. It also computes the movement of water into and out of aquifer storage |
| Well | WEL | Flows from or to wells |
| River | RIV | Flows to or from rivers or main canals via bed seepage |
| General Head Boundary | GHB | Flows to or from the model area from adjacent areas of the aquifer |
| Constant Head Boundary | CHD | The constant head boundary condition is used to fix the head value in selected grid cells regardless of the system conditions in the surrounding grid cells |
| Alternate Solver Packages | | |
| Strongly Implicit | SIP | Solves the finite-difference equations using the iterative SIP procedure |
| Slice-successive Over-relaxation | SOR | Solves the finite-difference equations using the iterative SOR procedure |

Groundwater flow modelling is a complex process, which need many inputs, and required some pre-requisites for the computation of the groundwater level. The groundwater flow modelling process chart is shown in **Figure 4.8**.

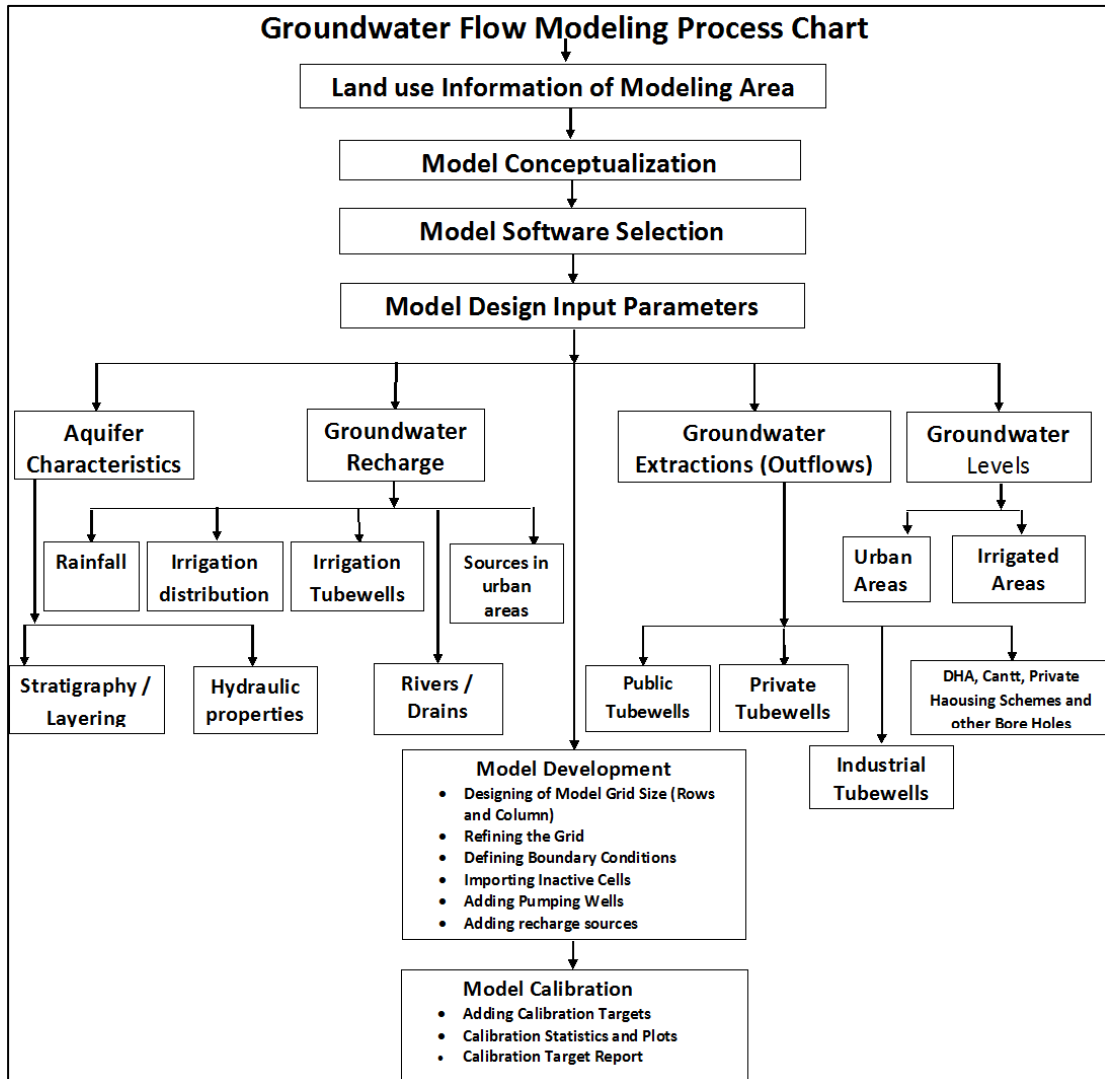


Figure 4.8: Visual MODFLOW Modelling Process flow chart for the study

4.4.1 Model Settings

After starting the model, first of all units for different parameters for the model and select the numeric model for the groundwater calculation are selected through the screen shown in **Figure 4.9**.

The screenshot shows a software interface for configuring a groundwater model. It is titled "Project Outline" and contains the following sections:

- Project Information:** Project Title: "Groundwater model", Description: (empty text box), and a "Details" button.
- Flow Simulation:**
 - Flow Type: Radio buttons for "Saturated (Constant Density)" (selected), "Saturated (Variable Density)", "Variably Saturated", and "Vapor".
 - Numeric Engine: "USGS MODFLOW 2000 from SWS" (dropdown menu).
 - Simulation Type: "Groundwater flow" (dropdown menu).
- Transport Simulation:**
 - Transport: Radio buttons for "No" (selected) and "Yes".
 - Numeric Engine: "NONE" (dropdown menu).
 - Description: (empty text box).
- Units:** A list of units for various parameters: Length (m), Time (day), Conductivity (m/d), Pumping Rate (m³/d), Recharge (mm/yr), Mass (kg), and Concentration (mg/L).

At the bottom of the dialog are four buttons: "< Back", "Next >", "Cancel", and "Help".

Figure 4.9: Model setting for the groundwater model of study area

4.4.2 Flow Settings

The next step for the model was to input the initial conditions to start the model, and define the simulation period for the model. Then it is selected whether model is to run steady state or transient condition. For this present study steady state analysis was selected for to make the computaion easy, because in transient condition variabillity of parameters over time is needed but in steady state parameteres remain constant for the whole period. The model was run for a period of 12 years and steady state condition. The remaining input parameters were assumed constant for the starting of the model and after that they were be changed according to the study area characteristics, as shown in **Figure 4.10**.

| Parameter Name | Value | Units |
|--------------------|-------|-------|
| Kx | 8.64 | m/d |
| Ky | 8.64 | m/d |
| Kz | 0.864 | m/d |
| Ss | 1E-5 | 1/m |
| Sy | 0.20 | |
| Eff. Por. | 0.15 | |
| Tot. Por. | 0.30 | |
| Recharge | 0 | mm/yr |
| Evapotranspiration | 0 | mm/yr |
| Extinction Depth | 0 | m |

Figure 4.10: Flow setting for model input of study area

4.4.3 Defining Model Domain

The next step is to define the model area. This model was developed for the Lahore district containing an area of almost 1772 Km². Model consist of 221x151 columns and rows respectively. The background map of the study area could be imported also in the Visual MODFLOW. The map is helpful to define the boundary conditions for the model. The aquifer was considered a single layer with a thickness of almost 400 m. The image imported in the background of the model was georeferenced in UTM coordinate system.

A site map was georeferenced according to the model domain and world coordinate systems forming a uniform grid of 221x151 columns and rows, respectively as shown in **Figure 4.11**.

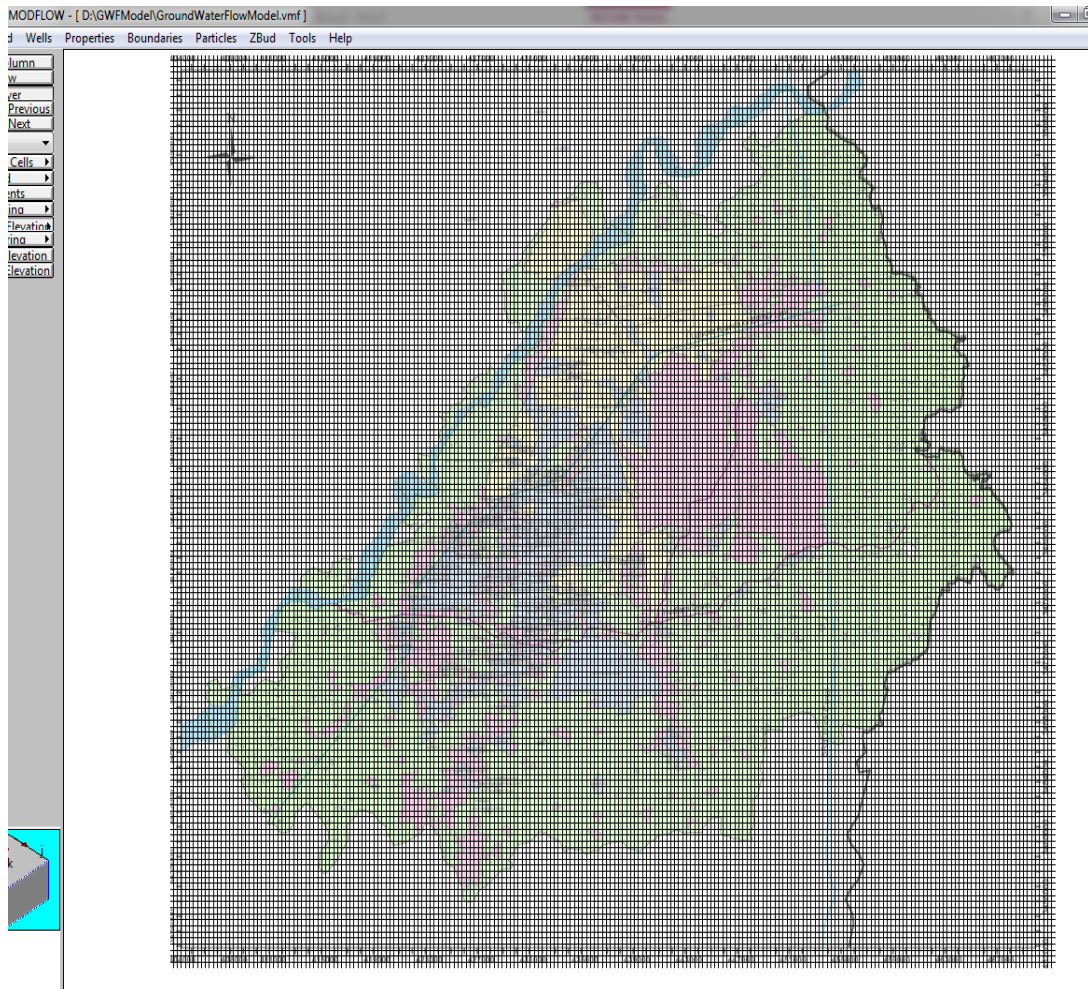


Figure 4.11: Defining model domain for study area in Visual Modflow

The grid could be refined according to the site condition in some points for the accuracy of model. The maximum number of rows and columns in VISUAL MODFLOW version used is 500x500.

4.4.4 Import Surface Elevation

Visual MODFLOW provides an improved set of tools for importing, creating, and editing and modifying layer elevation for the finite difference model grid. We can import data by using USGS DEM files to a specific number of grid cells. Visual MODFLOW supports importing of layer elevation files from a variety of data files including. The wells data, conductivity data and elevation data were imported from excel files and the interpolated using kriging method as shown in **Figure 4.12**.

- Points data (XYZ) ASCII files (.TXT, .ASC, .DAT);
- MS access database files (.MDB);
- MS excel files (.XLS);
- ESRI point files (.SHP);
- USGS DEM files (.DEM);
- Surfer grid files (.GRD); and
- ESRI grid files (.GRD).

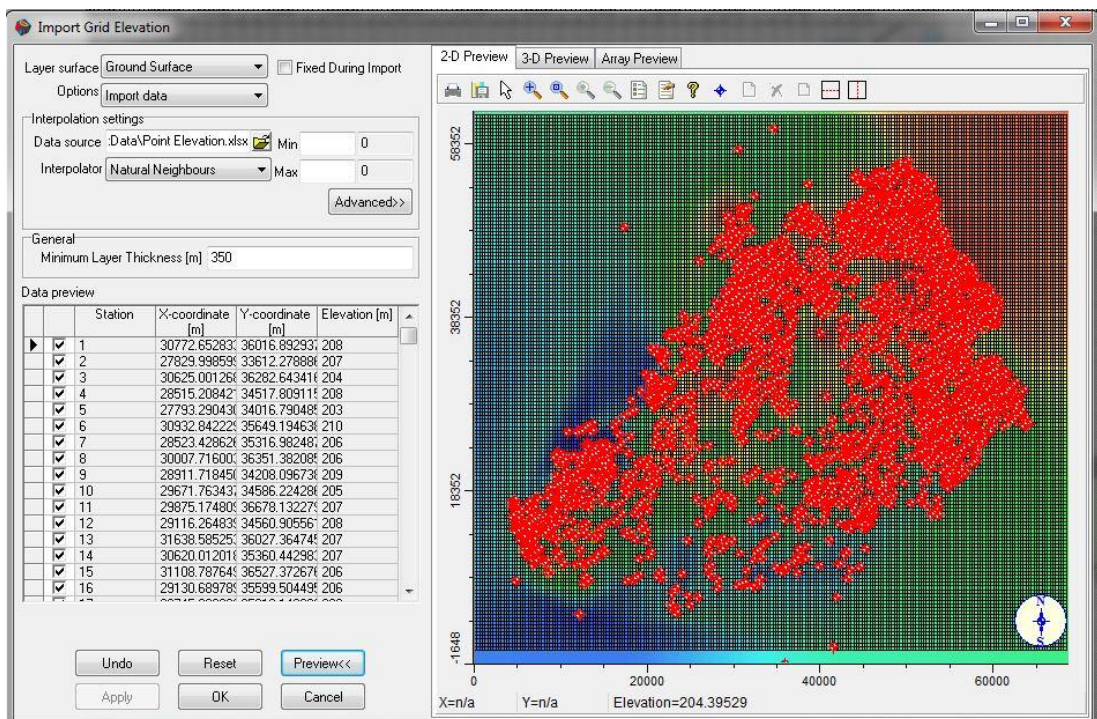


Figure 4.12: Importing surface elevation of the study area

4.4.5 Adding Pumping Wells

The pumping well data was obtained from WASA and MMP that was used for the model. About 484 WASA Tube wells and almost 2800 agricultural tube wells data was obtained and scheduling of these tube wells was done in Excel and then that excel sheet was imported in the model input. For agricultural tube wells the installed capacity is 1 Cusec, WASA, and non-WASA tube wells discharge capacity was varied 2-4 cusec and

they run for an average 10-18 hr a day. This type of the Excel sheet was prepared for agricultural and WASA tube wells and then it was imported in the VISUAL MODFLOW as shown in **Figure 4.13**.

Import Pumping Well Data

Line range
 All Use column names in line
 Lines:

Enter list of line numbers and line ranges separated by commas. For example: 1,3,5-12,23-END

Delimiters
 Space Comma
 Tab Slash
 User defined

Co-ordinate system
 World Model

Existing data
 Merge Replace

| Required Data | Match to column number... | Fill |
|----------------------------------|---------------------------|------|
| Well Name | 1 | |
| X co-ordinate [m] | 3 | |
| Y co-ordinate [m] | 4 | |
| Screen ID | 6 | |
| Top of Screen [m] | 9 | |
| Bottom of Screen [m] | 10 | |
| Stop Time [day] | 11 | |
| Pumping Rate [m ³ /d] | 15 | |

| Top Elevation(m) #13 | Bottom Elevation (m) #14 | Actual Pumpage #15 | Discharge #16 | Start Time #17 | End Time #18 | Column # 19 |
|----------------------|--------------------------|--------------------|---------------|----------------|--------------|-------------|
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 0 | 365 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 365 | 730 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 730 | 1095 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 1095 | 1460 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 1460 | 1825 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 1825 | 2190 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 2190 | 2555 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 2555 | 2920 | |
| 140.95 | 3.79 | -4566.93888 | 3.996926914 | 2920 | 3285 | |

Source Line No: 2

Help Next>> Cancel

Figure 4.13: Importing pumping wells data for study area

4.4.6 Adding Observation Wells

Observation wells are required to calibrate the groundwater model so that the relationship between calibrated and observed head is obtained. To import the observation wells an excel file shown in Annexure A was made and then imported in model. Thirteen observation wells were selected for the site and their excel file was

prepared to import in the model as shown in **Figure 4.14**. These observation wells were used to calibrate the groundwater model of the study Area.

Import Head Well Data

Line range
 All Use column names in line
 Lines:
 Enter list of line numbers and line ranges separated by commas. For example: 1,3,5-12,23-END

Delimiters
 Space Comma
 Tab Slash
 User defined

Co-ordinate system
 World Model

Existing data
 Merge Replace

| Required Data | Match to column number... | Fill |
|------------------------|---------------------------|------|
| Well Name | 1 | |
| X co-ordinate [m] | 2 | |
| Y co-ordinate [m] | 3 | |
| Screen ID | 4 | |
| Screen Elevation [m] | 5 | |
| Observation Time [day] | 6 | |
| Head [m] | 7 | |

| Well Name #1 | X [m] #2 | Y [m] #3 | Screen ID #4 | Screen Elev. [m] #5 | Obs. Time [day] #6 | HEAD [m] #7 |
|--------------|------------|------------|--------------|---------------------|--------------------|-------------|
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 365 | 195.938032 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 730 | 195.598032 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 1095 | 195.398032 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 1460 | 195.4468 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 1825 | 194.251984 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 2190 | 194.38 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 2555 | 194.178832 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 2920 | 194.050816 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 3285 | 193.874032 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 3650 | 193.58 |
| 11 | 437976.868 | 3470581.58 | 12 | 183.7 | 4015 | 193.31 |

Source Line No: 2

Help Next>> Cancel

Figure 4.14: Importing observation wells for study area

4.4.7 Conductivity Data

The conductivity map of the Lahore city was obtained from the NESPAK past study in which NESPAK developed conductivity zones for Lahore city with respect to conductivity. A cluster of conductivity values were obtained from these zones and then imported to Visual MODFLOW. After doing interpolation in the model a conductivity map was developed for the study area. The cluster of the conductivity data was

distributed with respect to latitude - longitude, and their conductivity values which imported to Visual MODFLOW and then, interpolated for the remaining area by using kriging method as shown in **Figure 4.15**.

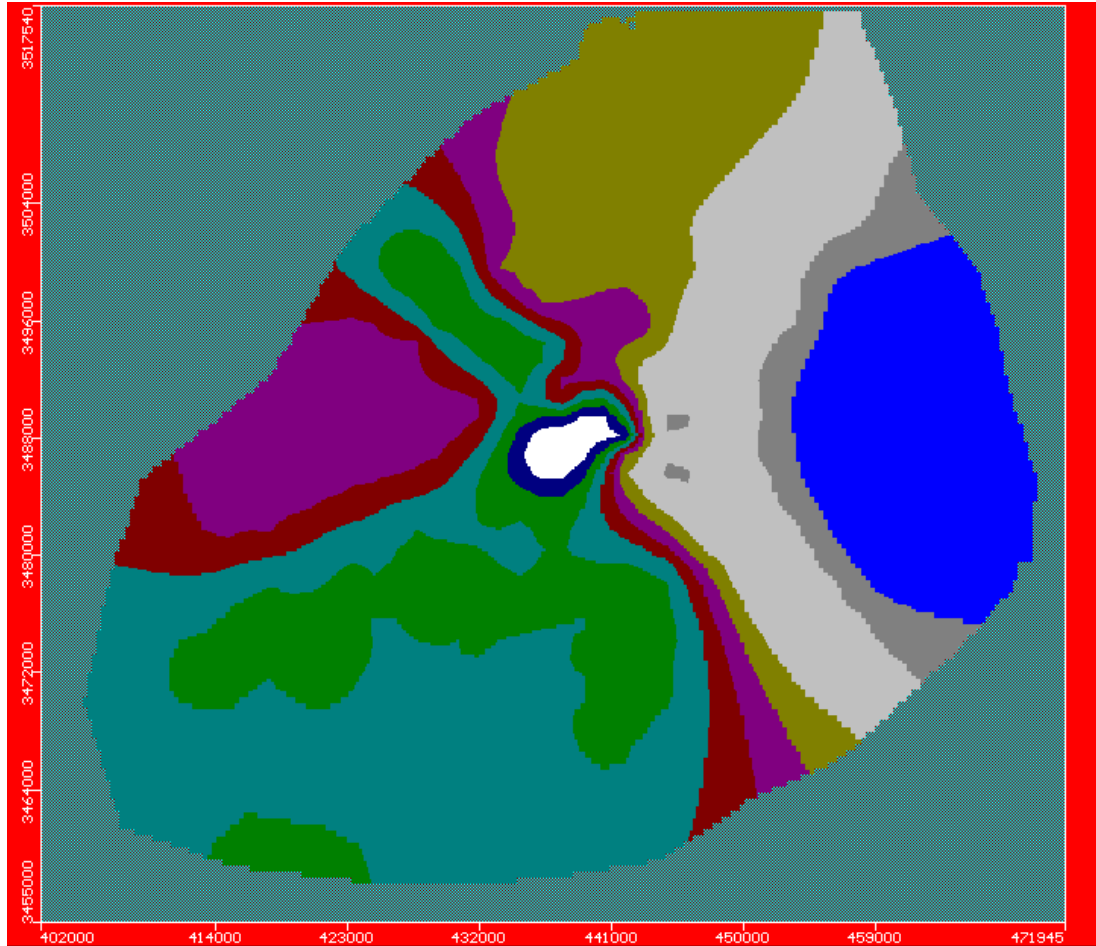


Figure 4.15: Conductivity map of the study area

4.4.8 Importing Initial Heads

Initial heads were observed from the observation wells. As observation wells record was collected for year 2001-2013. The heads obtained for the year 2001 were imported as initial head condition to run the model. The initial heads were imported and interpolated then to find out the initial head at remaining places as the initial head available were at the locations of observation wells only so remaining points were calculated with the help of kriging method with in the VISUAL MODFLOW model.

4.4.9 General Head Boundary

The function of the General-Head Boundary (GHB) Package is mathematically similar to that of the River, Drain, and ET Packages. Flow into or out of a cell from an external source is provided in proportion to the difference between the head in the cell and the reference head assigned to the external source. The application of this boundary condition is intended to be general, as indicated by its name, but the typical application of this boundary conditions is to represent heads in a model that are influenced by a large surface water body outside the model domain with a known water elevation. The purpose of using this boundary condition is to avoid unnecessarily extending the model domain outward to meet the element influencing the head in the model. As a result, the General Head boundary condition is usually assigned along the outside edges of the model domain. The boundary conditions for the study area is shown in **Figure 4.16**.

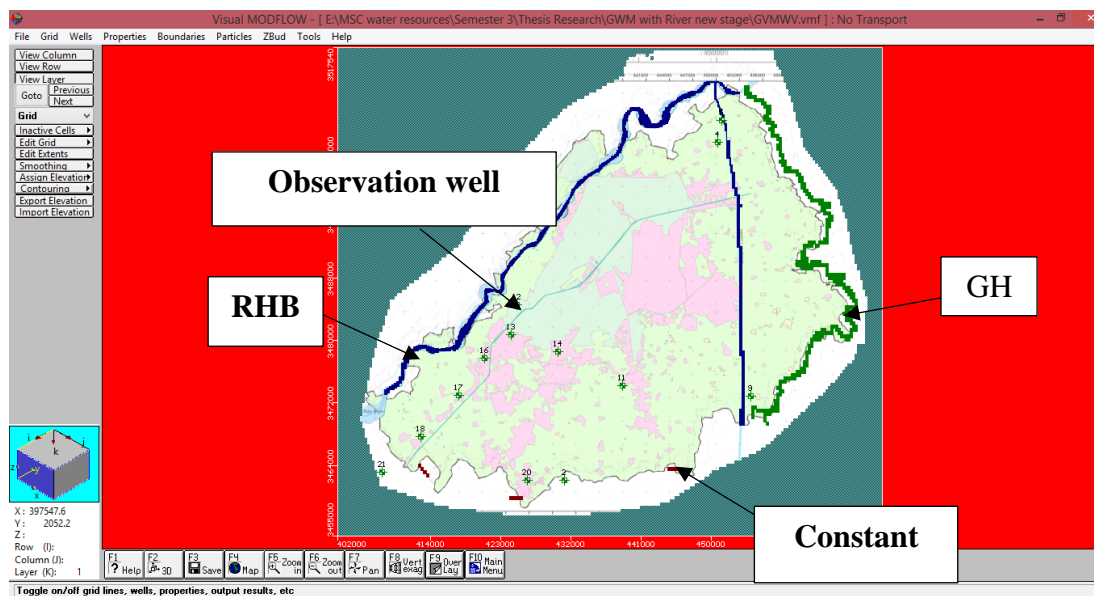


Figure 4.16: Boundary conditions in study area

4.4.10 Constant Head Boundary

Constant head boundary condition was assigned in the South East corner of the study area where head is fixed throughout the computation time as shown in **Figure 4.16**. The constant head boundary condition can be input as shown in **Figure 4.17**.

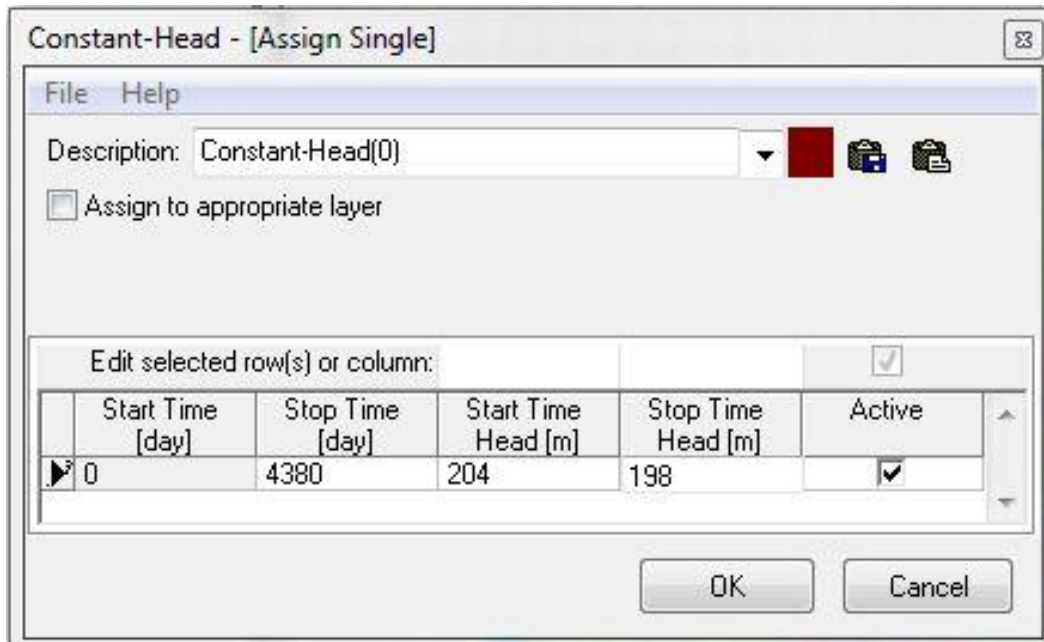


Figure 4.17: Constant head boundary condition for study area

4.4.11 River Head Boundary

River boundary condition was assigned in the study area for the River Ravi Reach and BRBD canal. Linear gradient method was selected to make the computation easier. It was assumed that river is flowing with a linear slope so a line was used for River input boundary condition where the input is required for the U/S and D/ S X-section of the River for the computation. Riverbed thickness of 1 m was used to make computation easier. The River data was input at starting point and ending point for linear gradient. River stage and river bottom were obtained from Ravi x-sections obtained from Irrigation Department. The input is shown in **Figure 4.18**.

File Help

Description: River(75)

Assign to appropriate layer

Use default conductance formula

Linear gradient

Start point End point

| Start Time [day] | Stop Time [day] | River Stage [m] | Riverbed Bottom [m] | Conductance [m ² /day] | Riverbed Thickness [m] | Riverbed Kz [m/d] | River Width [m] | Active |
|------------------|-----------------|-----------------|---------------------|-----------------------------------|------------------------|-------------------|-------------------------------------|--------|
| 0 | 4380 | 211.81 | 208.84 | (\$RCHLNG*\$RBW)/1 | 0.25 | 1000 | <input checked="" type="checkbox"/> | |

OK Cancel

Figure 4.18: River ravi boundary condition for the study area

River Ravi condition was input and then model was run to check the effect of its stage under normal condition. After the proposed dam simulation, the River Ravi stage was changed to check the effect on the groundwater recharge and groundwater level contours were simulated.

CHAPTER V

RESULTS AND DISCUSSIONS

5.1 Flow Calculation

5.1.1 Flow Duration Curve

Flow duration curve on daily basis for the year 2001 to 2013 was plotted to check the amount of water available in River Ravi for the whole year. Flow duration curves for the summer season is shown below in **Figure 5.1**, which shows that a minimum discharge of 1754 cusec is available even if no water is coming from the link canal (MRL) for the whole summer season.

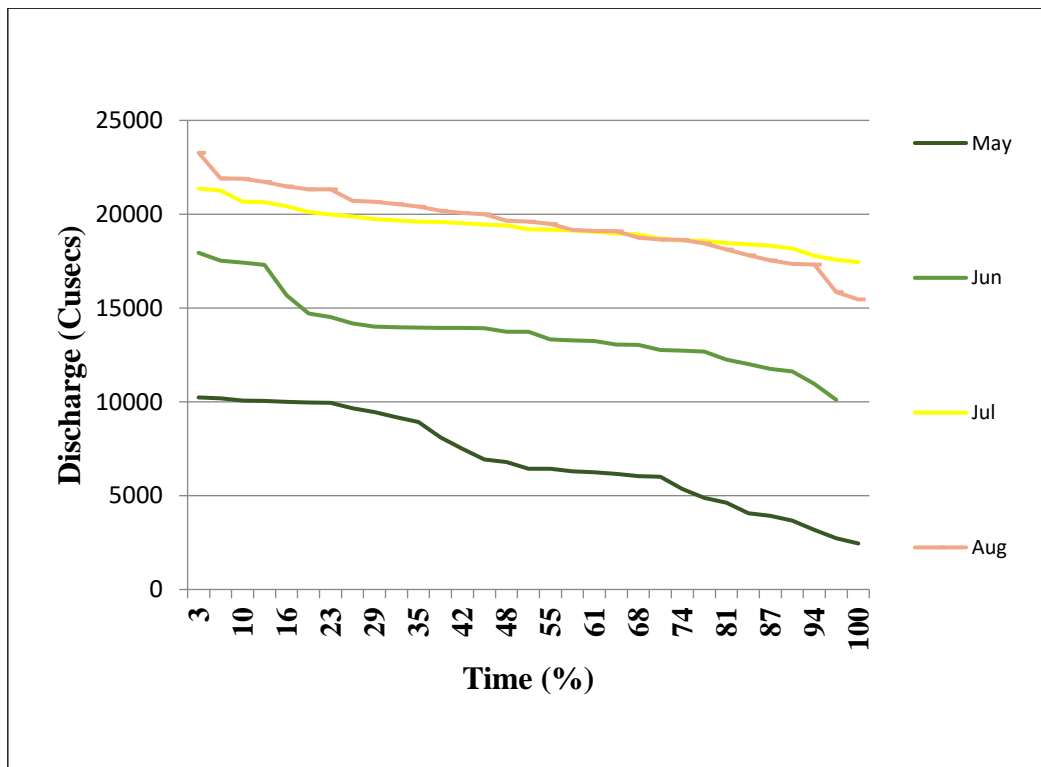


Figure 5.1: Avg. Flow Duration Curve of River Ravi for Peak months in year (2001-2013)

In addition, the graph shows the rise in discharge during the Jun-Sep, and then again reduces to a minimum value of 1754 cusec in December.

The average discharge for the year 2000-2014 on monthly basis is shown in Figure 5.2. The average flow in August is 22906 cusec and that for December is only 1754 cusec. This graph also shows the average flow available in the River Ravi for the whole year as shown in **Figure 5.2**.

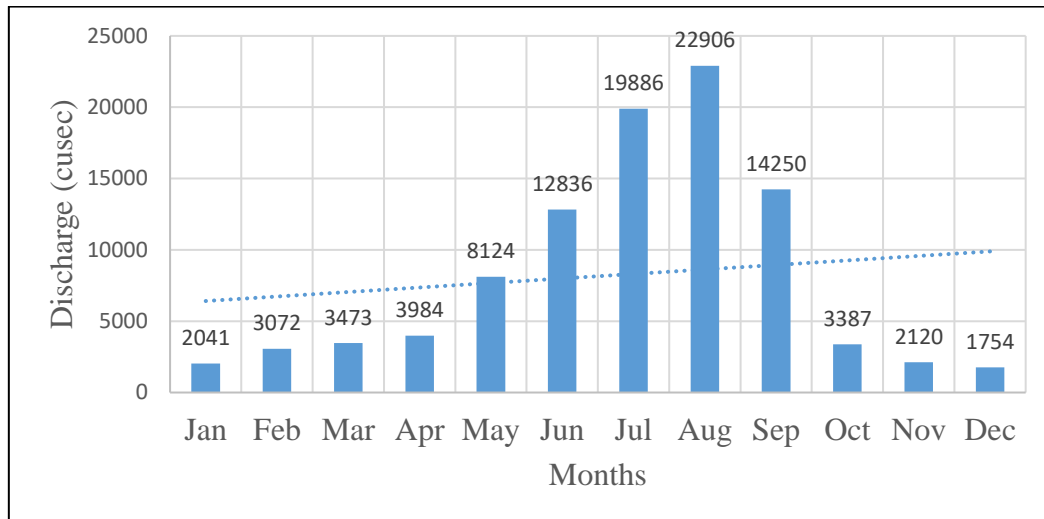


Figure 5.2: Average daily discharge in river Ravi for (2000-2014)

The flow analysis for the River Ravi was done and graphs for max, minimum and average discharge was plotted for better visualization of available data. The max discharge on the river Ravi on monthly basis is 23274 cusec in August and that of december is 1472 cusec as shown in **Figure 5.3**.

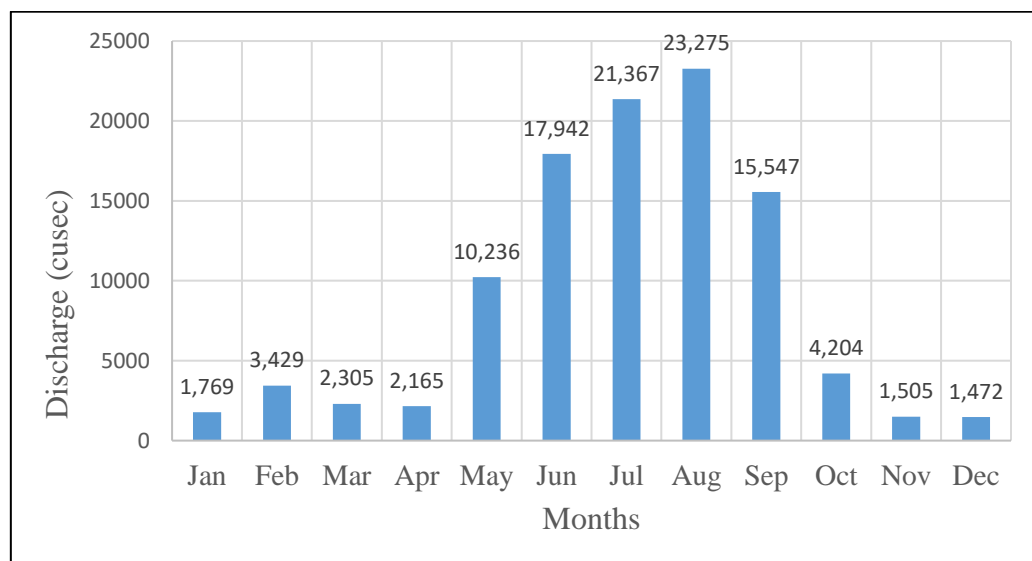


Figure 5.3: Max daily flow in monthly basis of river Ravi (2000-2014)

Flow analysis for the river Ravi was done and discharge available in river Ravi throughout the year was observed. Discharge at 90% of the time of whole year was calculated and it was found that a minimum value of 1200 cusec of water is available in the river Ravi throughout the year and a max discharge of 18000 cusec is available for the month of July in the river Ravi as shown in **Figure 5.4**.

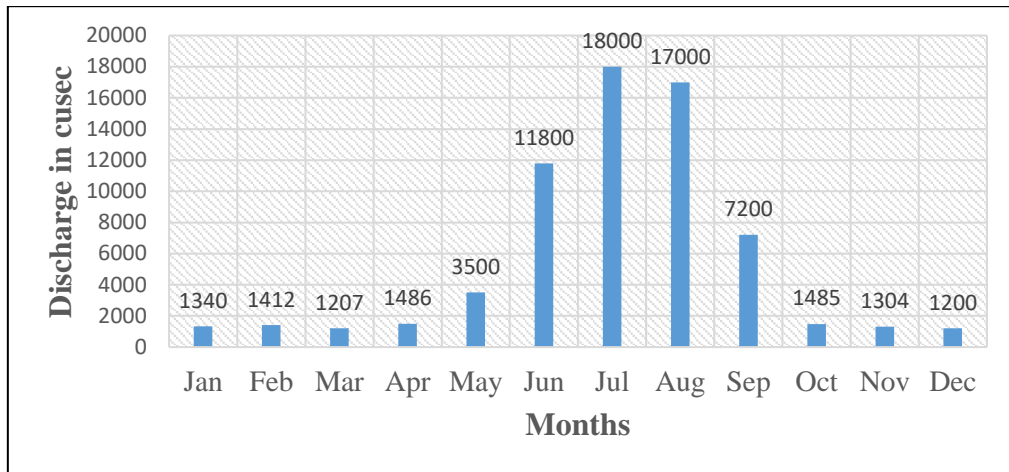


Figure 5.4: River Ravi discharge available 90% of the time for (2000-2014)

The minimum discharge of the river Ravi was also determined from the year 2000-2014. The discharge was determined for the Ravi Siphon and Shahadra. The minimum discharge of 200 cusec per day was observed for the month of December and similarly for moonsoon season a minimum discharge of 10000 cusec is observed in the month of July for the year 2000-2014 as shown in **Figure 5.5**.

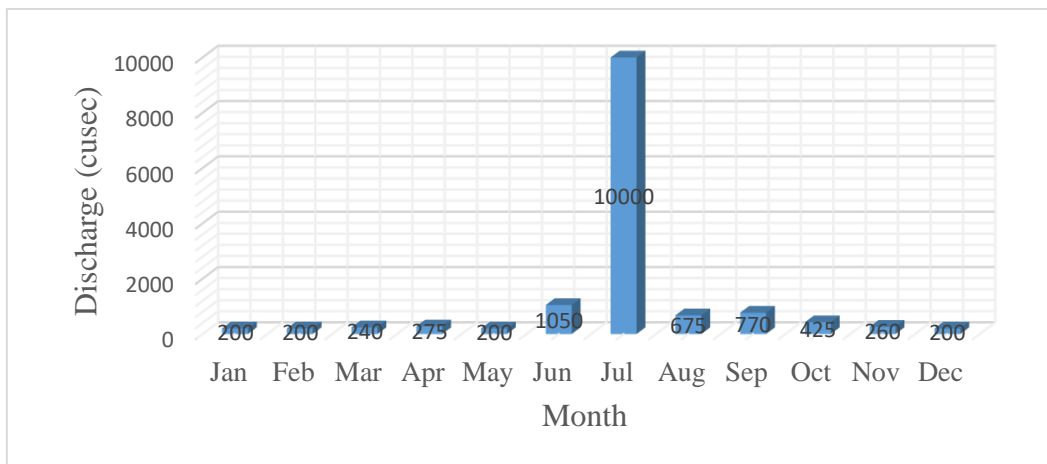


Figure 5.5: Minimum discharge throughout the year (2000-2014)

5.1.2 Frequency Analysis

Frequency analysis was used to calculate the peak discharge at different return period for the model input. The methods used for the frequency analysis was normal distribution, log normal distribution, extreme value type I, Frequency Factor, Log Pearson III and Gumbel Distribution. In Pakistan Gumbel, method is used mostly for the computation of the Probabilistic max flow or rainfall in different return period based on the structure designed. The Gumbel method shows a max discharge of 194902.31 cusec for a return period of 100-yr. while for the normal distribution it was 159676.12 cusec. As shown in **Table 5.1**.

Table 5.1: Frequency Analysis of Discharge data of River Ravi

| Sr. No. | Time period | Normal Distribution | Log Normal Distribution | Extreme Value Type I | Frequency Factor | Log Pearson III | Gumbel |
|---------|-------------|---------------------|-------------------------|----------------------|------------------|-----------------|--------|
| | Year | Cusec | | | | | |
| 1 | 2 | 58473 | 47948 | 51328 | 51328 | 43888 | 51328 |
| 2 | 5 | 95072 | 79836 | 89765 | 89765 | 76321 | 89765 |
| 3 | 10 | 114221 | 104244 | 115214 | 115214 | 107723 | 115214 |
| 4 | 20 | 130031 | 129927 | 139626 | 139626 | 147623 | 139626 |
| 5 | 25 | 134635 | 138535 | 147369 | 147369 | 162688 | 147369 |
| 6 | 50 | 147819 | 166464 | 171224 | 171224 | 217878 | 171224 |
| 7 | 100 | 159676 | 196360 | 194902 | 194902 | 288490 | 194902 |
| 8 | 500 | 183673 | 274305 | 249619 | 249619 | 537702 | 249619 |
| 9 | 1000 | 192894 | 311906 | 273142 | 273142 | 696933 | 273142 |
| 10 | 10000 | 220235 | 456495 | 351244 | 351244 | 161089 | 351244 |

5.2 Recharge Calculation for River Ravi

Recharge from River Ravi was estimated on the basis of individual x-sectional area and seepage rate. Maximum discharge of 18000 cusec was observed in the period of 2000-2014 so that discharge is used to check the maximum river Ravi recharge capacity. The observed recharge in the River Ravi is about 6978784 ft³ per annum, if the flow of

18000 cusec pass through the River Ravi. The recharge due to DAM structure was estimated as 24666069 ft³/yr. as shown in **Table 5.2**.

Table 5.2: Recharge Calculation for River Ravi

| Description | Units | River Ravi |
|---------------------|----------|-------------|
| Discharge | Cusec | 18000 |
| X-Sectional Area | sq. m | 16843291.26 |
| Velocity | ft/sec | 0.037710916 |
| Permeability | m/day | 0.086 |
| Seepage Rate | Cusec | 569.66 |
| Normal Recharge | Cubic ft | 6978783 |
| Recharge Due to DAM | Cubic ft | 24666069 |

5.3 HEC-RAS Results

HEC RAS model was run to calculate the water surface profile for the flow condition in the river Ravi and to assess the effect of ponding in the river Ravi discharge and hydraulic characteristics. HEC-RAS model was run on three scenarios for river Ravi.

- 1- Steady flow analysis without installing dam
- 2- Unsteady flow analysis without installing dam at max flow
- 3- Unsteady flow analysis with dam installed and maximum flow in the year.

5.3.1 Steady flow analysis

Steady flow analysis was used for model calibration and then that model was used for unsteady analysis. For steady flow analysis, the velocity profile was drawn for the whole channel and it was found that max velocity in the channel is 6 ft/sec at 19 mile D/S of the Indian boarder and minimum velocity observed was 0.05ft/sec in the channel. The steady state analysis showed that flow in the channel is subcritical as shown in **Figure 5.6**.

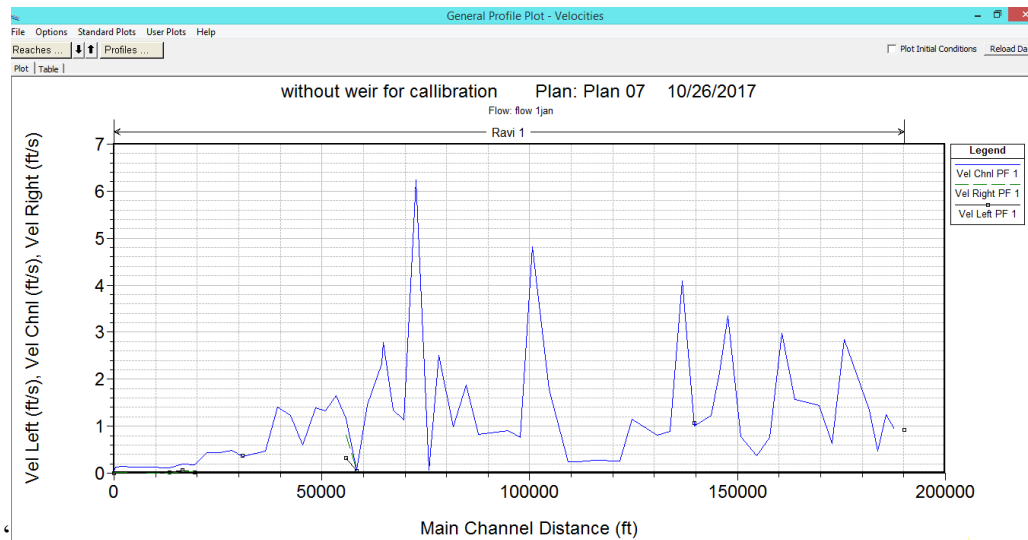


Figure 5.6: Velocity profile in river Ravi for steady flow condition

Water surface profile for the 38 mile reach of river Ravi was determined in HEC-RAS for the steady state conditions and run on the flow observed at 1 Jan 2014 at Shahdra station and then simulated the water surface profile for the channel. The flow for 1 Jan was assumed as base line for the callibratiion of HEC RAS model and water surface depth changed 37ft from U/S to D/S of the river channel as shown in **Figure 5.7**.

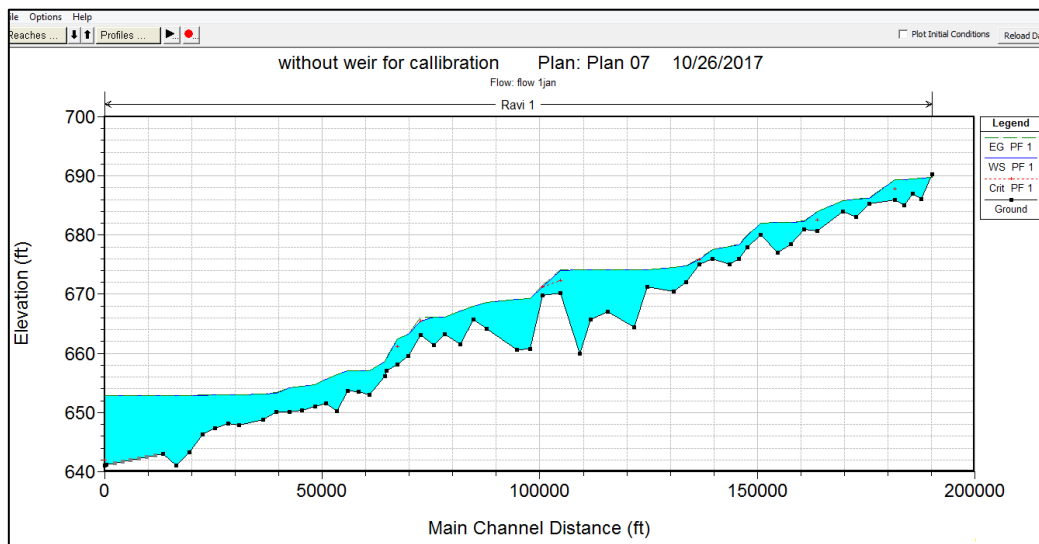


Figure 5.7: Water Surface Profile of river Ravi for steady flow

5.4 Model Calibration

HEC-RAS model was calibrated using steady state flow analysis for the study area reach. For calibration purpose, values of Manning roughness ‘n’ was changed between 0.015-0.025 for channel left over bank (LOB) and right over bank (ROB), while for the main channel the value of Manning’s n was changed between 0.018- 0.03, and then flow was calculated at that section and simulated flow was compared for the better results. The value of Manning’s n for main channel was 0.018 and for banks, a value of 0.015 was selected as these results at 99.78% accuracy as shown in **Table 5.3**.

Table 5.3: HEC-RAS Model Calibration

| Manning’s roughness coefficient “n” | | | | | | | |
|-------------------------------------|-------|-------|-----------|-------------|--------------|-------|----------|
| LOB | Main | ROB | Flow Date | Observed WL | Simulated WL | Error | Accuracy |
| 0.025 | 0.030 | 0.025 | 1-01-09 | 672.6 | 674.66 | 0.305 | 99.69 |
| 0.027 | 0.031 | 0.027 | 1-01-09 | 672.6 | 674.77 | 0.322 | 99.68 |
| 0.023 | 0.029 | 0.023 | 1-01-09 | 672.6 | 674.62 | 0.299 | 99.70 |
| 0.030 | 0.032 | 0.030 | 1-01-09 | 672.6 | 674.74 | 0.322 | 99.68 |
| 0.020 | 0.022 | 0.020 | 1-01-09 | 672.6 | 674.29 | 0.250 | 99.75 |
| 0.018 | 0.020 | 0.018 | 1-01-09 | 672.6 | 674.18 | 0.234 | 99.76 |
| 0.015 | 0.018 | 0.150 | 1-01-09 | 672.6 | 674.04 | 0.214 | 99.78 |

5.5 Unsteady Flow Analysis without installation of dam

When hec ras model was run for the unsteady conditon and water surface profile was computed. The maximum flow observed in the whole year was 18000 cusec and against that flow a water depth of 5-18 ft was observed in the River Ravi reach as shown in **Figure 5.8**.

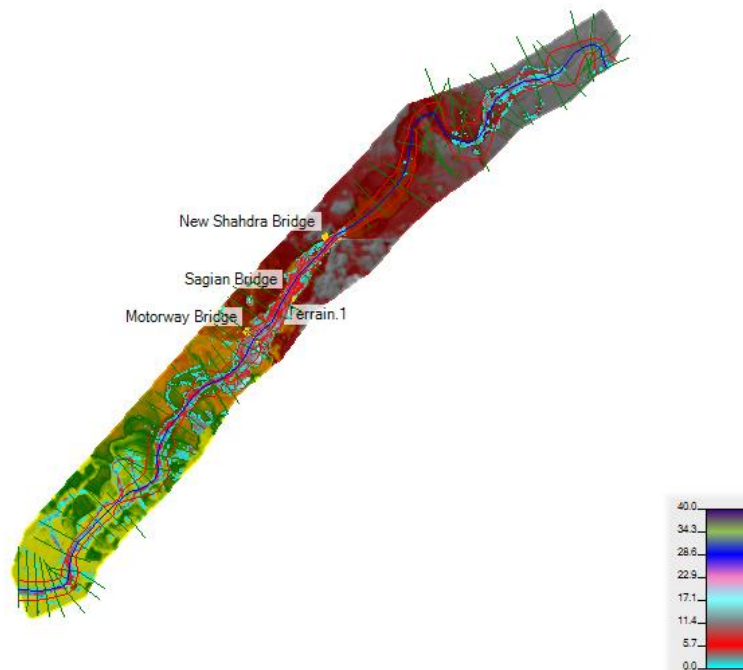


Figure 5.8: Water Surface Profile at max flow available

The river stage and discharge curve was calculated in HEC- RAS for every x-section in the river channel and it was found that the maximum river flow of 18000 in 5 July and the water depth of 18 ft was calculated. The max water surface elevation of 697 ft was observed at the flow of 18000 cusec. The water surface profile is shown in

Figure 5.9

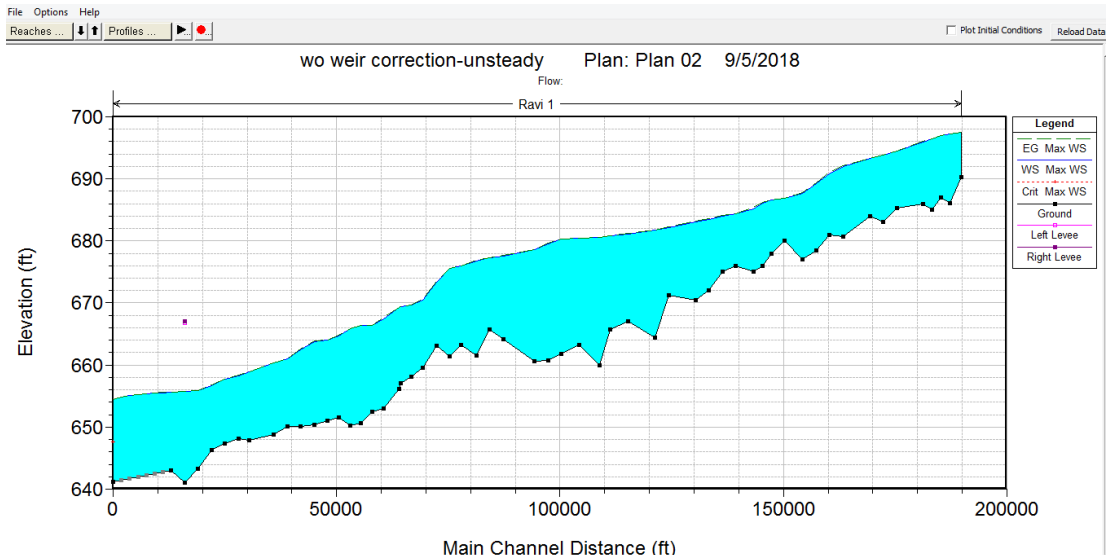


Figure 5.9: Water Surface Profile for the River Ravi

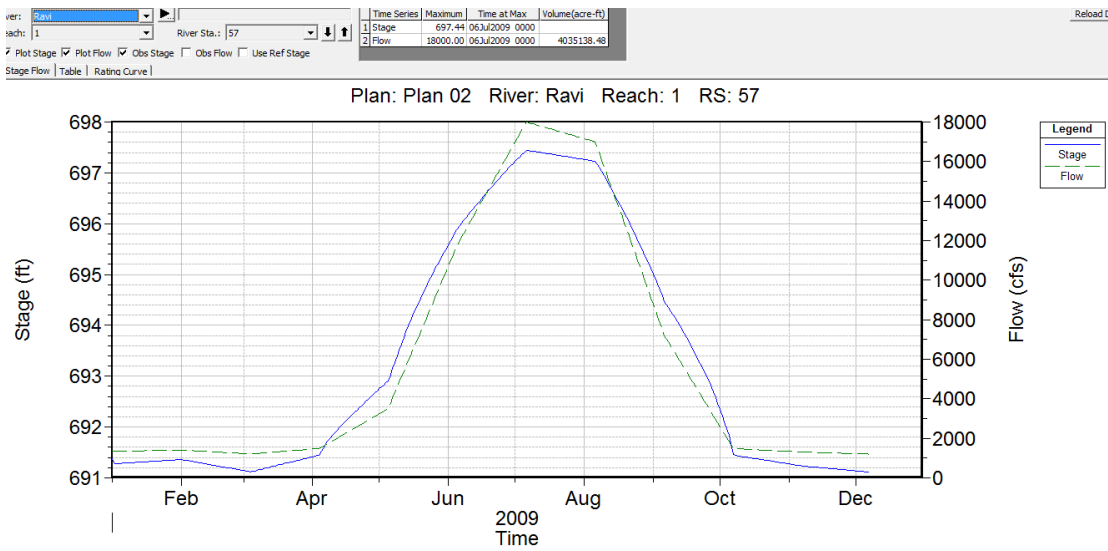


Figure 5.10: River Stage and Flow graph at D/S

The rating curve at the dam site is observed and found that the at flow less than 1200 cusec the stage is increased at a steep rate but after that flow the stage is increased steady and a maximum of 697 ft stage was observed for the flow of 18000 cusec at Downstream as shown in **Figure 5.10**

5.6 Unsteady flow with Maximum flow conditon and dam

HEC-RAS model was run for the high flow condition and water surface profile was estimated at each cross section and it was found that in the month of July the maximum water surface elevation was calculated. During the moonsoon season a flow of about

18000 cusec was calculated in the month of July and for the month of May and August the water surface profile deplete again as shown in **Figure 5.10**. The max water depth observed is 15 ft against a flow of 18000 cusec in case of dam installed as shown in **Figure 5.11**. Water surface elevation observed in July is 697 for the flow of 18000 cusec at the U/S of the river Ravi due to dam as shown in **Figure 5.12**

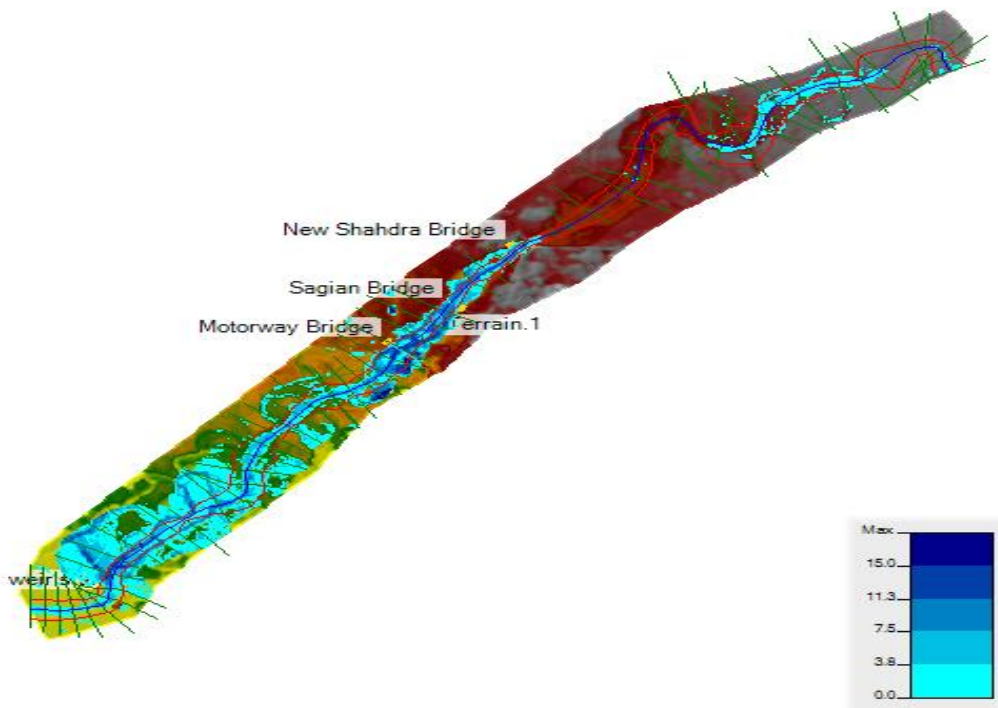


Figure 5.11: Water Surface Profile for the River Ravi after installation of DAM

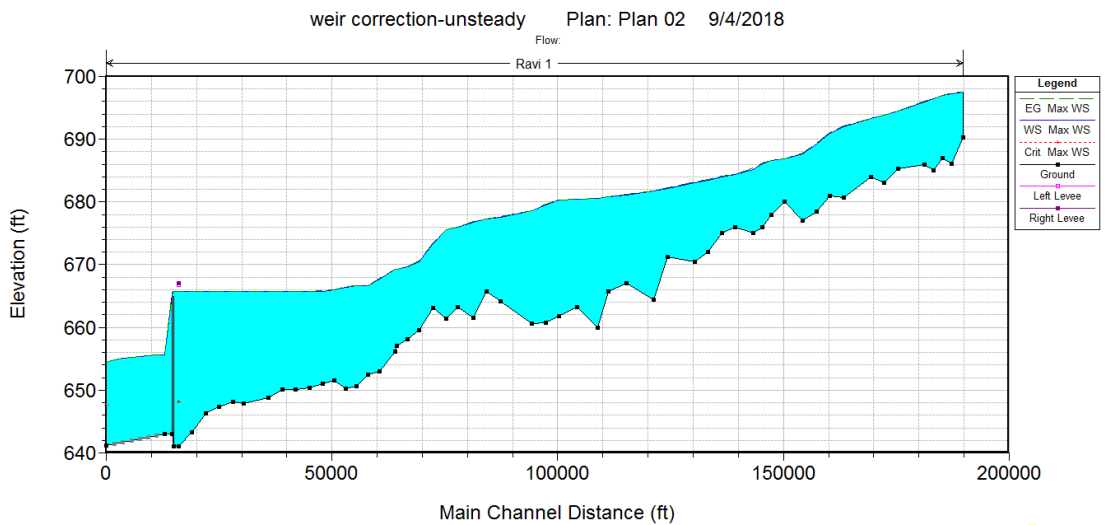


Figure 5.12: Maximum water surface profile with dam installed

The rating curve for the River Ravi was calculated in all section and curve for the dam installed is shown in **Figure 5.13**. Figure shows that at the beginning for a flow of less than 180 cusec, the stage in river rises steep until a Water surface of 656.62ft and after that, a constant trend between water surface and discharge was observed for a flow of 10000 cusec. The stage observed is only 667 ft. and after that, the water surface remains almost constant even for a flow of 10000 cusec as shown in **Figure 5.13**.

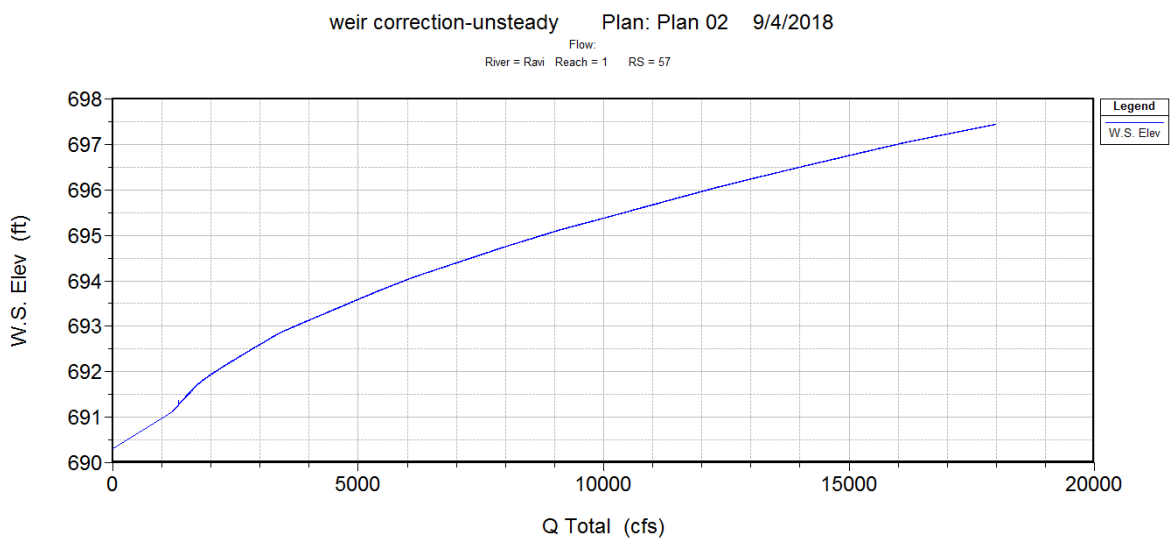


Figure 5.13: Rating Curve for D/S after installation of dam in river Ravi

5.7 Visual MODFLOW

Visual MODFLOW was run for the data of 2003-2014 for pumping wells discharge and recharge from river Ravi and rainfall for different scenarios and then calibrated using 13 observation wells in the vicinity of district Lahore. The Water Level contour for the year 2014 is shown in **Figure 5.14**. The maximum water level observed was 210 m in the vicinity of BRBD canal and along the river Ravi the water level observed varied between 177 m to 210 m. The flow data and pumping data was assumed constant for the whole year that is the model was based on no pumping change scenario.

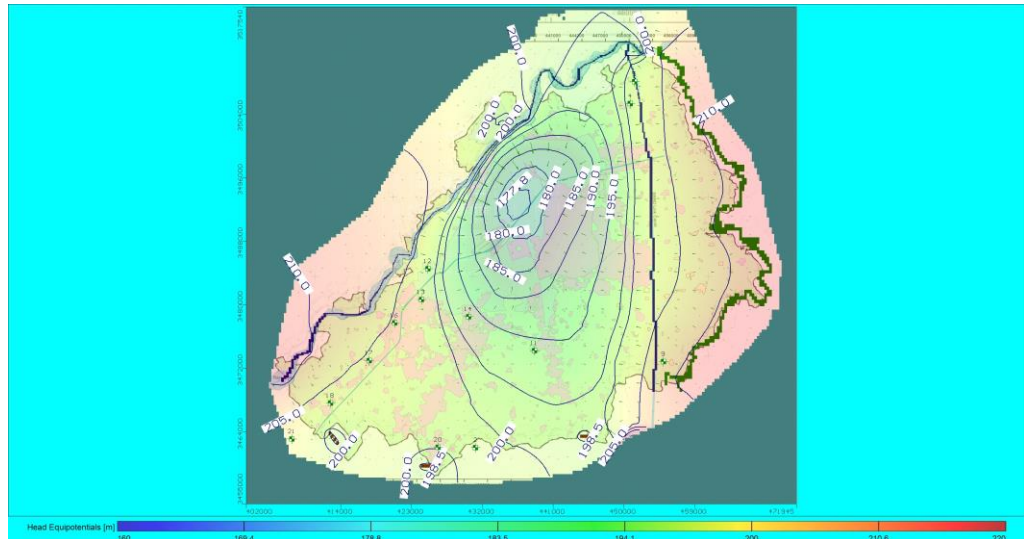


Figure 5.14: Groundwater contours for the year 2014

The groundwater model developed for the River Ravi was calibrated on the 2003-2014 observation wells data and the correlation coefficient for the 12 Year period was 0.94 with the standard error of 0.523 m for the groundwater model. The calibration was based on R^2 & standard error and 95% confidence interval based values was observed for the 12 Year period as shown in Figure 5.15

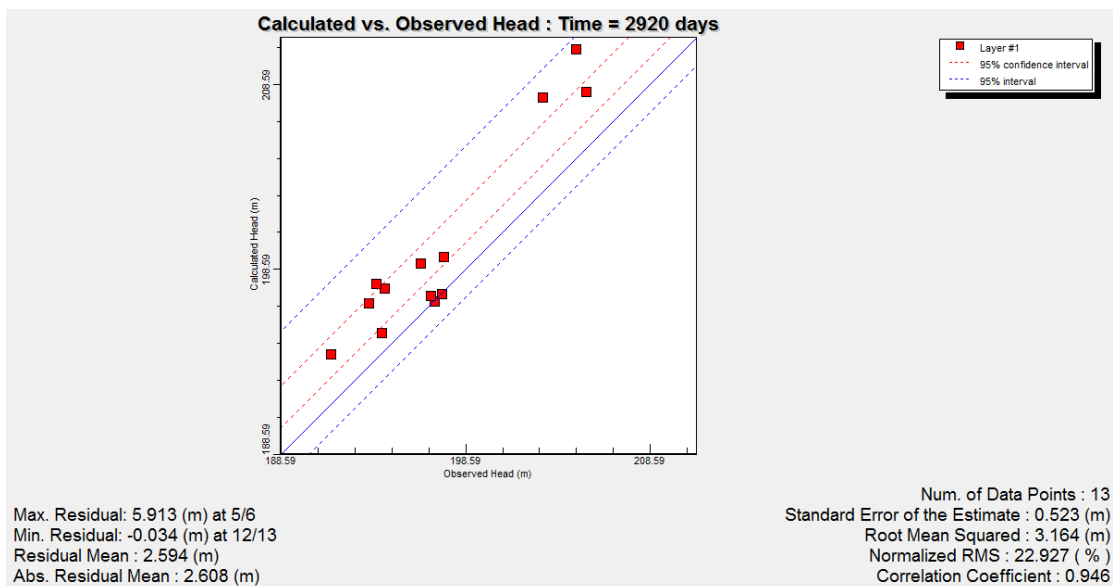


Figure 5.15: Model Calibration for year 2003-2010

The model was run for 12 years data from the year 2003-2014. Then after calibration river stage was changed to check the effect of dam in River Ravi. Since due to dam the

river stage rise the effect will be visible. The groundwater model was validated for the year 2010-2014 and a R^2 value of 0.91 and standard error of 0.63 m was found for the year 2014 as shown in **Figure 5.16**.

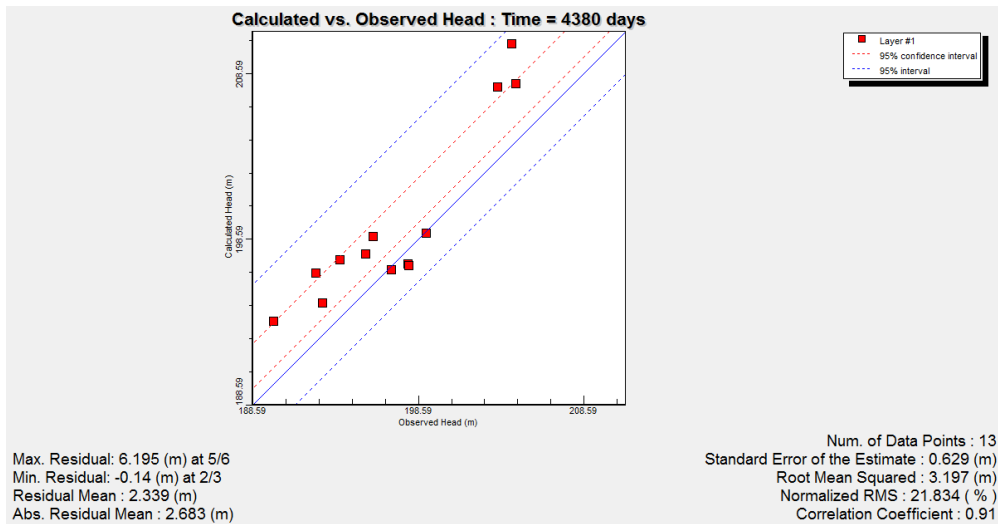


Figure 5.16: Model Validation for year 2014

Groundwater model contours were estimated for the current situation of the year 2017 without changing river stage and assuming the recharge and pumping data constant, as 2014 due to data availability constraint and a minimum water level observed is 172.4m for the year 2017 as shown in **Figure 5.17**.

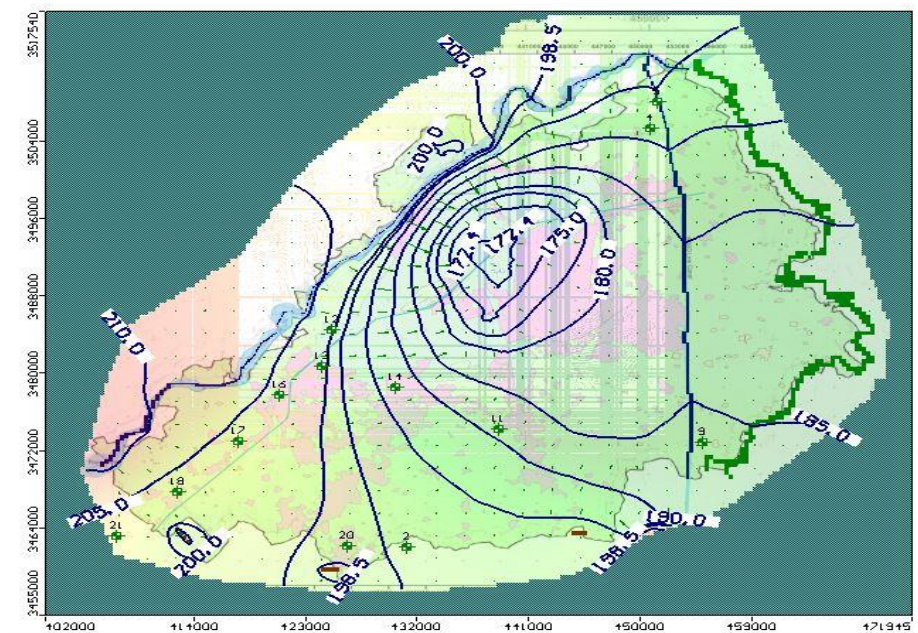


Figure 5.17: Groundwater contour for year 2017 situation

Groundwater Model was extrapolated to the year 2020 for the estimation of Water level. Considering the entire parameters constant and to check what will be future scenario if there is no change in river stage. Tube wells discharge or recharge from the rainfall assuming as constant and it was found that the contour of the 175 m is expanded exponentially and more area is observed where water level was dropped and a minimum of 172 m water level was observed in this scenario as in **Figure 5.18**.

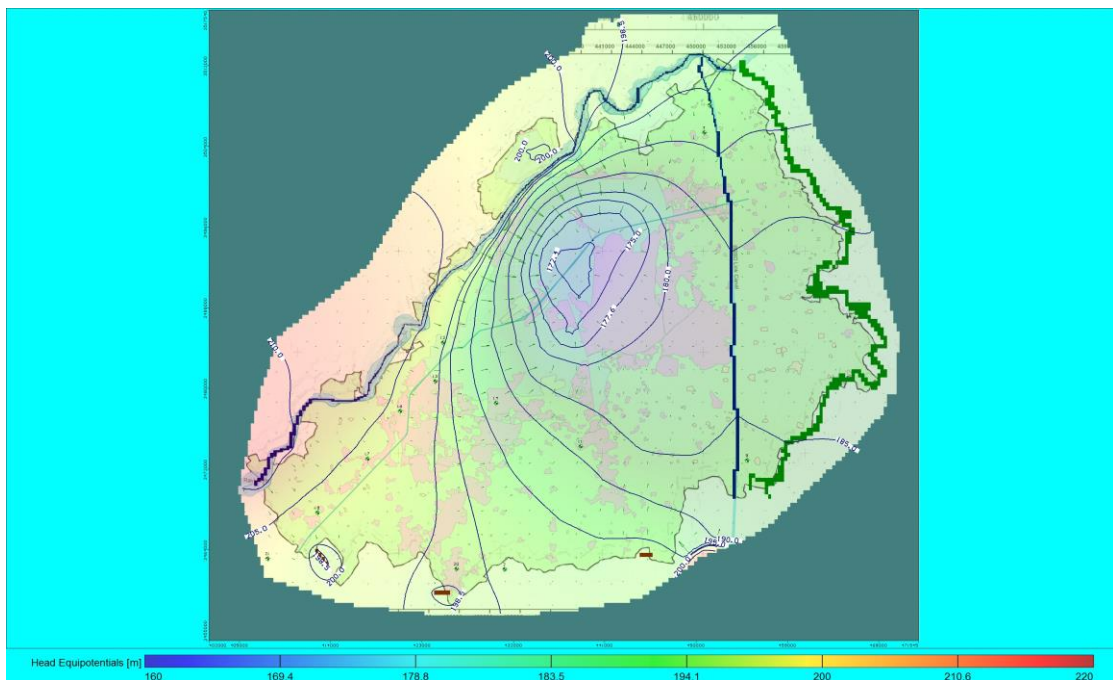


Figure 5.18: Contour 2020 with no changes in River stage

5.8 Groundwater Situation after Rising the River Stage

When groundwater model was run for the change in river stage the contour of 200 m was almost disappear and water level rise for the area. The river stage was increased to 3 m than the previous stage to check the effect better the parameters such as pumping and conductivity parameter was assumed constant and model was run for the steady state analysis for the better result.

The ground water level observed after one year of changing river stage and found that the contour of 172 m was disappeared and a minimum water level observed was 175 m near the vicinity of River Ravi as shown in **Figure 5.19**.

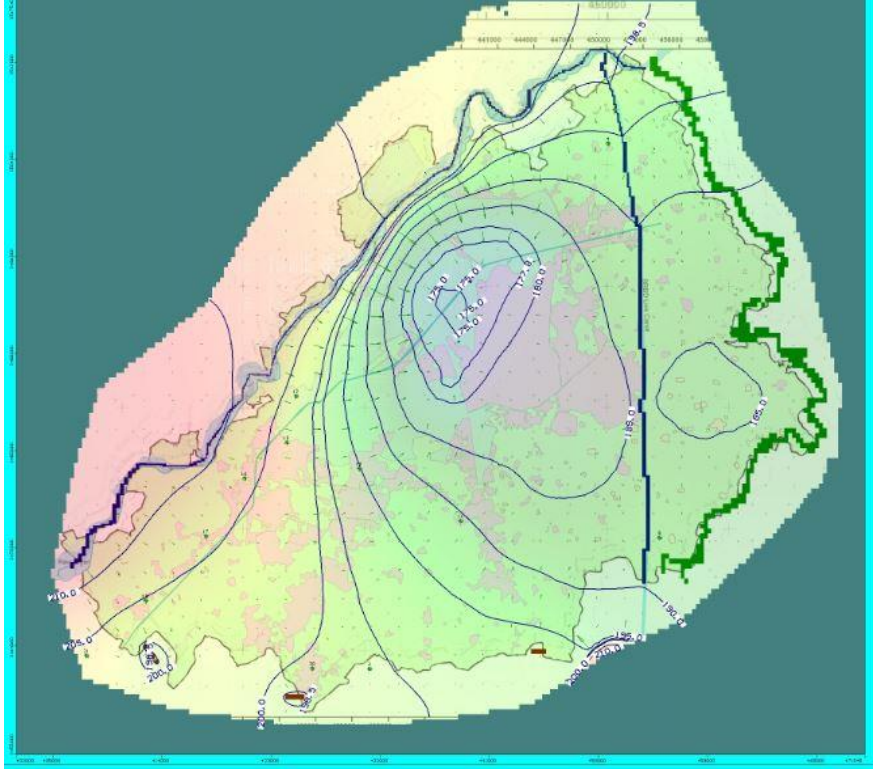


Figure 5.19 : Groundwater contour after 1 year of changing the river Stage

The model was extrapolated till year 2020 after changing the river stage and the area of the Lahore city which has water level less than 175 m was reduced near river Ravi and a maximum of 185 m contour was observed in the overall Lahore city vicinity as shown in **Figure 5.20**.

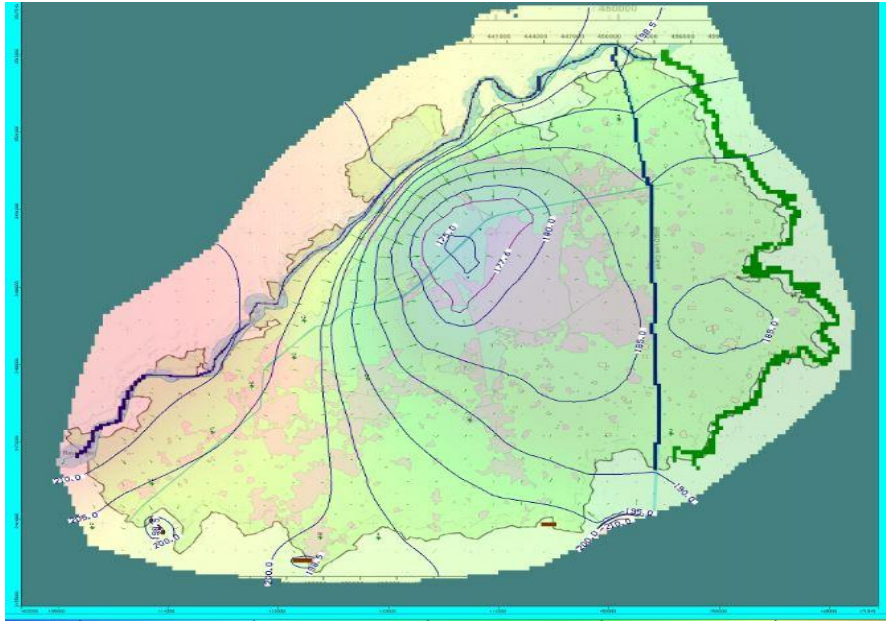


Figure 5.20: Contour after 3 years changing river stage (2020)

The maximum ground water lever observed in the Lahore city was about 80 ft in the center of Lahore city. Groundwater level was depleting at the rate of of about 1-2 ft per year in the Lahore city. The areas which are close to the River ravi has less groundwater depletion rate compare to the areas that are in the center of the city. Most of the observation wells which are close to River has less groundwater depletion. The water level in different observation wells are shown in Figure 5.21.

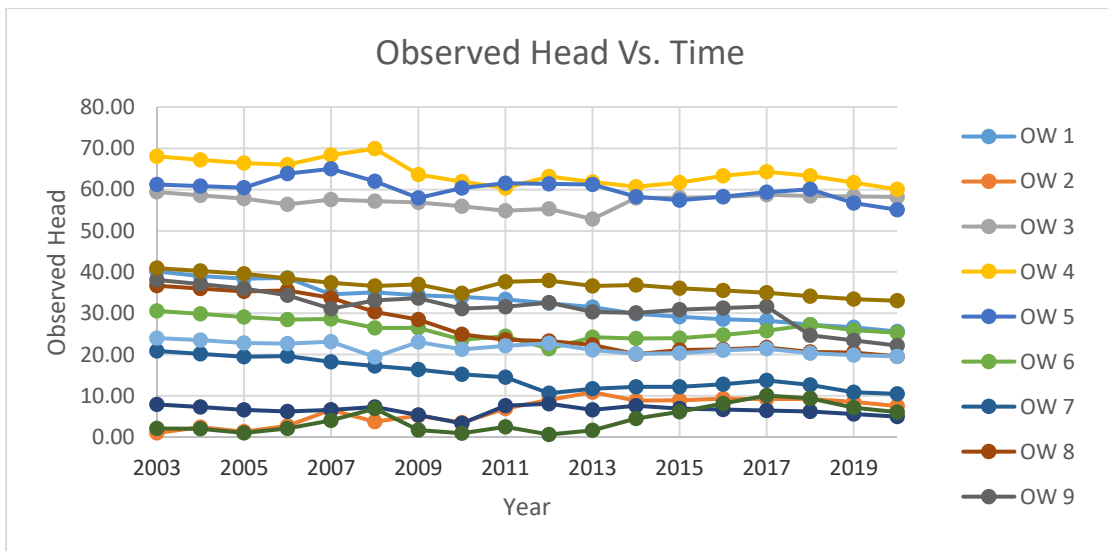


Figure 5.21: Simulated Head Vs. Time in different Location of Lahore city

A dam structure is installed at 3 mile U/S of the Mohlanwall, and 5 mile downstream of the Motorway bridge and water surface profile was computed in the HEC-RAS model. The proposed dam height was 10 ft and the x-section of River Ravi at the proposed location was about 1200 ft. The river x-section of 12000 ft was used for HEC-RAS modelling including flood banks in the study. The back water effect is shown in below Figure. The max height of 15' was observed in the channel after installation of the dam in the River Ravi. The recharge in the river Ravi in normal situation is about 6978783 ft^3 . When we construct a structure we can get an extra recharge of 24666069 ft^3 due to the dam structure as shown in **Figure 5.22**

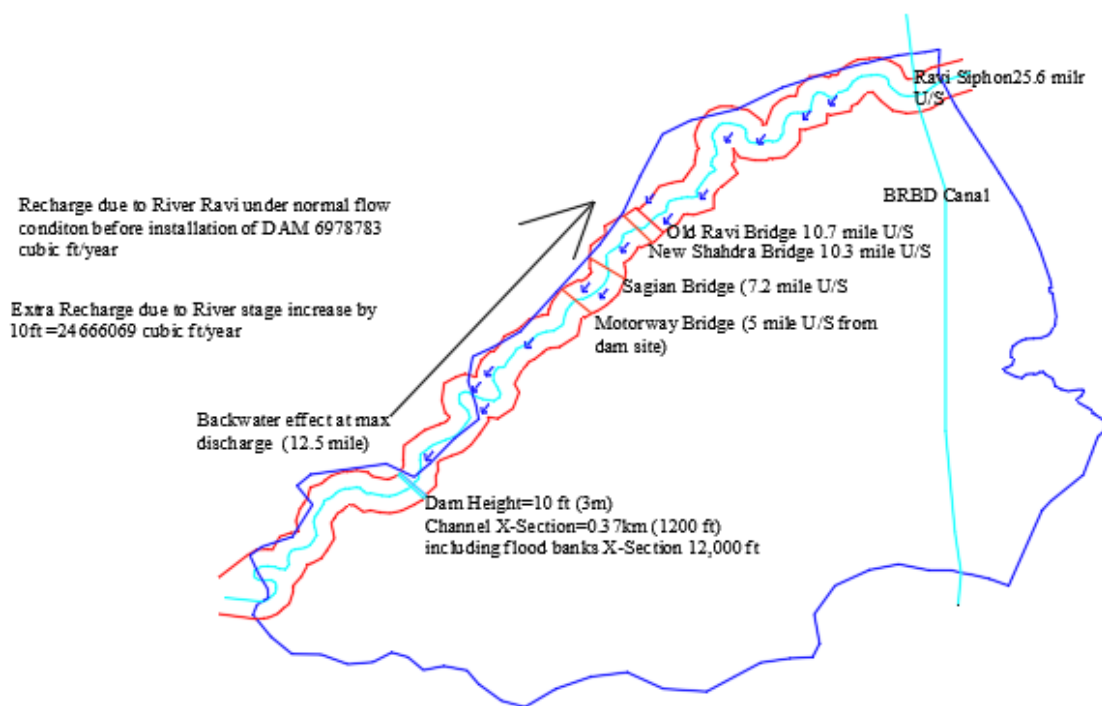


Figure 5.22: Study Area map including proposed dam site

5.9 Groundwater Level Interpolation in ArcMap

Groundwater level of Lahore city was estimated using Visual MODFLOW and then the water level from the ground is computed and groundwater maps of Lahore city was developed in ArcMap for year (2003-2020) before and installation of dam structure.

Thirteen observation well points were used to interpolate the groundwater level on simulated and observed case scenario.

The observed water level maps (2003 & 2017) with their area are shown in **Figure 5.23** and **Figure 5.24**.

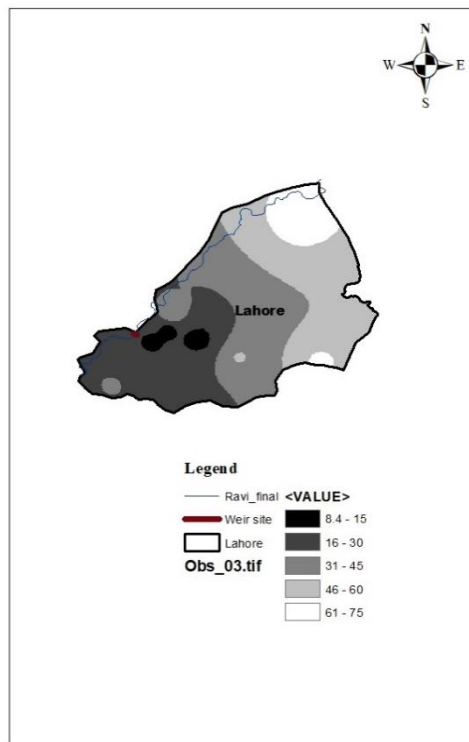


Figure 5.23: Observed GW level (ft.) map of Lahore city for year 2003

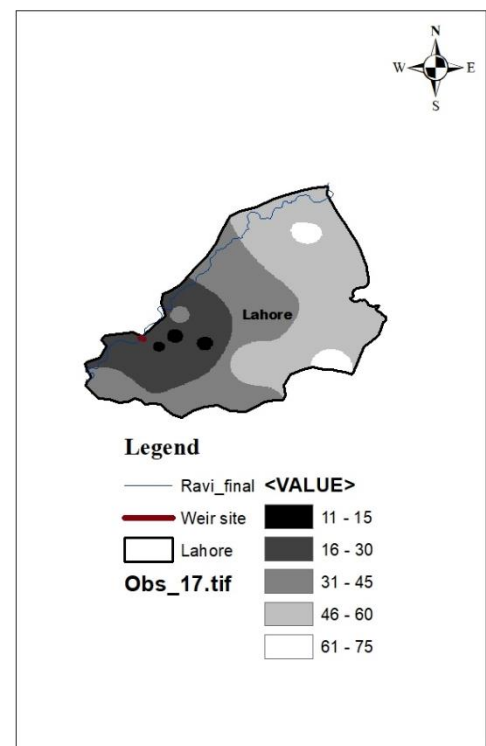


Figure 5.24: Observed GW level (ft.) map of Lahore city for year 2017

Variation of observed ground water level of Lahore city from 2003 to the current situation is shown in Table 5.4. the observed water level near the proposed Dam site is 1-15 ft during the year 2003 which is increase to 11-20 ft. during the year 2017 with a 4 mile² more area than the year 2003.

Table 5.4: Variation of observed water levels before installation of DAM in Lahore city (2003-2017)

| Year | WL (ft) | Area (mile ²) | Year | WL (ft) | Area (mile ²) |
|------|---------|---------------------------|------|---------|---------------------------|
| 2003 | 1-15 | 19 | 2017 | 1-15 | 14 |
| 2003 | 16-30 | 202 | 2017 | 16-30 | 148 |
| 2003 | 31-45 | 224 | 2017 | 31-45 | 219 |
| 2003 | 46-60 | 172 | 2017 | 46-60 | 250 |
| 2003 | 61-75 | 44 | 2017 | 61-75 | 30 |

Simulated groundwater maps are developed for the Lahore city as shown in **Figure 5.25** and **Figure 5.26**

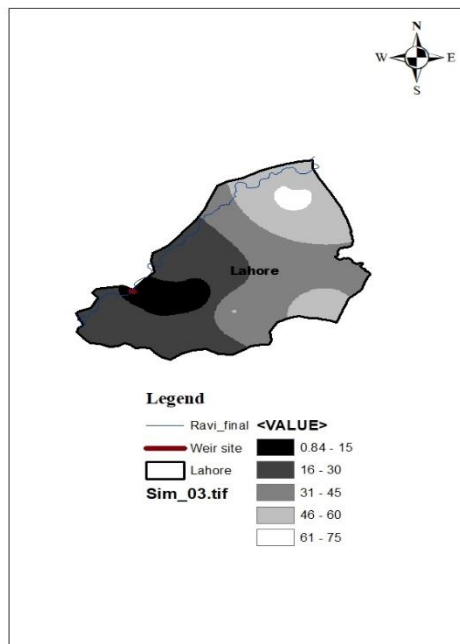


Figure 5.25: Simulated GW level (ft.) map of Lahore city for year 2003

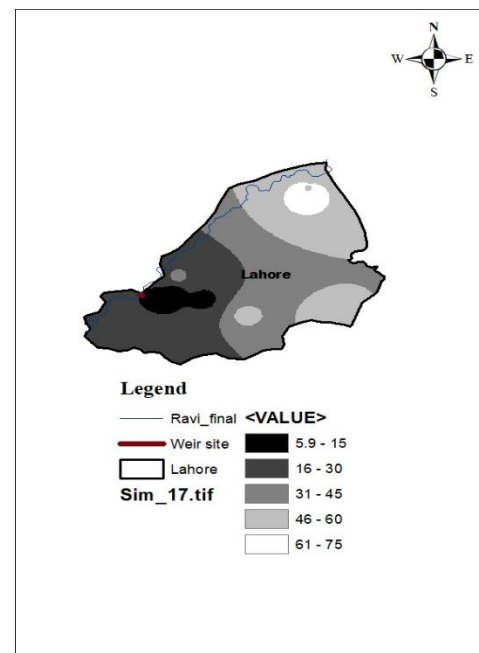


Figure 5.26: Simulated GW level (ft.) map of Lahore city for year 2017

Variation of simulated ground water level of Lahore city by using Visual MODFLOW is shown in **Table 5.5**. where the water level near the proposed dam site increase from 2-11 ft to 5-16 ft with an increase in area variation for year 2003-2017

Table 5.5: Variation of simulated water levels before installation of DAM in Lahore city (2003-2017)

| Year | WL (ft) | Area (mile ²) | Year | WL (ft) | Area (Km ²) |
|------|---------|---------------------------|------|---------|-------------------------|
| 2003 | 1-15 | 35 | 2017 | 1-15 | 16 |
| 2003 | 16-30 | 176 | 2017 | 16-30 | 145 |
| 2003 | 31-45 | 211 | 2017 | 31-45 | 264 |
| 2003 | 46-60 | 193 | 2017 | 46-60 | 214 |
| 2003 | 61-75 | 55 | 2017 | 61-75 | 24 |

Visual MODFLOW is run for three year 2018-2020 to check the effect of the DAM structure on groundwater levels. Then groundwater maps are developed in ArcMap showing water levels and their areas of influence by each year.

Groundwater level maps for the Lahore city is developed in ArcMap after raising the water level (DAM structure) and effect of water level fluctuation is checked in Visual MODFLOW which then mapped in ArcMap as shown in **Figure 5.27**, **Figure 5.28** and **Figure 5.29** for year 2018-2020 respectively. Groundwater interpolation maps are shown in Annexure E.

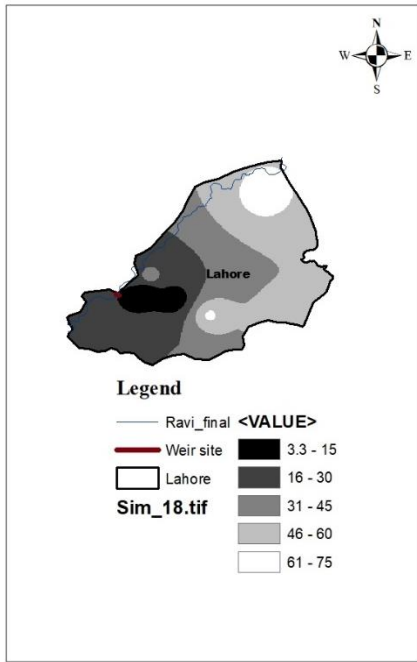


Figure 5.27: Simulated GW level (ft.) map of Lahore city after installation of DAM in year 2018

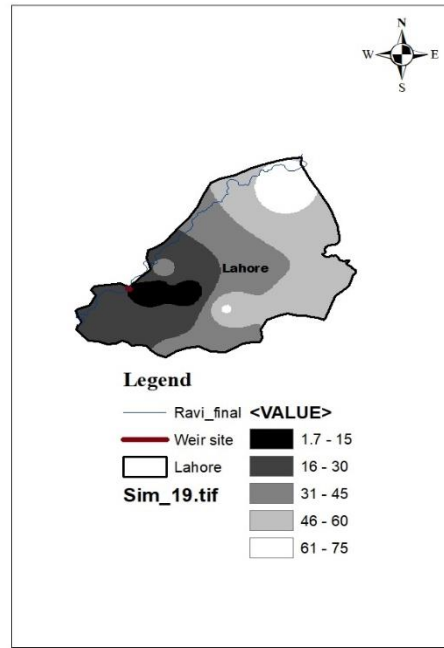


Figure 5.28: Simulated GW level (ft.) map of Lahore city after installation of DAM in year 2019

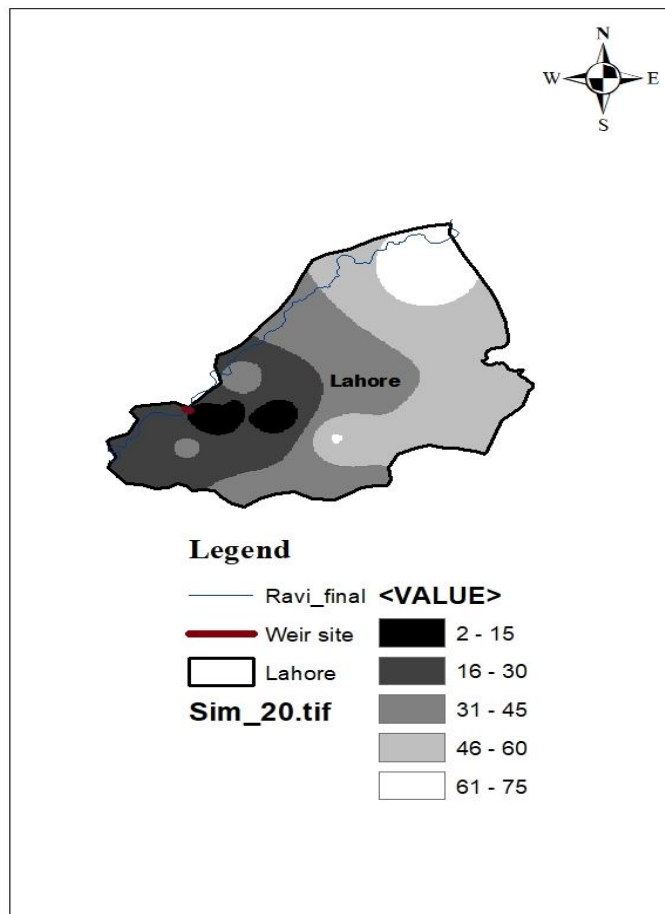


Figure 5.29: Simulated GW level (ft.) map of Lahore city after installation of DAM in year 2020

Groundwater maps were developed for year 2003-2017 for observed and simulated scenario the remaining maps are shown in Annexure D.

After raising the river stage by construction of dam the water level is estimated and water level is reduced for 5-16 ft to 2-12 ft during year 2017-2020 (3 years after construction of dam) the area reduction was calculated as shown in **Table 5.6** and **Table 5.7**

Table 5.6: Variation of simulated water levels after installation of dam in Lahore city (2017-2018)

| Year | WL (ft) | Area (mile ²) | Year | WL (ft) | Area (mile ²) |
|------|---------|---------------------------|------|---------|---------------------------|
| 2017 | 1-15 | 16 | 2018 | 1-15 | 14 |
| 2017 | 16-30 | 145 | 2018 | 16-30 | 105 |
| 2017 | 31-45 | 264 | 2018 | 31-45 | 195 |
| 2017 | 46-60 | 214 | 2018 | 46-60 | 310 |
| 2017 | 61-75 | 24 | 2018 | 61-75 | 26 |

Table 5.7: Variation of simulated water levels after installation of dam in Lahore city (2019-2020)

| Year | WL (ft) | Area (mile ²) | Year | WL (ft) | Area (mile ²) |
|------|---------|---------------------------|------|---------|---------------------------|
| 2019 | 1-15 | 12 | 2020 | 1-15 | 9 |
| 2019 | 16-30 | 103 | 2020 | 16-30 | 97 |
| 2019 | 31-45 | 198 | 2020 | 31-45 | 204 |
| 2019 | 46-60 | 312 | 2020 | 46-60 | 318 |
| 2019 | 61-75 | 25 | 2020 | 61-75 | 22 |

The difference in water level throughout the Lahore city was estimated using ArcMap and showing in **Figure 5.30**, **Figure 5.31** and **Figure 5.32** for 2018, 2019 and 2020 respectively

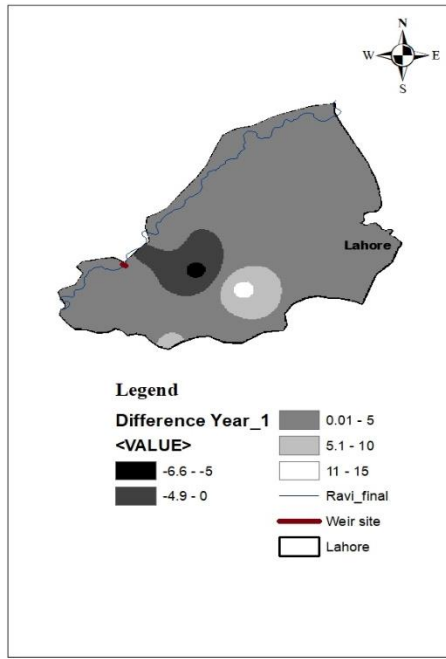


Figure 5.29: water level difference map of Lahore for year 2018

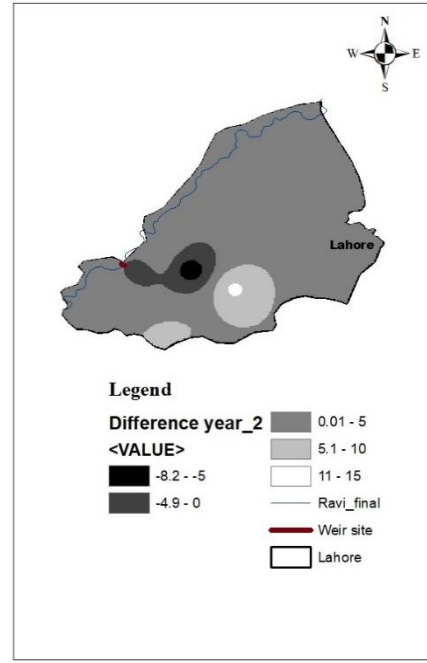


Figure 5.30: water level difference map of Lahore for year 2019

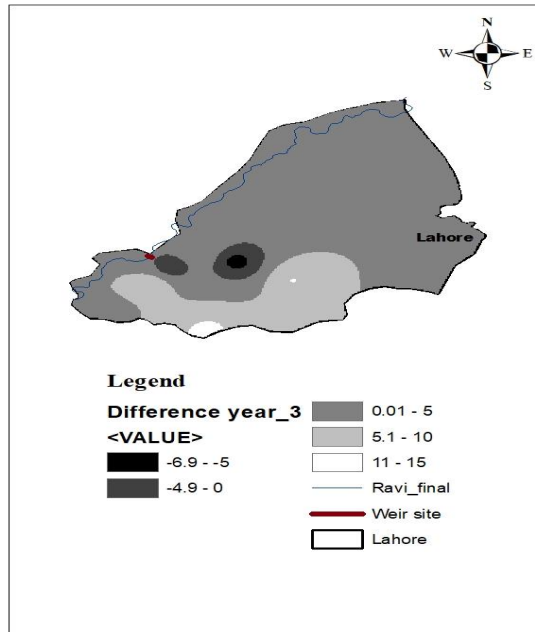


Figure 5.30: water level difference map of Lahore for year 2020

CHAPTER VI

CONCLUSIONS AND RECOMMENDATIONS

6.1 Conclusions

- Under normal condition recharge capacity of river Ravi was calculated as 6978783 ft³/year.
- After installation of dam, the extra recharge capacity 24666069 ft³/year of River Ravi was estimated.
- The water depth in river Ravi under normal condition varies between 2-10 ft depending on the available flow and location of river Ravi from proposed dam site.
- When a dam structure was installed in the River Ravi the head of the river Ravi increase and a backwater effect of about 12.5-mile U/S was simulated.
- The maximum water depth of 15 ft was simulated using HEC-RAS in river Ravi after construction of dam..

6.2 Recommendations

- To maintain aquifer level of Lahore city other recharge options like rainwater harvesting, recharging wells and recharging ponds along with river Ravi should also be consider.
- Regular monitoring of groundwater levels and more piezometric wells should be installed in Lahore city.
- A policy for the private wells should be developed and enforced to control the installation of private tube wells in the city boundary.

- There should be data sharing facility for Govt. & Private institutes that do research for the betterment of city. The data should be combined in one department to reduce the time waste and people can learn better this way.
- As river Ravi's water is polluted so a full analysis of water quality is required for the future study and then a law should be enforced on industries surrounding the River boundary that recycled water dumped in the River to reduce quality issues.

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ANNEXURES

| | |
|------------|--|
| Annexure A | Observation Wells Data |
| Annexure B | Data analysis of discharge and rainfall |
| Annexure C | River x-section from Indian Boarder to Shahdra |
| Annexure D | River x-section from Shahdra to Head Balloki |
| Annexure E | Groundwater level interpolation maps |

Annexure A
Observation Wells Data

Table A.1: Observation Wells Data

| Well Name | X [m] | Y [m] | Screen ID | Screen Elev. [m] | Obs. Time [day] | HEAD [m] |
|------------------|--------------|--------------|------------------|-------------------------|------------------------|-----------------|
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 365 | 195.938 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 730 | 195.598 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 1095 | 195.398 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 1460 | 195.4468 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 1825 | 194.252 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 2190 | 194.38 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 2555 | 194.1788 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 2920 | 194.0508 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 3285 | 193.874 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 3650 | 193.58 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 4015 | 193.31 |
| 11 | 437976.9 | 3470582 | 12 | 183.7 | 4380 | 192.8 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 365 | 199.0194 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 730 | 198.7994 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 1095 | 198.5794 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 1460 | 198.3782 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 1825 | 198.427 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 2190 | 197.7656 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 2555 | 197.7656 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 2920 | 196.903 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 3285 | 197.156 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 3650 | 196.2172 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 4015 | 197.0828 |
| 12 | 424441.9 | 3480890 | 13 | 192.7 | 4380 | 196.9792 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 365 | 195.0574 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 730 | 194.8374 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 1095 | 194.6274 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 1460 | 194.6792 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 1825 | 194.2464 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 2190 | 193.9416 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 2555 | 193.6886 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 2920 | 193.332 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 3285 | 193.1308 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 3650 | 192.9266 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 4015 | 192.2652 |
| 13 | 423653.6 | 3477039 | 14 | 185.7 | 4380 | 192.4176 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 365 | 194.8806 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 730 | 194.6706 |

| Well Name | X [m] | Y [m] | Screen ID | Screen Elev. [m] | Obs. Time [day] | HEAD [m] |
|------------------|--------------|--------------|------------------|-------------------------|------------------------|-----------------|
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 1095 | 194.4506 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 1460 | 194.5268 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 1825 | 193.9934 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 2190 | 192.9358 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 2555 | 192.3688 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 2920 | 191.302 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 3285 | 190.8936 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 3650 | 190.793 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 4015 | 190.5126 |
| 14 | 429646.1 | 3474879 | 15 | 185.7 | 4380 | 189.8268 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 365 | 196.3092 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 730 | 195.9992 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 1095 | 195.6792 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 1460 | 195.1702 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 1825 | 194.1796 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 2190 | 194.7892 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 2555 | 194.9416 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 2920 | 194.1796 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 3285 | 194.332 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 3650 | 194.6368 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 4015 | 193.951 |
| 16 | 420209.4 | 3474105 | 17 | 186.7 | 4380 | 193.8748 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 365 | 199.1754 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 730 | 198.9654 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 1095 | 198.7554 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 1460 | 198.4262 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 1825 | 198.11 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 2190 | 197.8654 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 2555 | 197.9934 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 2920 | 197.332 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 3285 | 198.1702 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 3650 | 198.2708 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 4015 | 197.8654 |
| 17 | 416986.5 | 3469341 | 18 | 189.7 | 4380 | 197.9416 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 365 | 198.1236 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 730 | 197.9136 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 1095 | 197.7036 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 1460 | 197.6 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 1825 | 197.7036 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 2190 | 197.9322 |

| Well Name | X [m] | Y [m] | Screen ID | Screen Elev. [m] | Obs. Time [day] | HEAD [m] |
|------------------|--------------|--------------|------------------|-------------------------|------------------------|-----------------|
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 2555 | 197.3226 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 2920 | 196.713 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 3285 | 198.0084 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 3650 | 198.1608 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 4015 | 197.7036 |
| 18 | 412052.6 | 3464017 | 19 | 188.7 | 4380 | 198.0084 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 365 | 197.0724 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 730 | 197.1029 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 1095 | 197.4768 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 1460 | 198.3424 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 1825 | 198.9276 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 2190 | 199.7902 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 2555 | 198.2144 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 2920 | 197.428 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 3285 | 198.446 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 3650 | 197.8852 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 4015 | 198.19 |
| 2 | 430497.1 | 3458403 | 3 | 183.7 | 4380 | 199.08 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 365 | 197.011 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 730 | 196.861 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 1095 | 196.651 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 1460 | 196.5992 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 1825 | 196.7516 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 2190 | 195.618 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 2555 | 196.7272 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 2920 | 196.1664 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 3285 | 196.4468 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 3650 | 196.6236 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 4015 | 196.1176 |
| 20 | 425765 | 3458435 | 21 | 184.7 | 4380 | 195.8372 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 365 | 193.6028 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 730 | 193.4328 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 1095 | 193.1128 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 1460 | 193.5212 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 1825 | 194.6978 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 2190 | 193.826 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 2555 | 194.2832 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 2920 | 193.7468 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 3285 | 194.7892 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 3650 | 195.4506 |

| Well Name | X [m] | Y [m] | Screen ID | Screen Elev. [m] | Obs. Time [day] | HEAD [m] |
|------------------|--------------|--------------|------------------|-------------------------|------------------------|-----------------|
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 4015 | 196.0084 |
| 21 | 407149.1 | 3459470 | 22 | 186.7 | 4380 | 195.4232 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 365 | 203.812 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 730 | 203.572 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 1095 | 203.332 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 1460 | 202.9022 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 1825 | 203.2558 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 2190 | 203.1308 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 2555 | 203.0272 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 2920 | 202.75 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 3285 | 202.4176 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 3650 | 202.57 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 4015 | 201.808 |
| 4 | 450147 | 3501695 | 5 | 195.7 | 4380 | 203.3808 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 365 | 206.4416 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 730 | 206.1916 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 1095 | 205.9416 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 1460 | 205.8166 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 1825 | 206.5268 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 2190 | 207.0084 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 2555 | 205.1034 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 2920 | 204.57 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 3285 | 204.1128 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 3650 | 204.951 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 4015 | 204.55 |
| 5 | 450710.1 | 3504464 | 6 | 197.7 | 4380 | 204.2 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 365 | 205.3588 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 730 | 205.2488 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 1095 | 205.1288 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 1460 | 206.1712 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 1825 | 206.5248 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 2190 | 205.6134 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 2555 | 204.3668 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 2920 | 205.1044 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 3285 | 205.461 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 3650 | 205.4092 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 4015 | 205.3604 |
| 9 | 454404.4 | 3469200 | 10 | 189.7 | 4380 | 204.4704 |

Annexure B:
Data analysis of discharge and rainfall

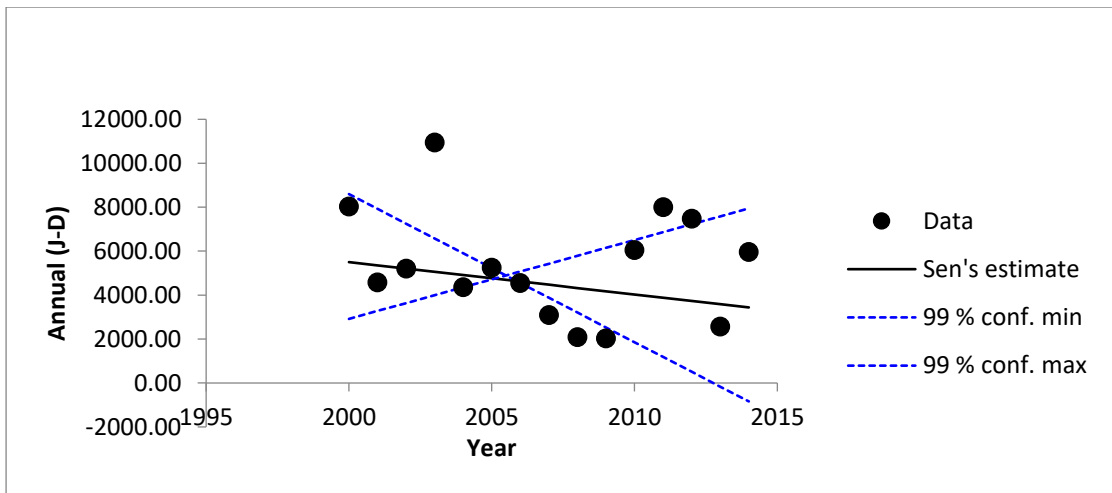


Figure B.1: Annual average daily discharge in cusec trend (2003-2014)

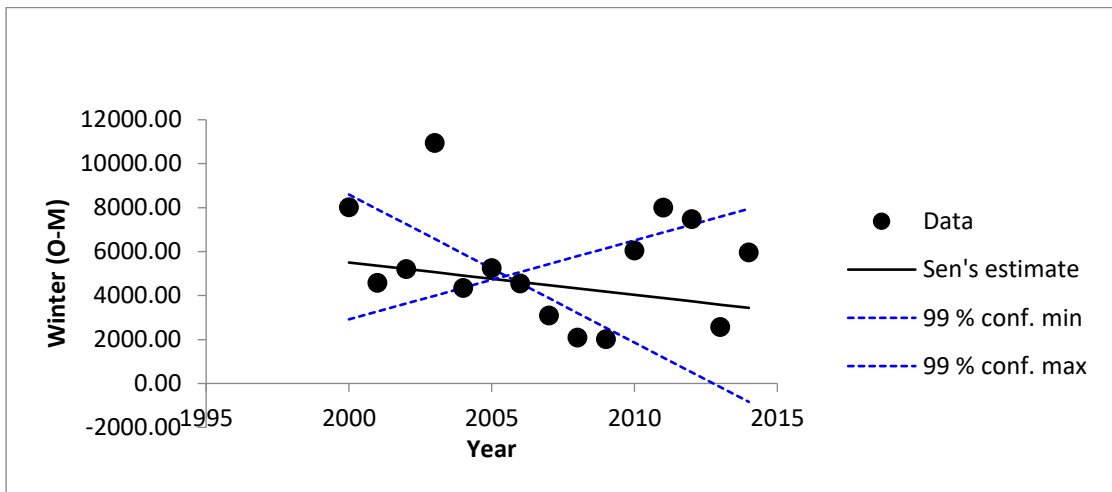


Figure B.2: Winter average daily discharge in cusec trend (2003-2014)

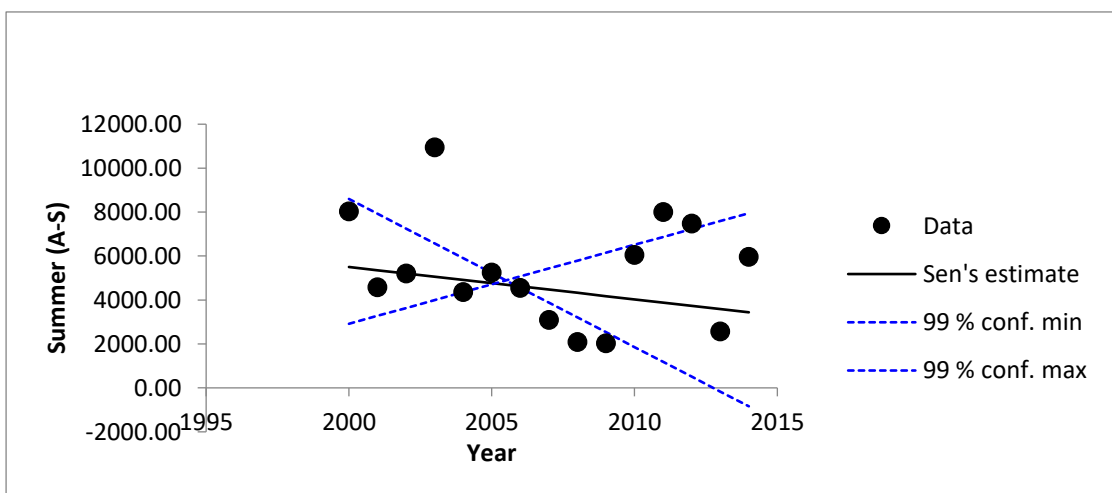


Figure B.3: Summer average daily discharge in cusec trend (2003-2014)

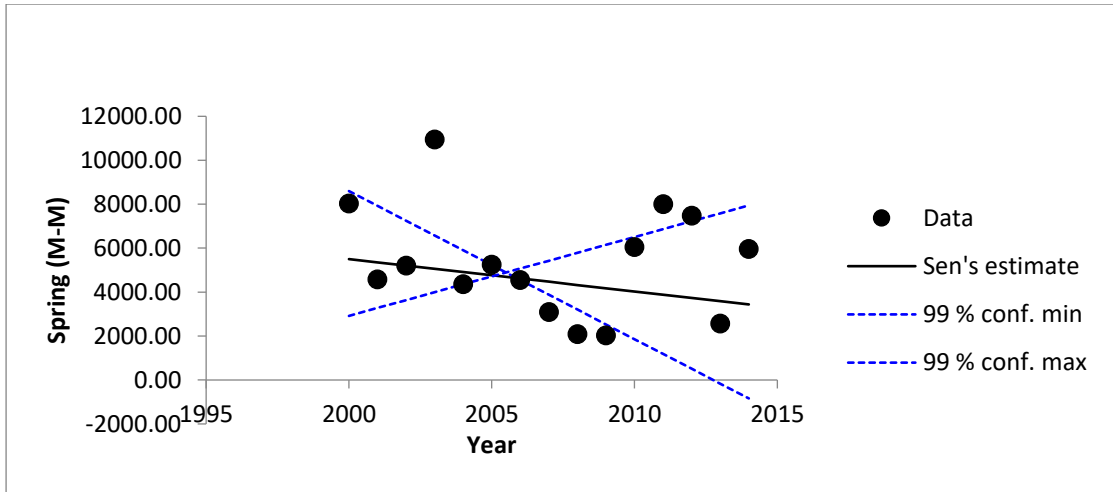


Figure B.4: Spring average daily discharge in cusec trend (2003-2014)

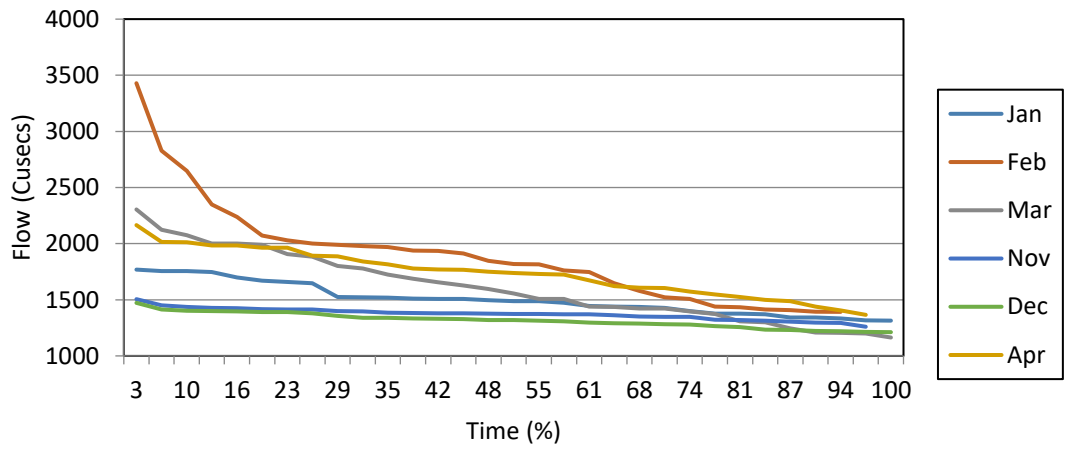


Figure B.5: Flow duration curve November - April (2003-2014)

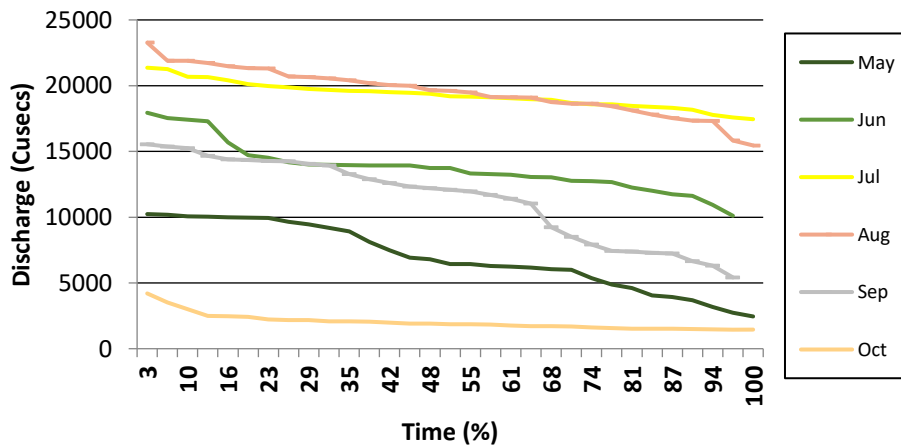


Figure B.6: Flow duration curve May - October (2003-2014)

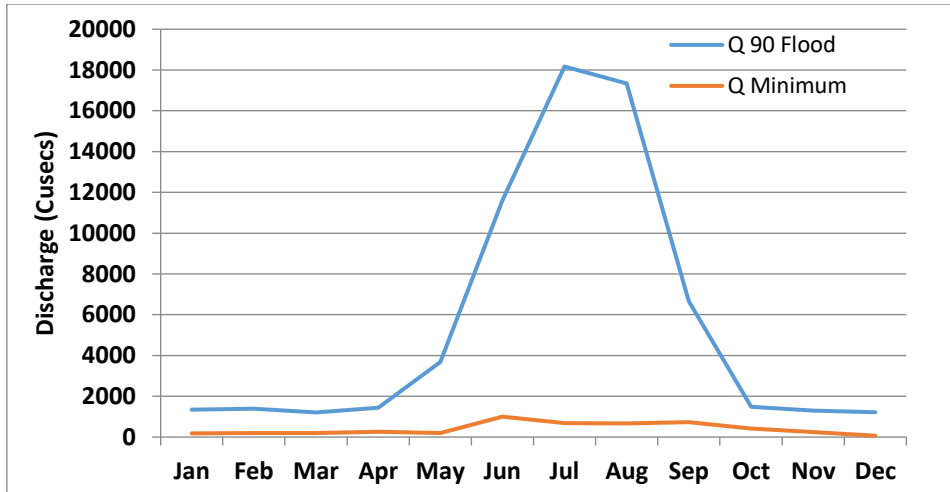


Figure B.7: Minimum and Q90 flood in river Ravi discharge in cusec (2003-2014)

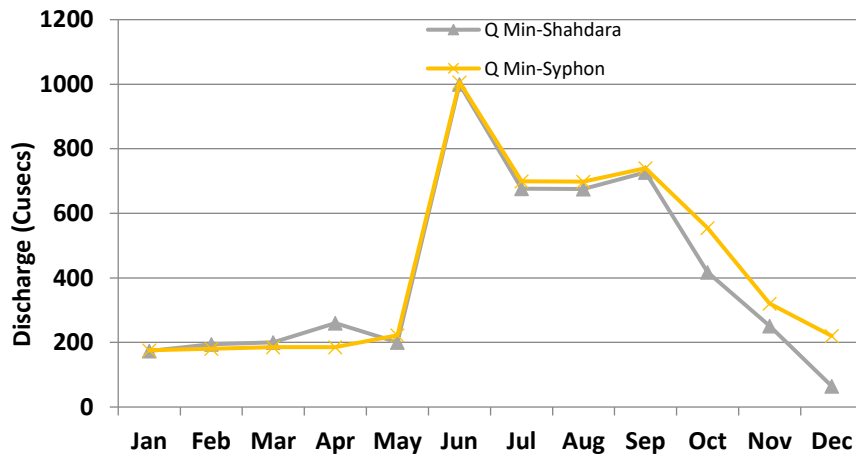


Figure B.8: Minimum flow in river Ravi discharge in cusec (2003-2014)

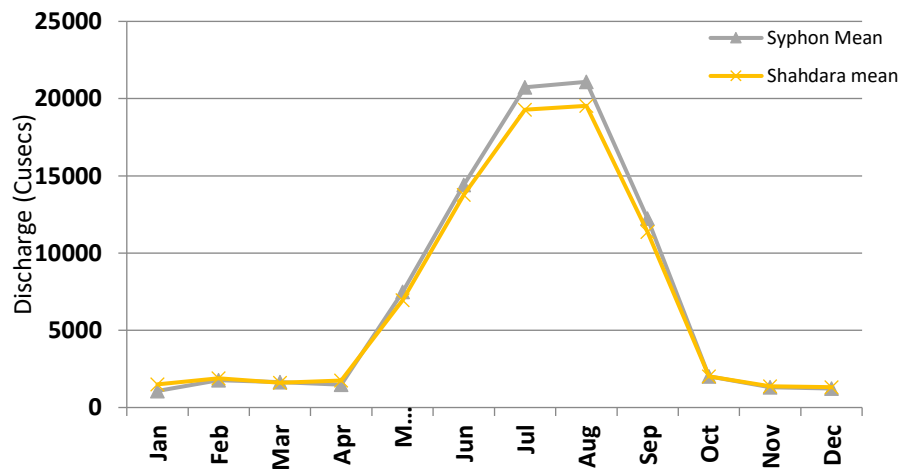


Figure B.8: Mean flow in river Ravi discharge in cusec (2003-2014)

Annexure C

River x-section from Indian Boarder to Shahdra

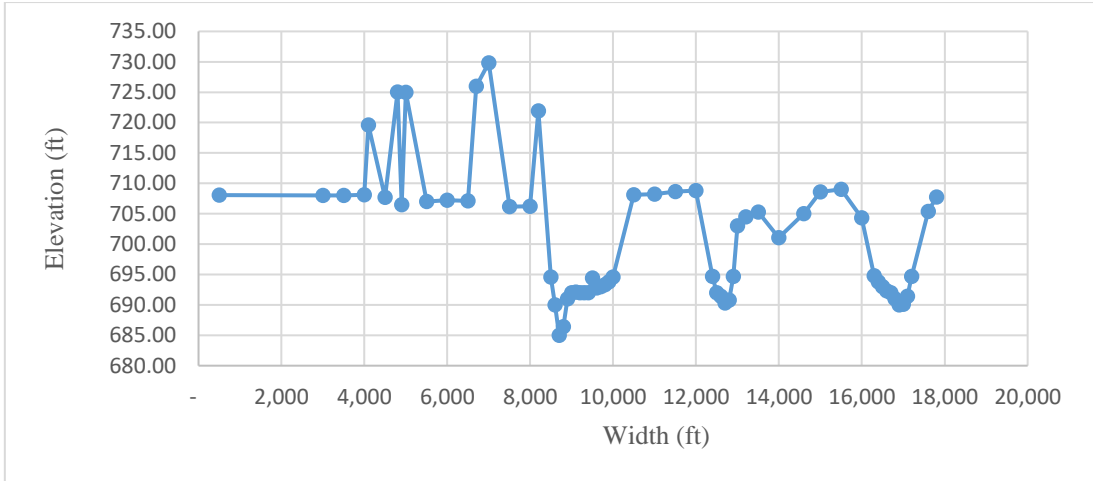


Figure C.1: Ravi Section at RD 81500

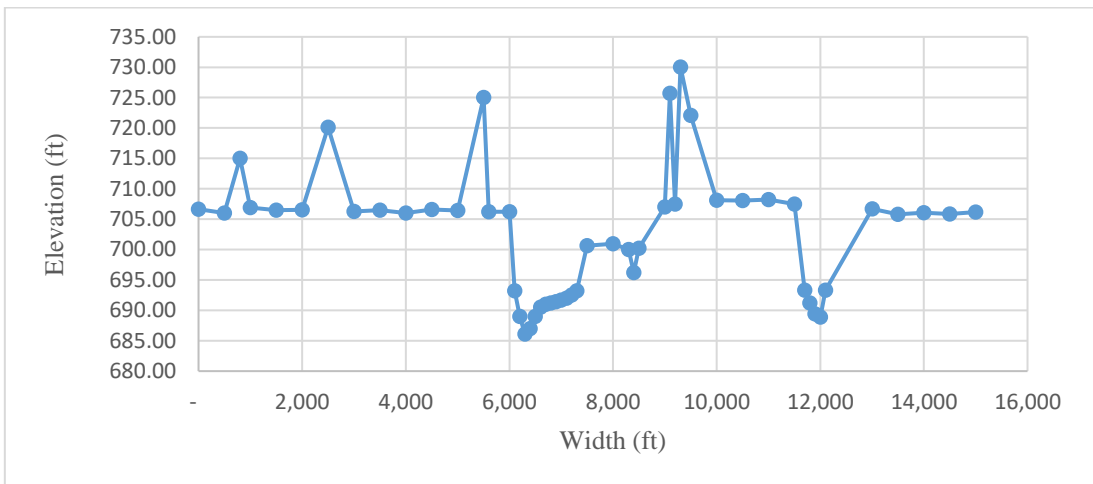


Figure C.2: Ravi Section at RD 79000

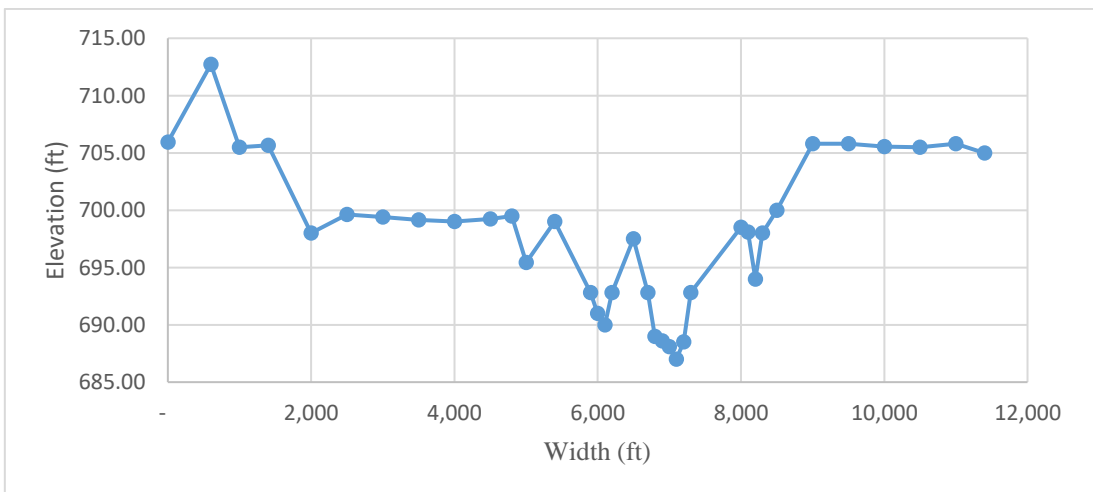


Figure C.3: Ravi Section at RD 77000

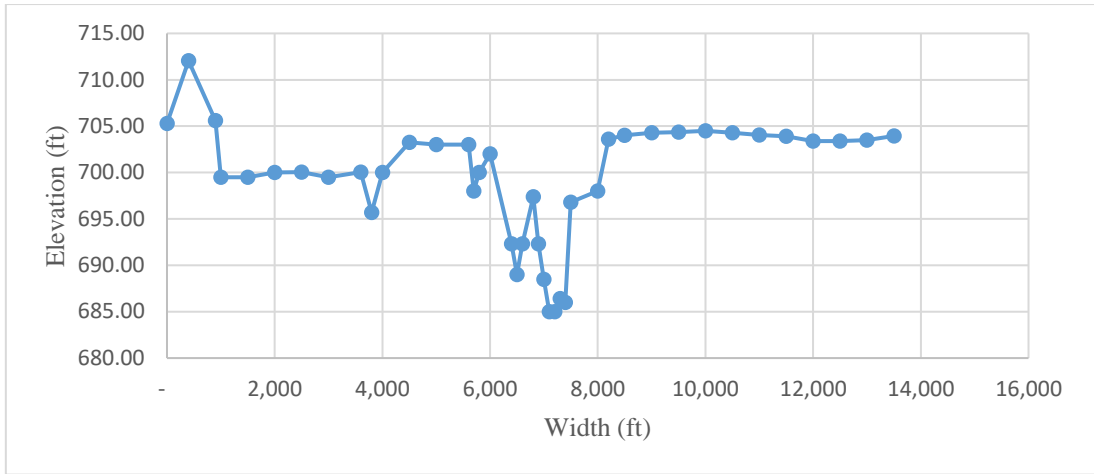


Figure C.4: Ravi Section at RD 75000

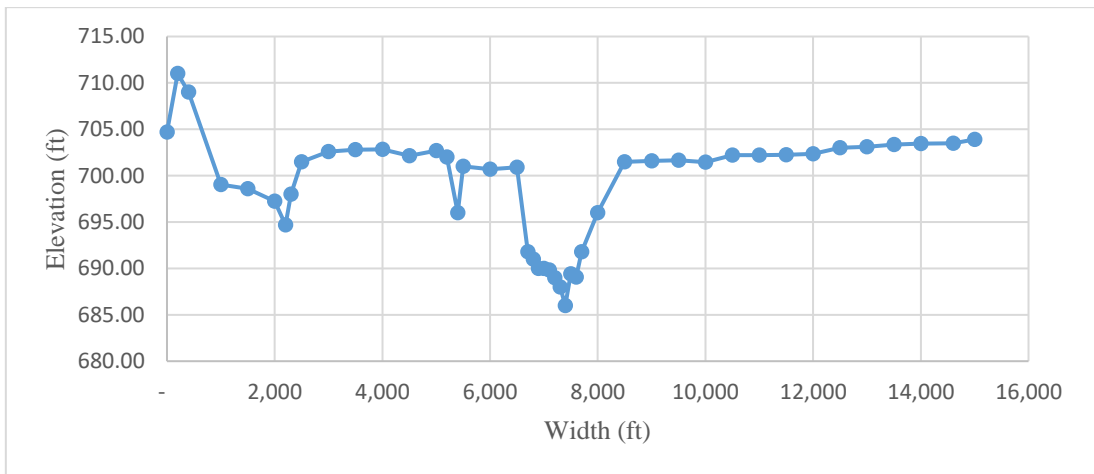


Figure C.5: Ravi Section at RD 73000

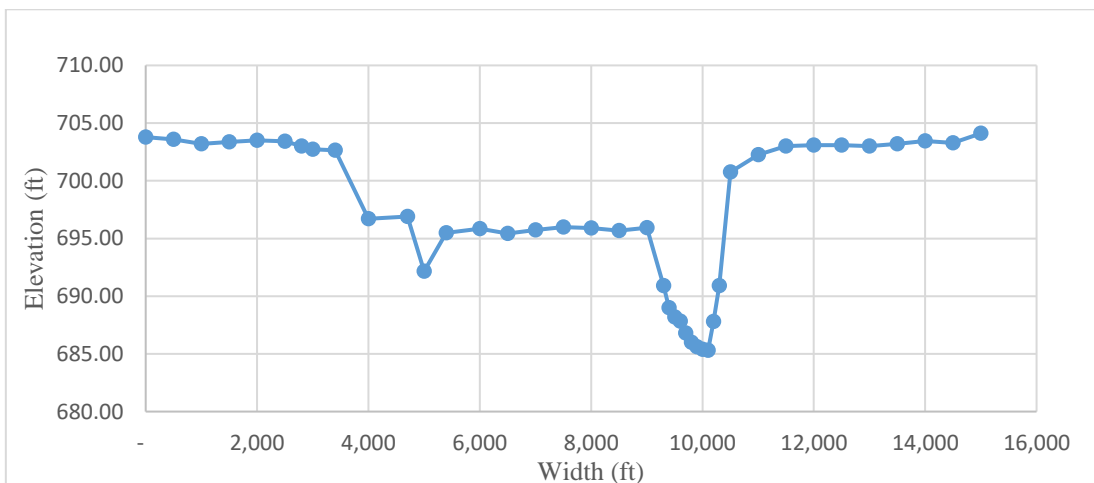


Figure C.6: Ravi Section at RD 67000

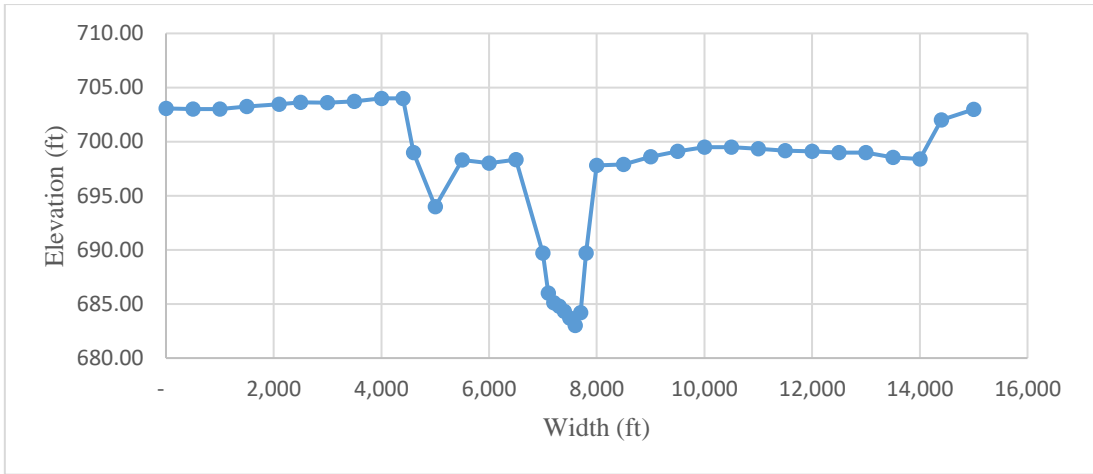


Figure C.7: Ravi Section at RD 64000

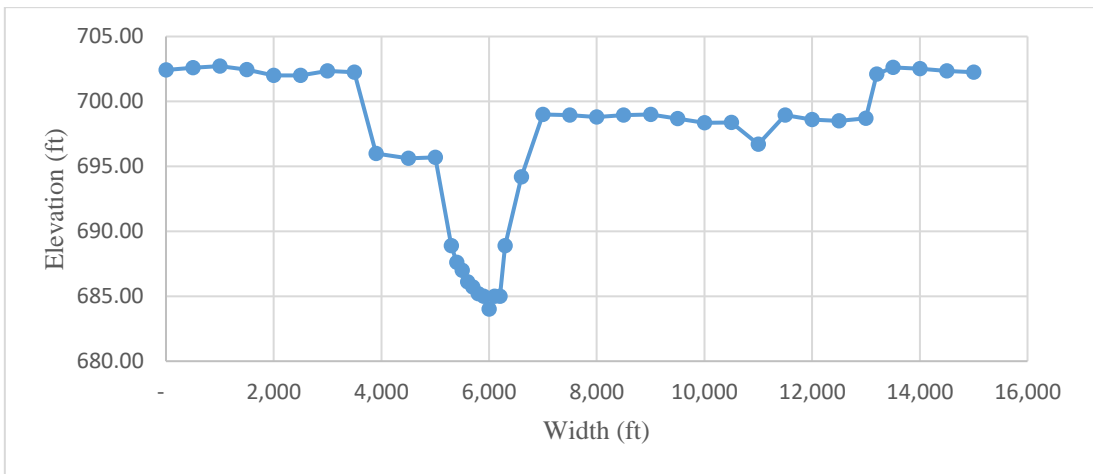


Figure C.8: Ravi Section at RD 61000

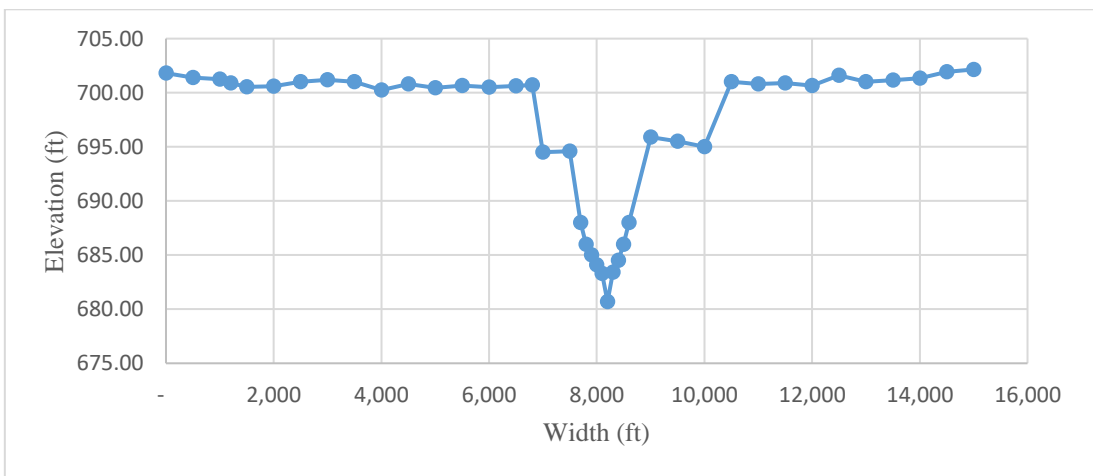


Figure C.9: Ravi Section at RD 55000

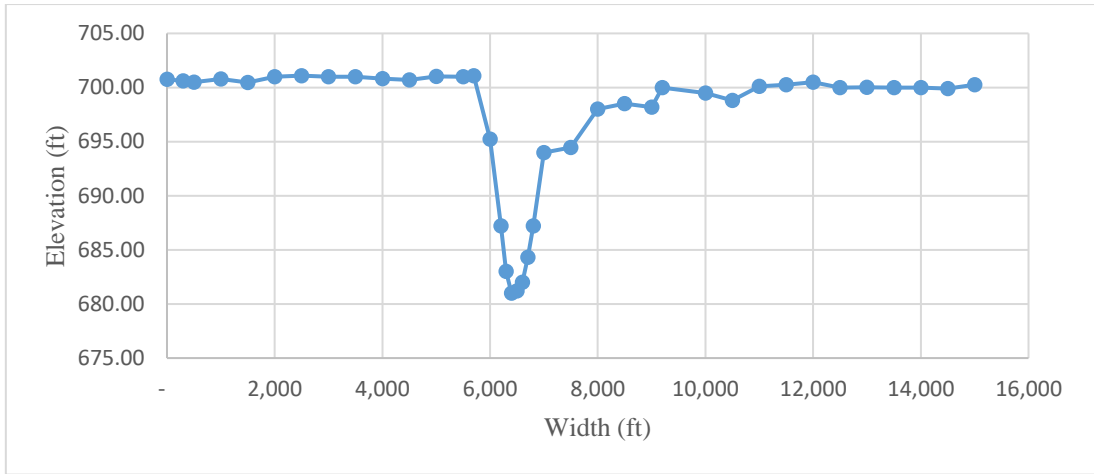


Figure C.10: Ravi Section at RD 52000

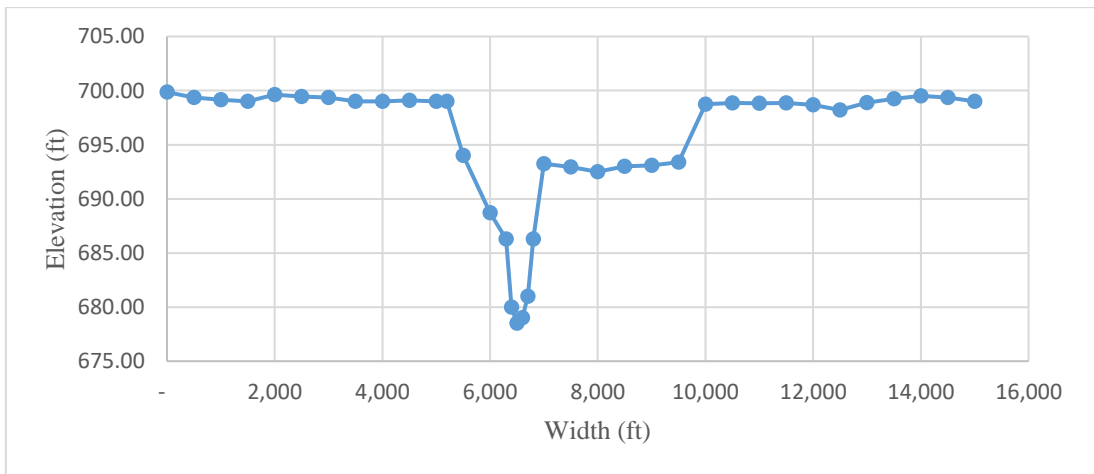


Figure C.11: Ravi section at RD 49000

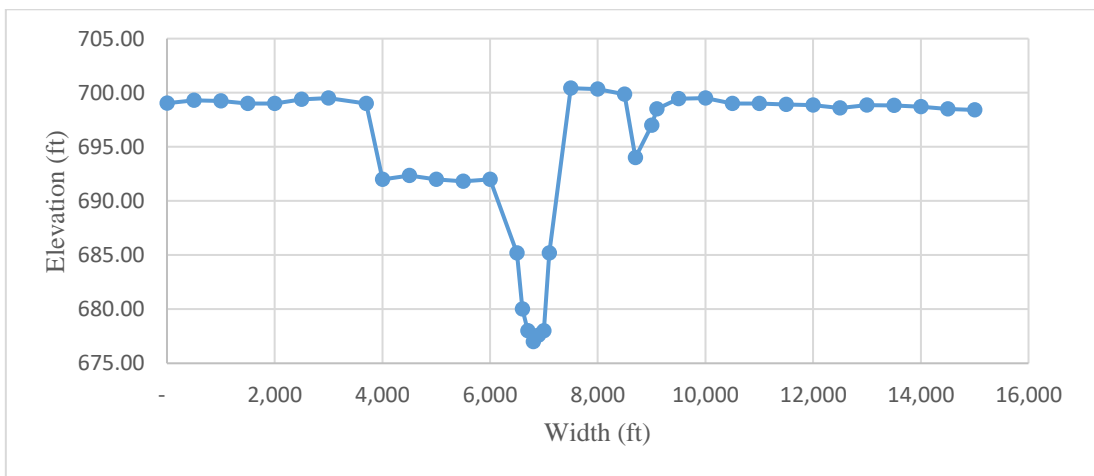


Figure C.12: Ravi Section at RD 46000

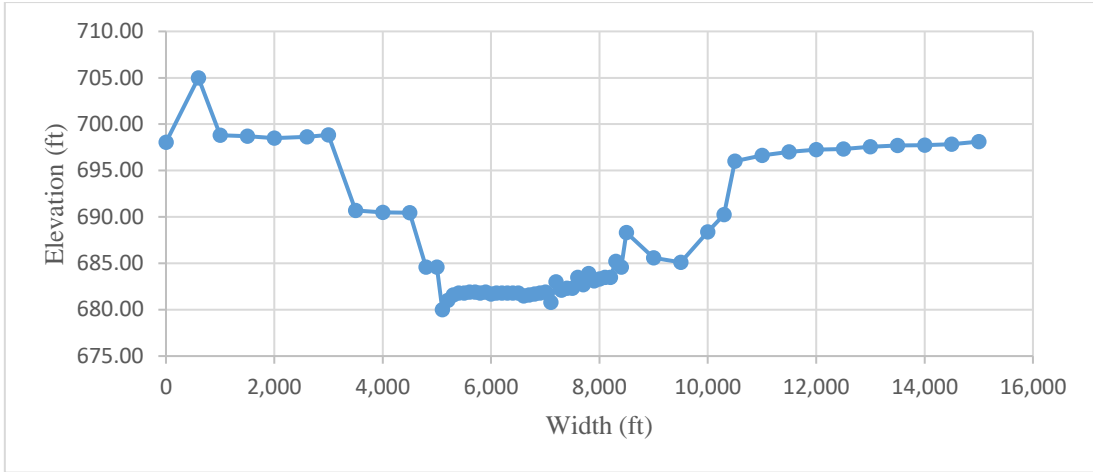


Figure C.13: Ravi Section at RD 42000

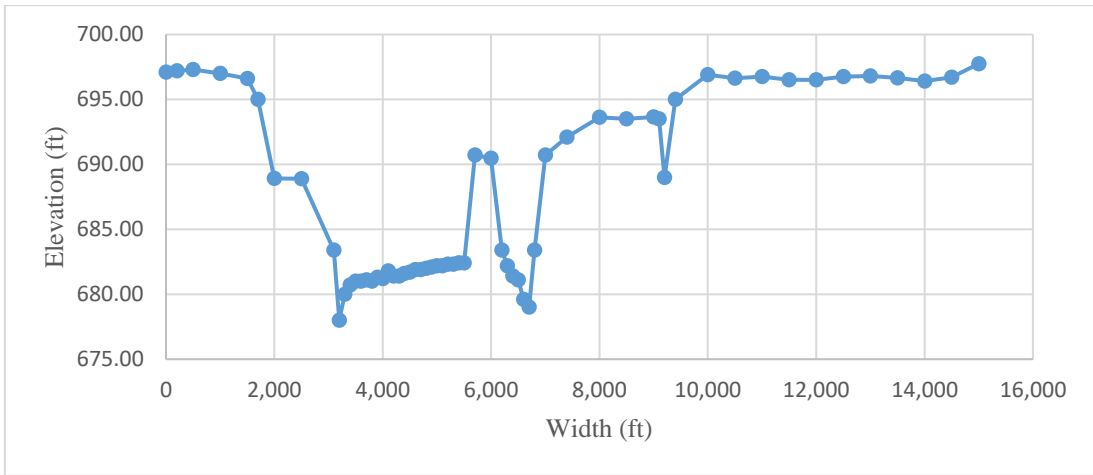


Figure C.14: Ravi Section at RD 39000

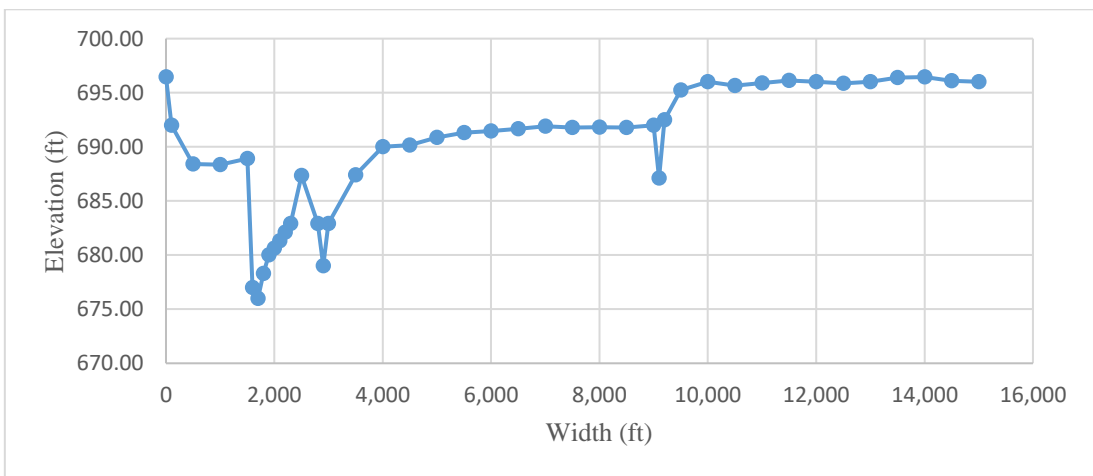


Figure C.15: Ravi Section at RD 37000

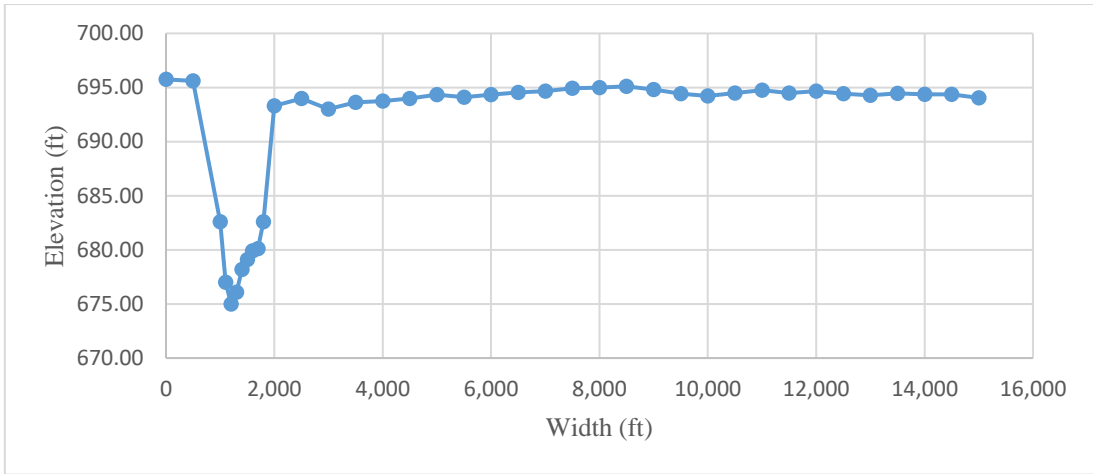


Figure C.16: Ravi Section at RD 35000

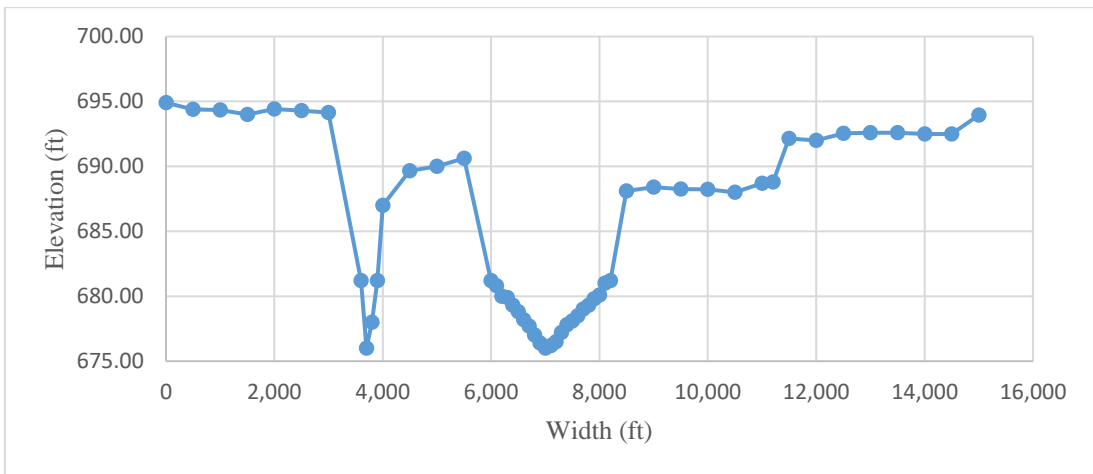


Figure C.17: Ravi Section at RD 31000

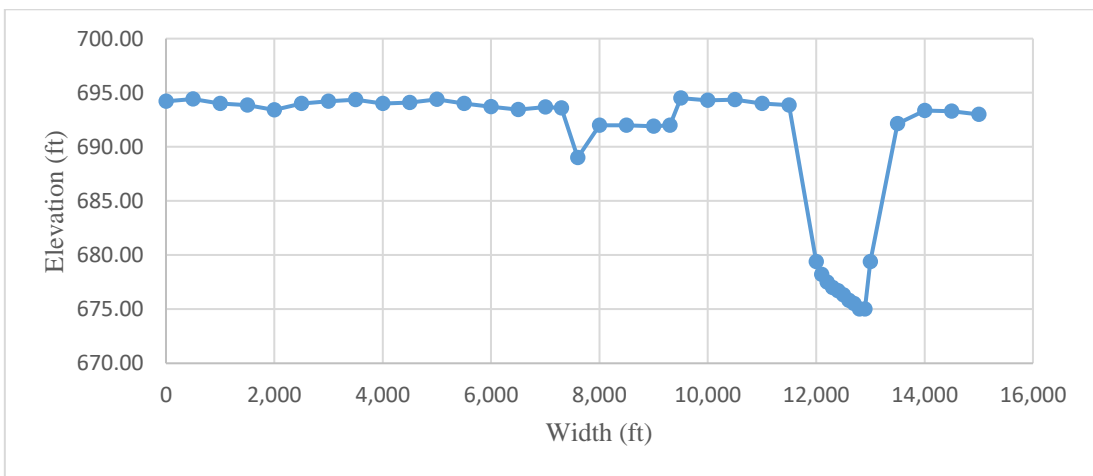


Figure C.18: Ravi Section at RD 28000

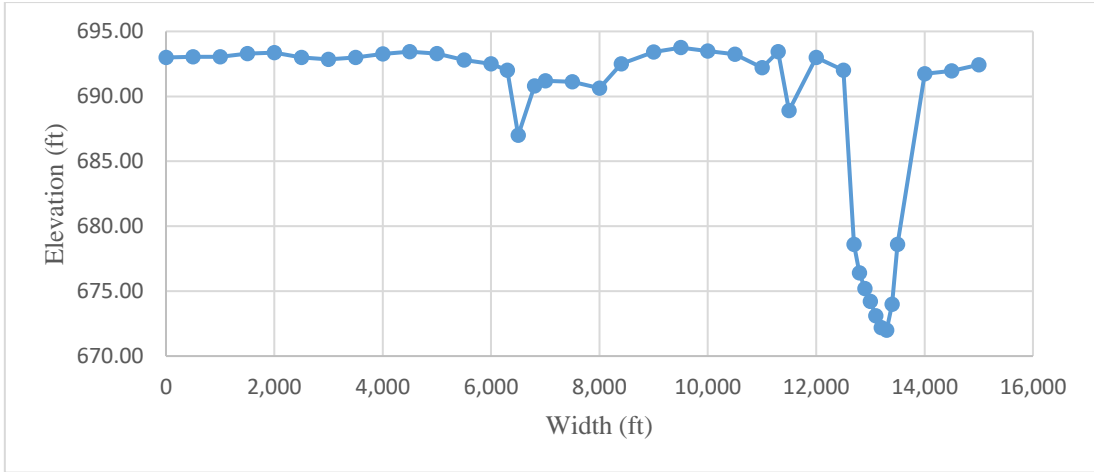


Figure C.19: Ravi Section at RD 25000

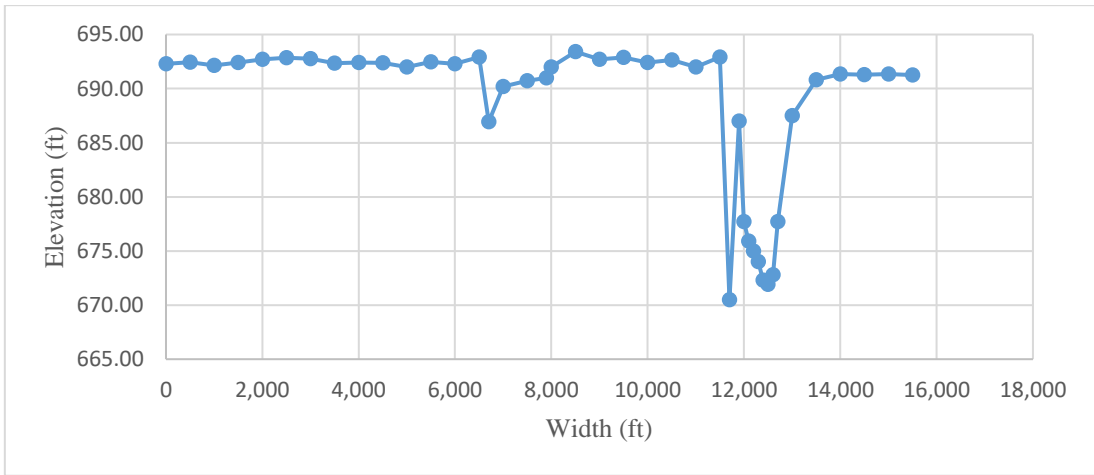


Figure C.20: Ravi Section at RD 22000

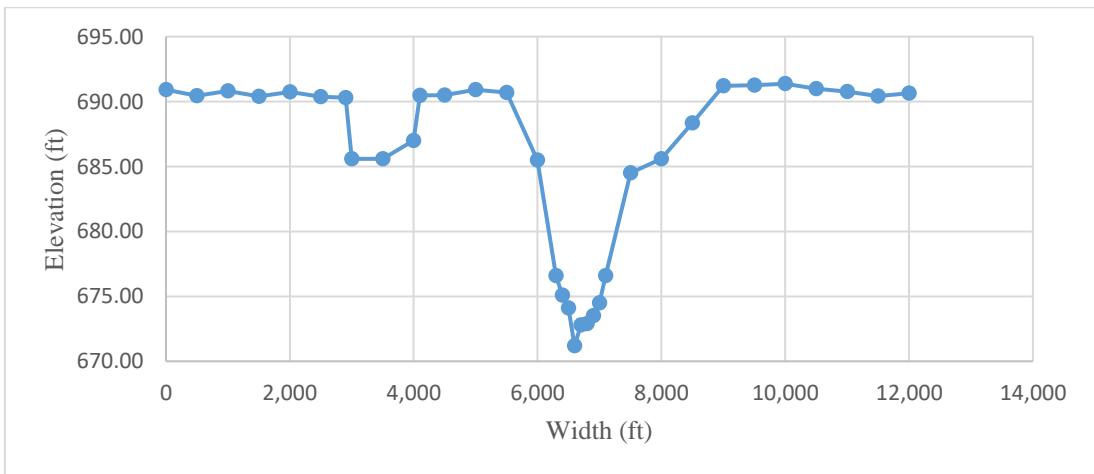


Figure C.21: Ravi Section at RD 16000

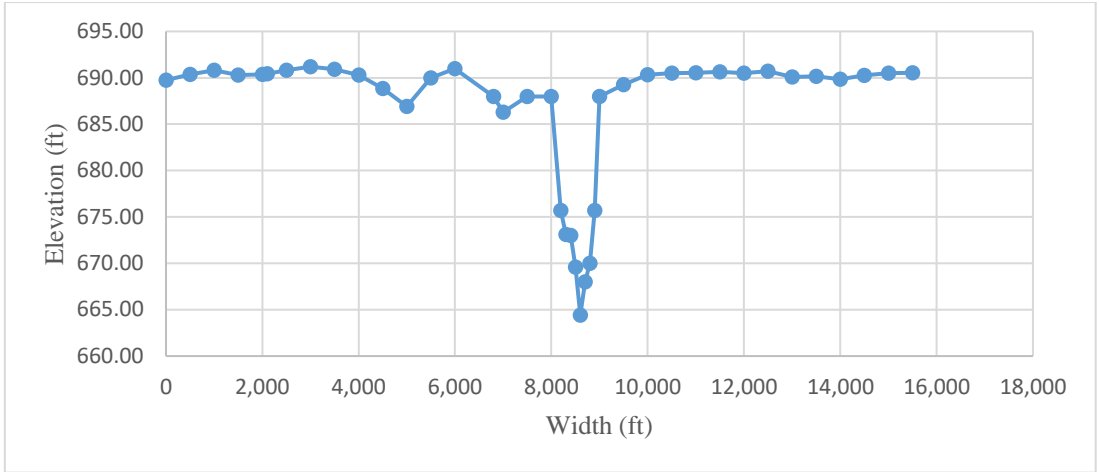


Figure C.22: Ravi Section at RD 13000

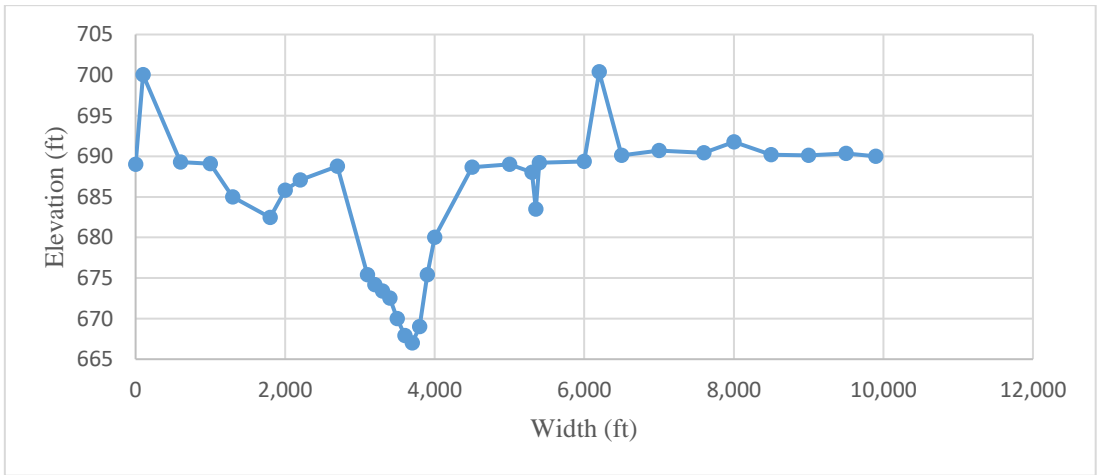


Figure C.23: Ravi Section at RD 7000

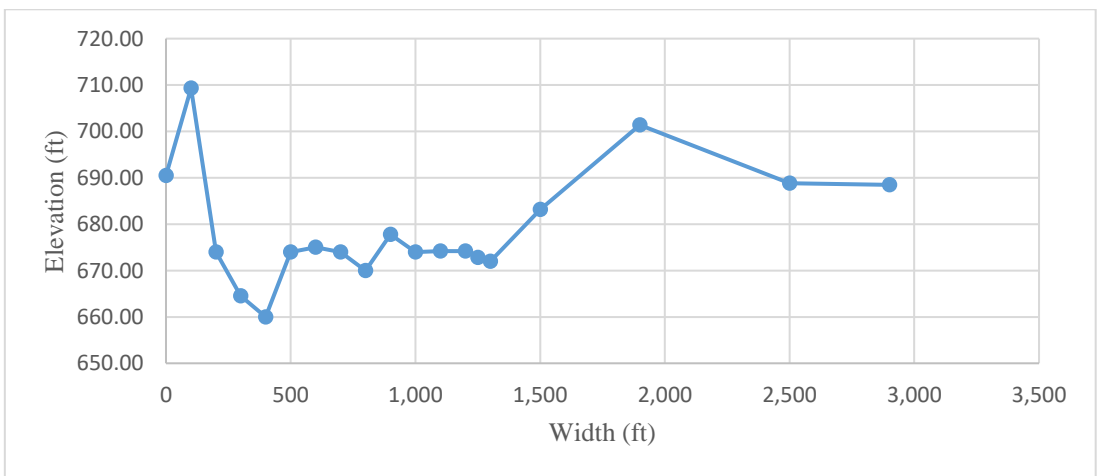


Figure C.24: Ravi Section at RD 500

Annexure D

River Section Shahdra to head Balloki

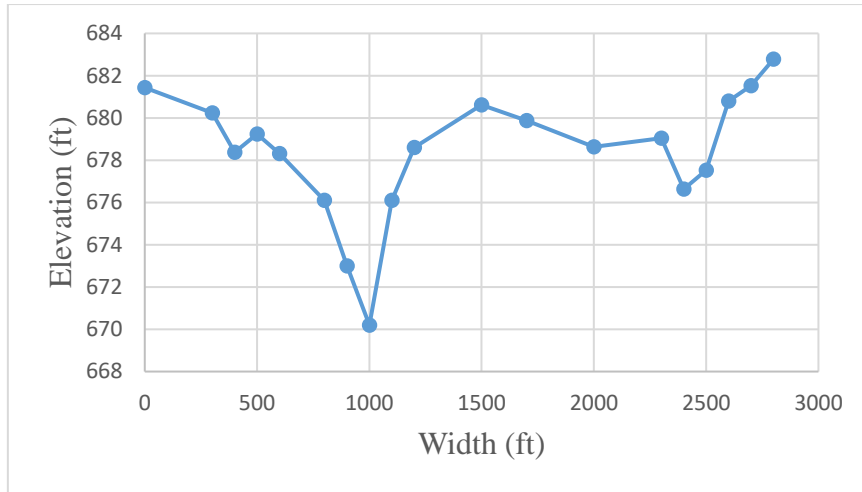


Figure D.1: Ravi Section at RD 4000

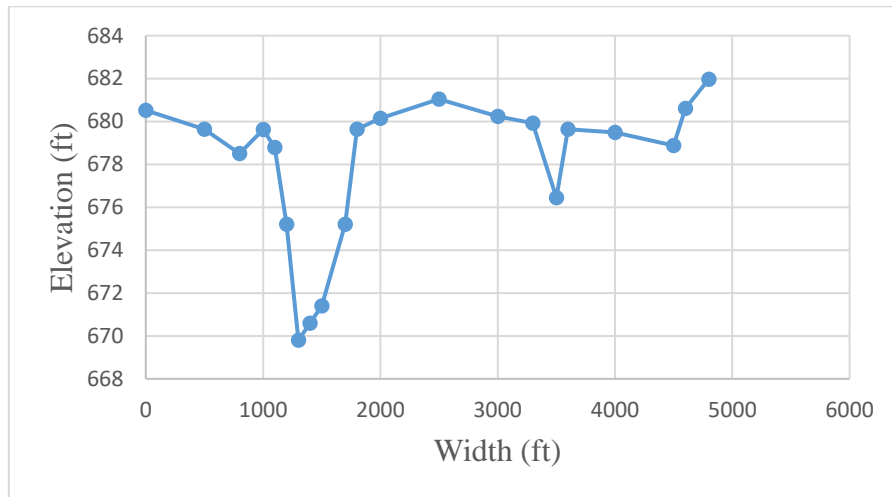


Figure D.2: Ravi Section at RD 8000

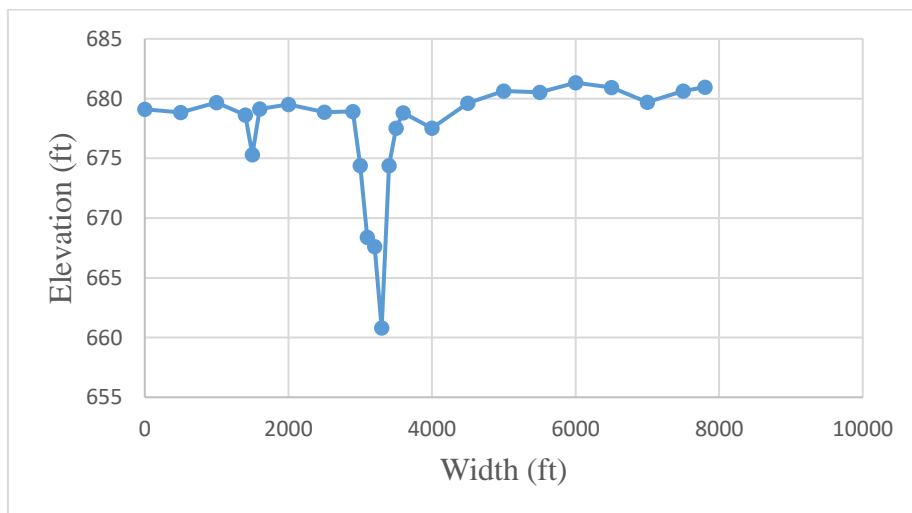


Figure D.3: Ravi Section at RD 11000

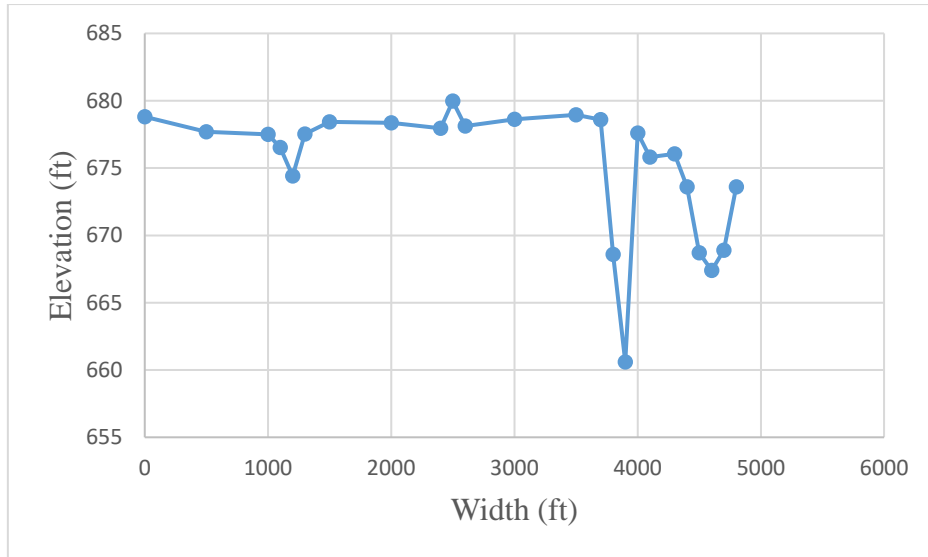


Figure D.4: Ravi Section at RD 14000

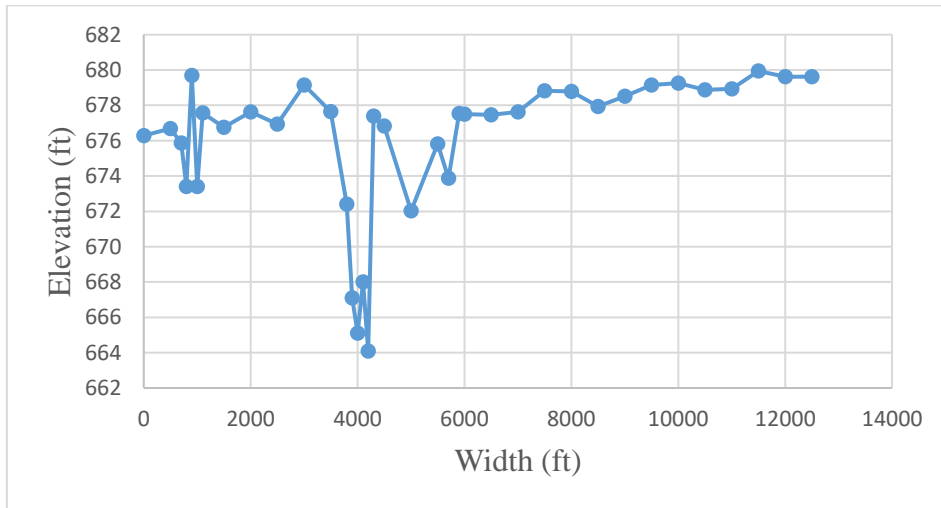


Figure D.5: Ravi Section at RD 21000

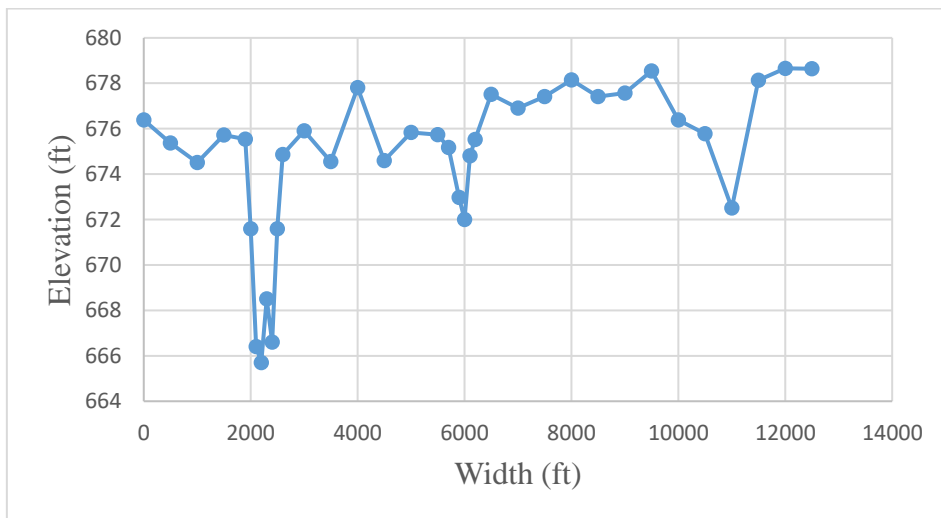


Figure D.6: Ravi Section at RD 24000

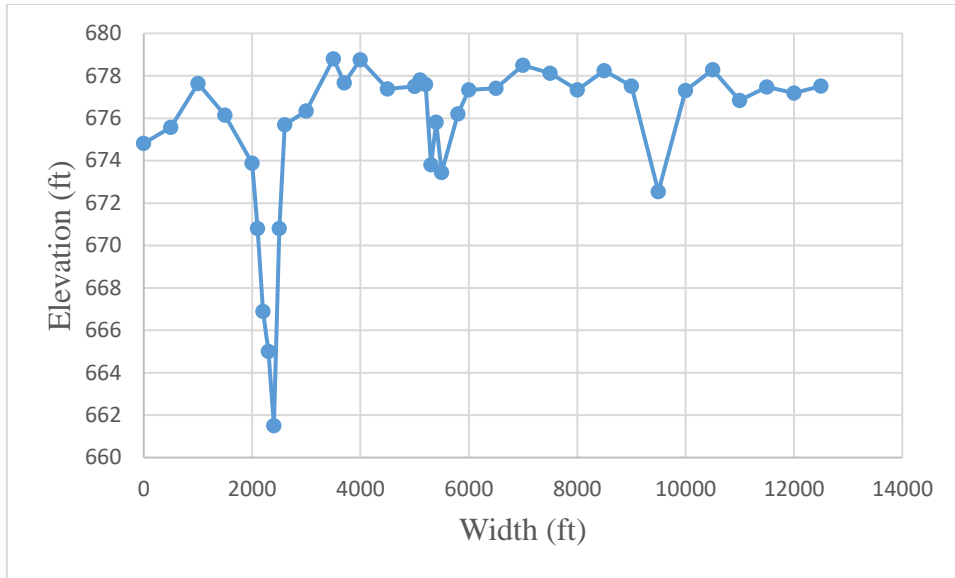


Figure D.7: Ravi Section at RD 27000

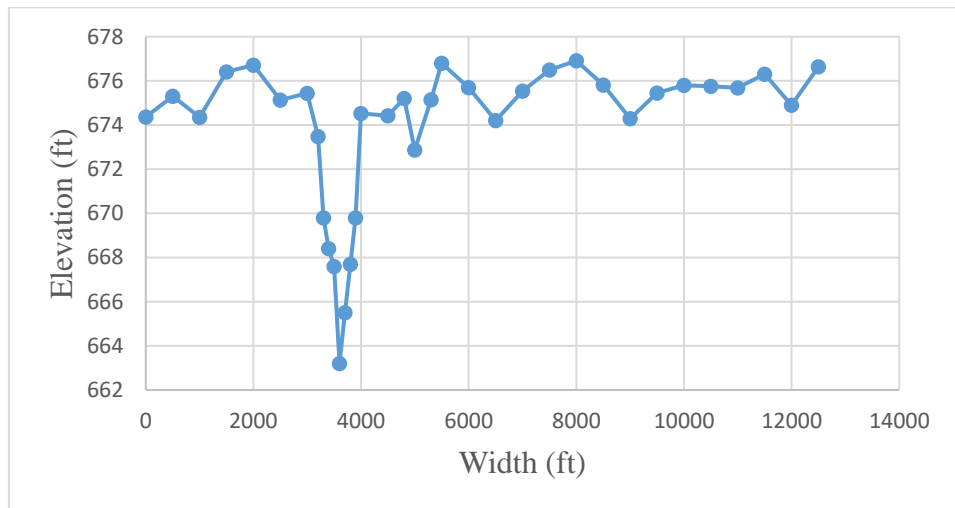


Figure D.8: Ravi Section at RD 30500

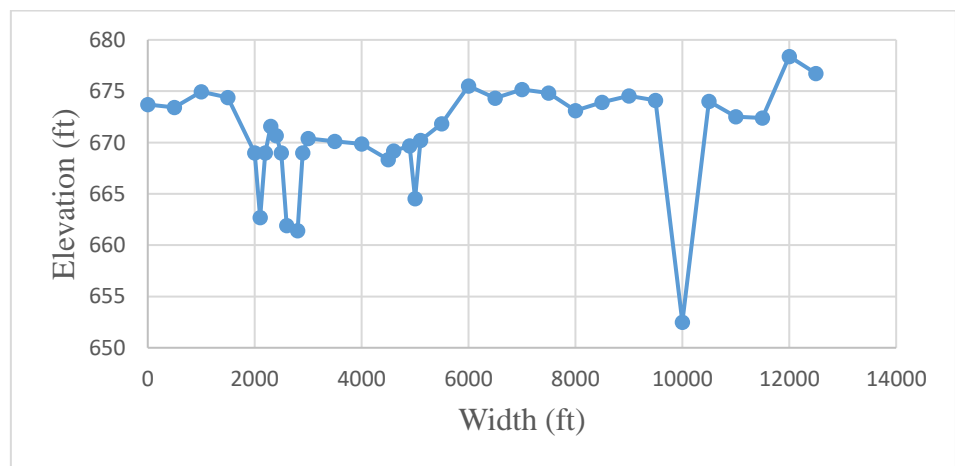


Figure D.9: Ravi Section at RD 33000

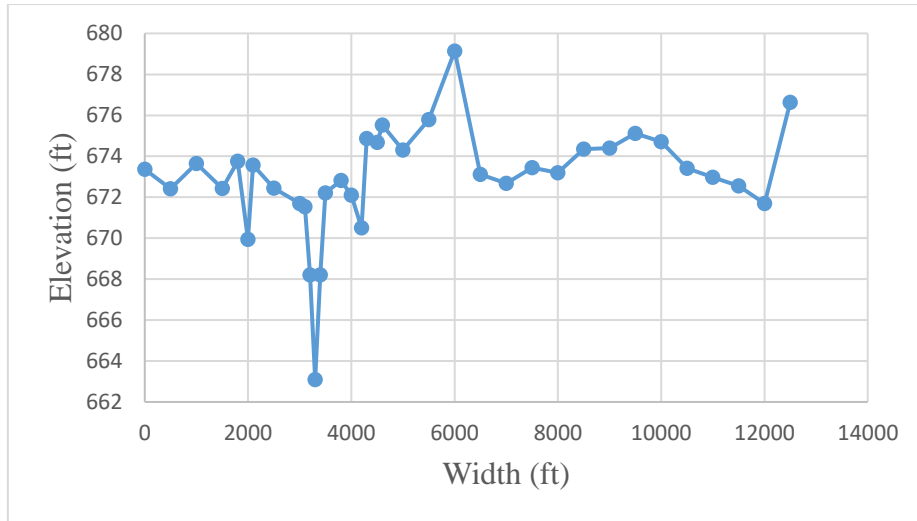


Figure D.10: Ravi Section at RD 36000

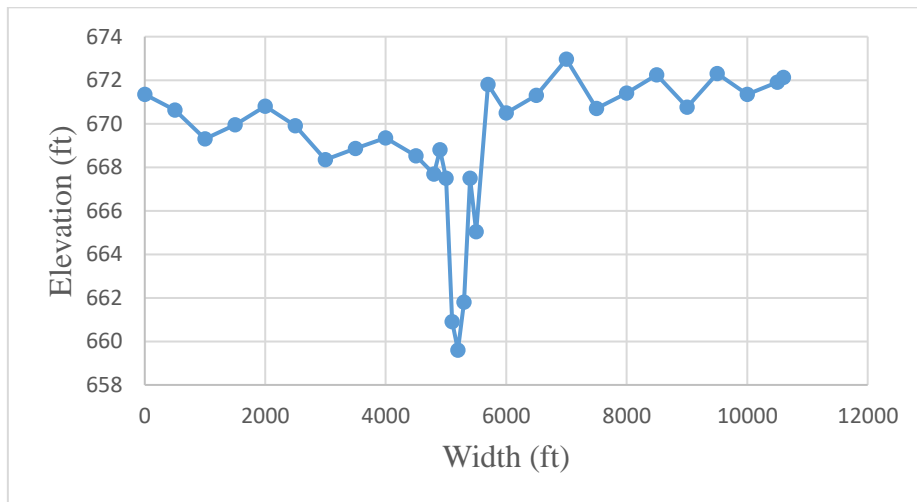


Figure D.11: Ravi Section at RD 39000

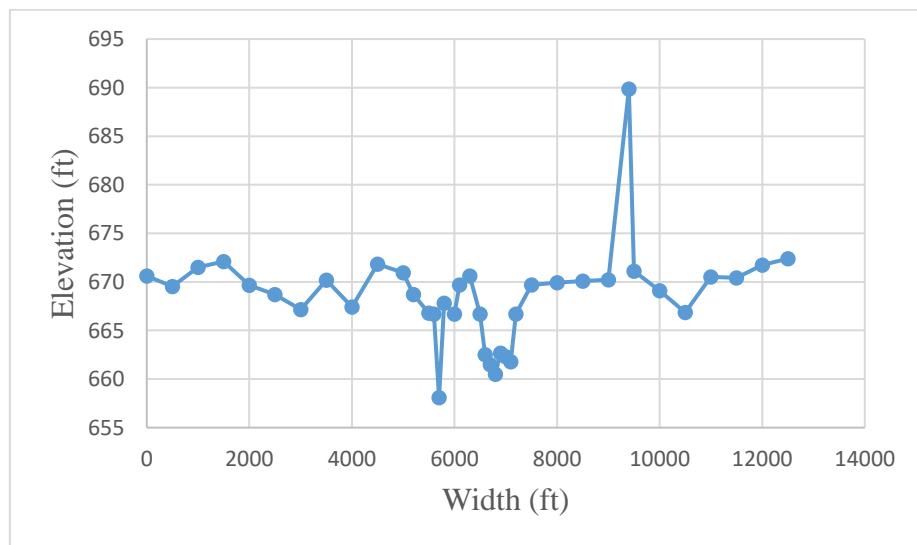


Figure D.12: Ravi Section at RD 41500

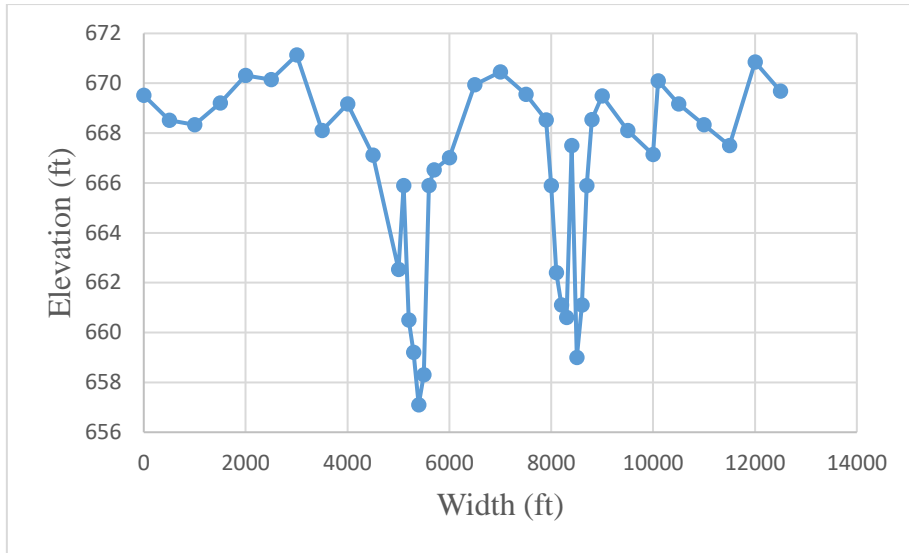


Figure D.13: Ravi Section at RD 44000

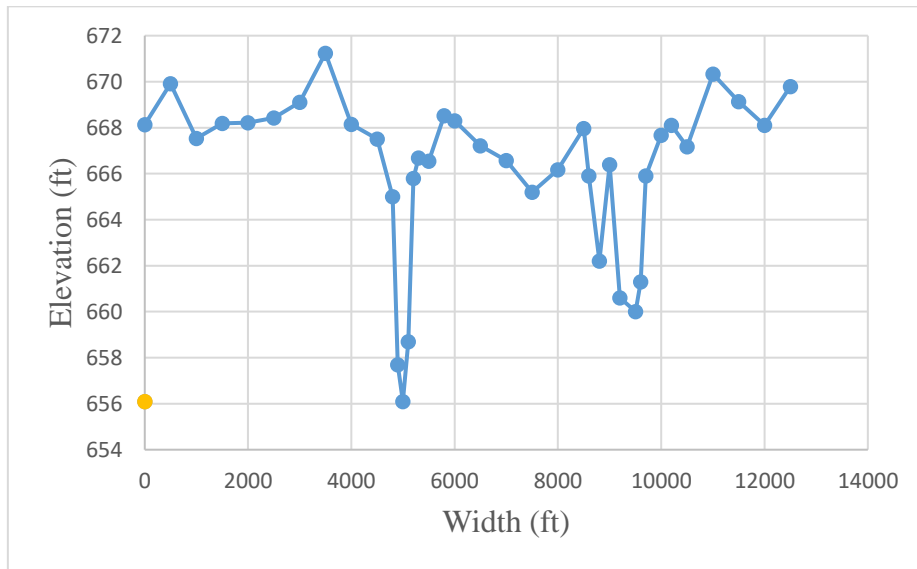


Figure D.14: Ravi Section at RD 47000

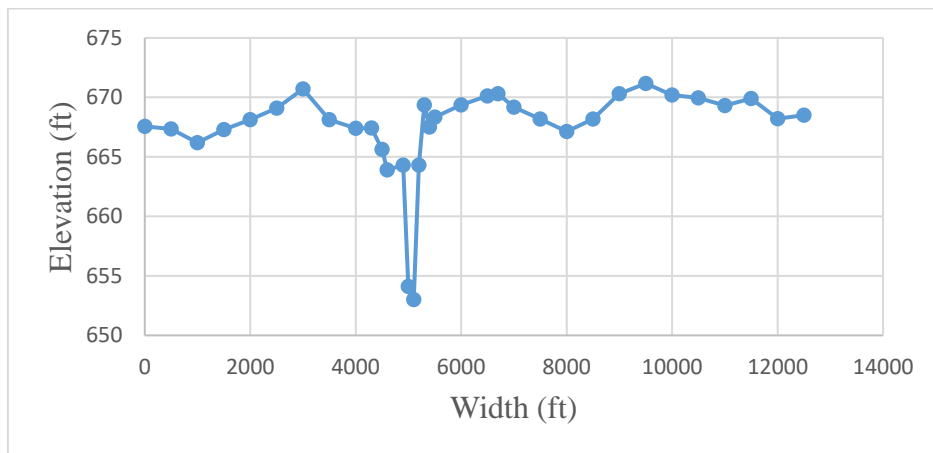


Figure D.15: Ravi Section at RD 50500

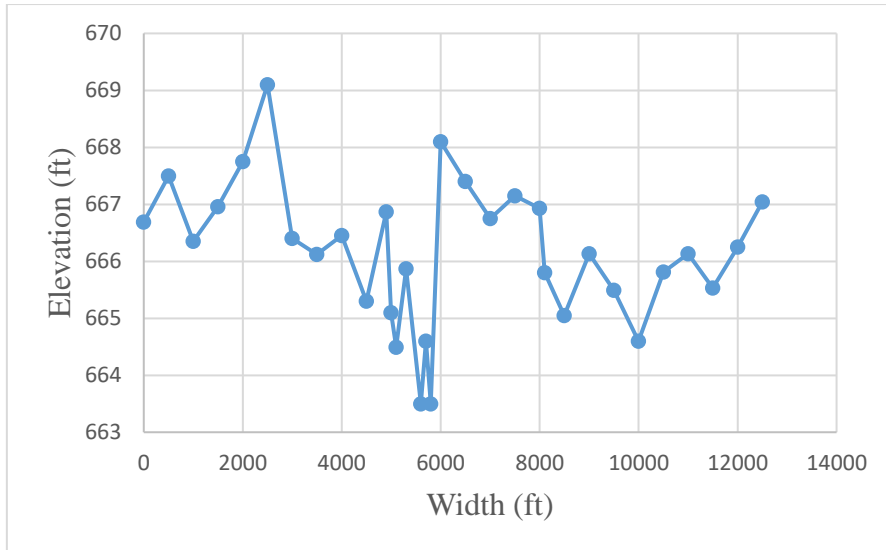


Figure D.16: Ravi Section at RD 53000

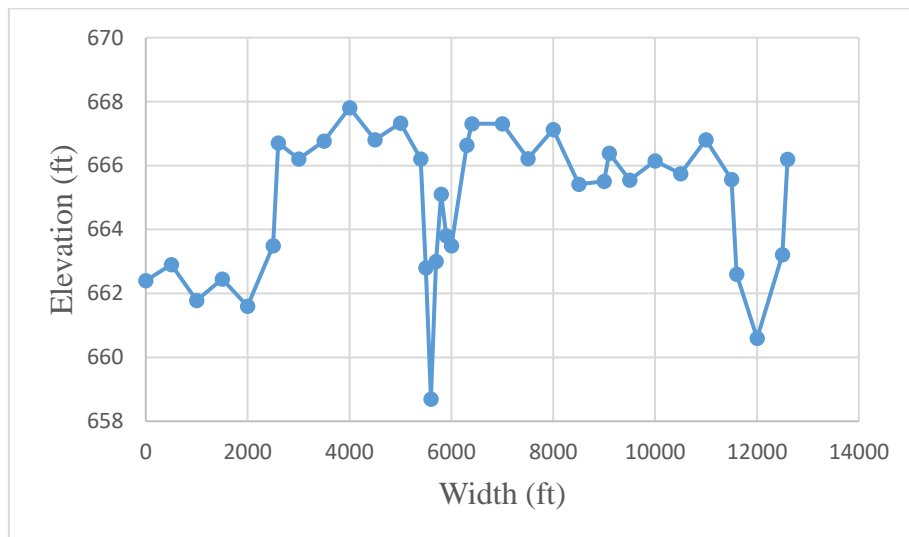


Figure D.17: Ravi Section at RD 55500

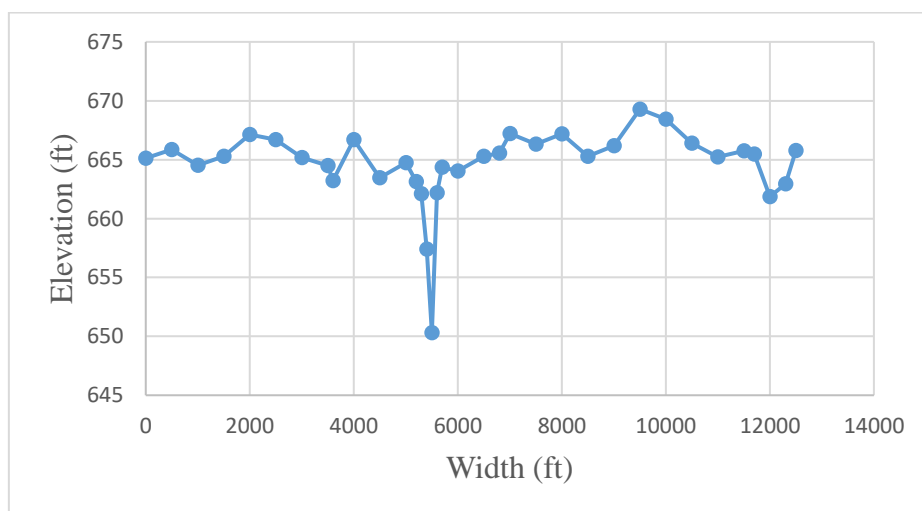


Figure D.18: Ravi Section at RD 58000

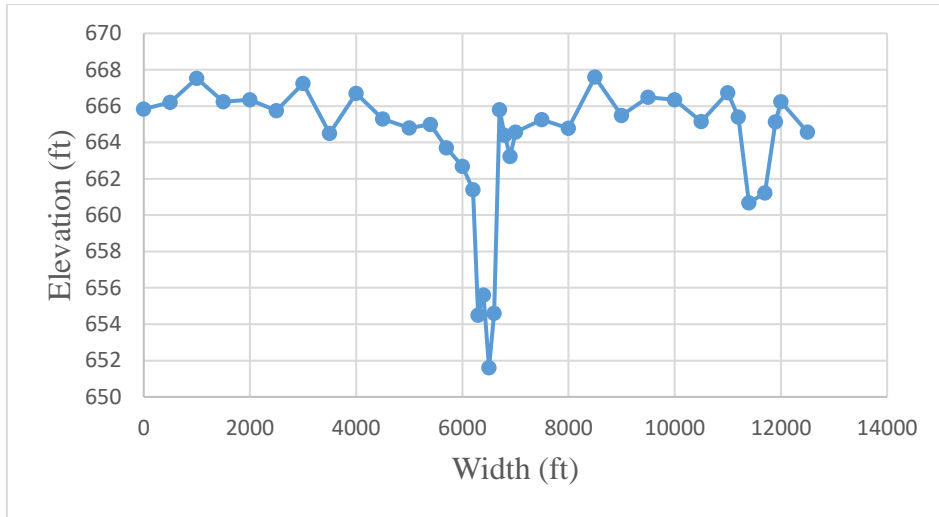


Figure D.19: Ravi Section at RD 60500

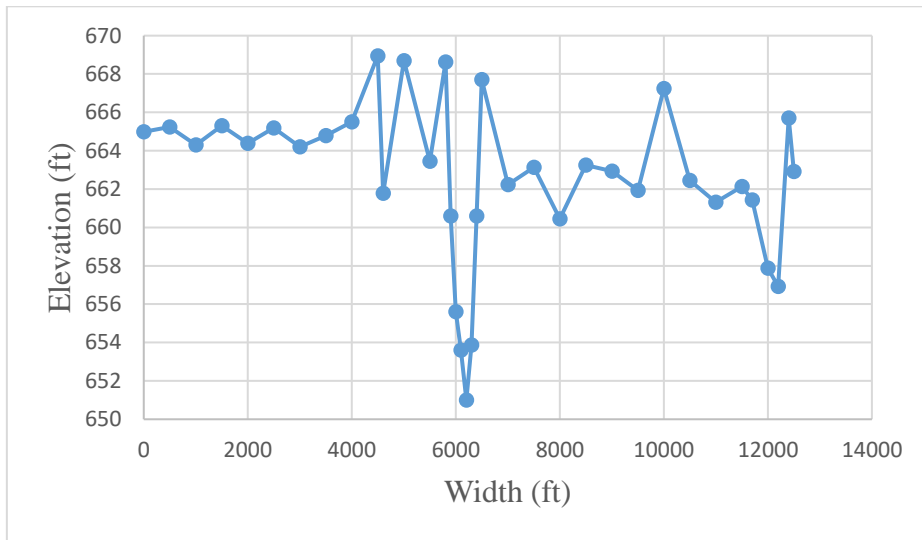


Figure D.20: Ravi Section at RD 63000

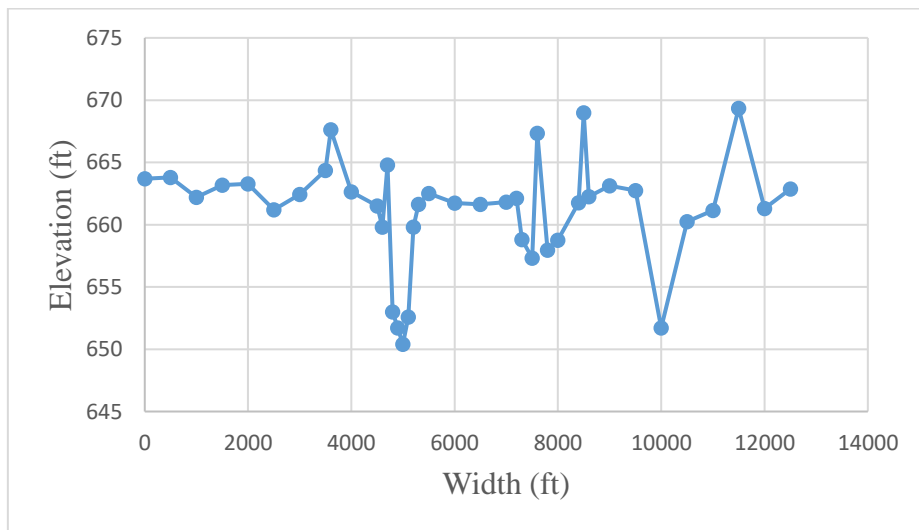


Figure D.21: Ravi Section at RD 66000

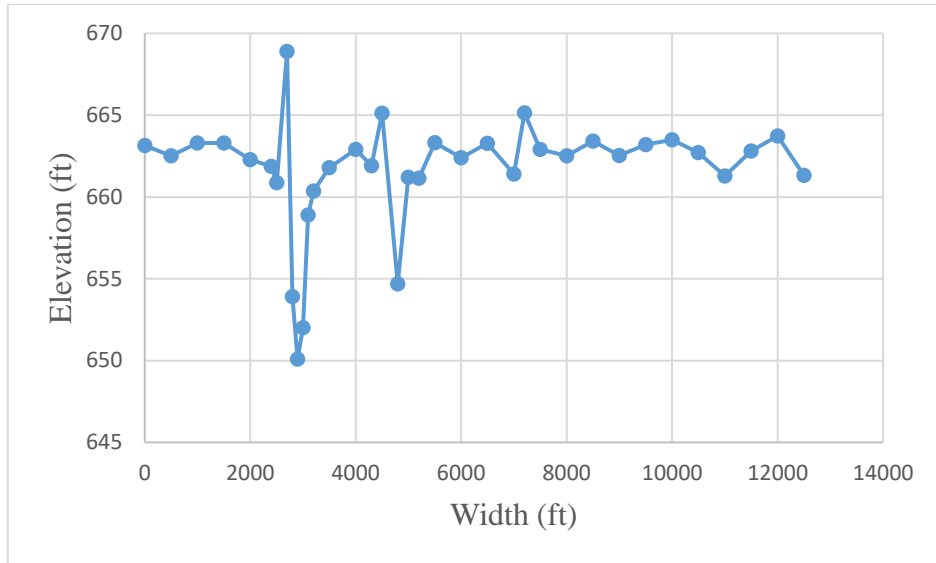


Figure D.22: Ravi Section at RD 69000

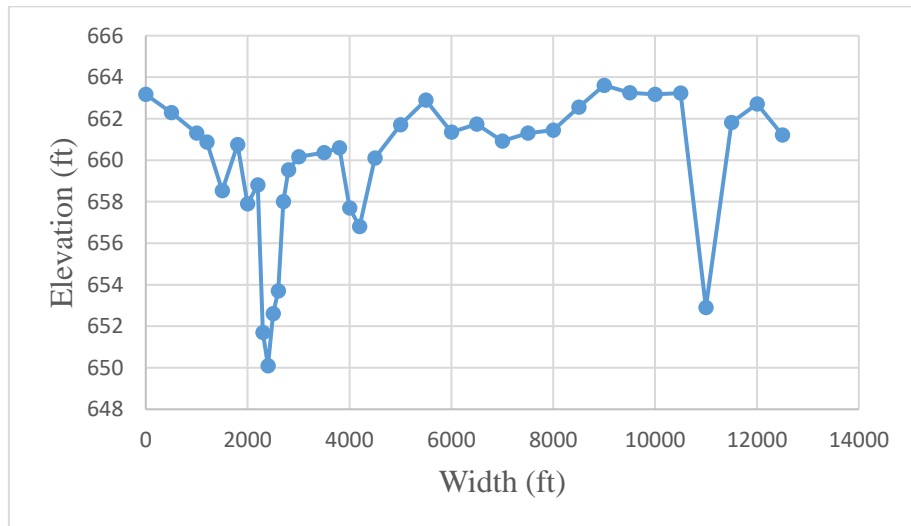


Figure D.23: Ravi Section at RD 72000

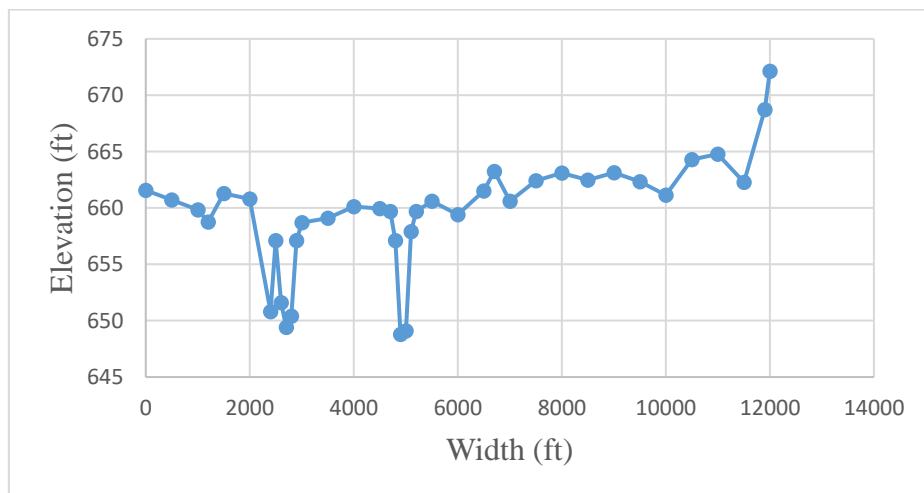


Figure D.24: Ravi Section at RD 75000

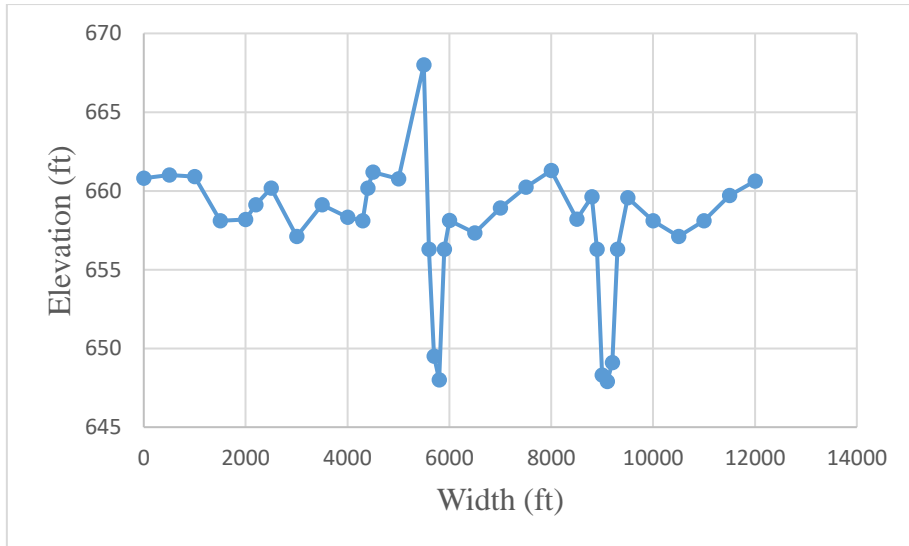


Figure D.25: Ravi Section at RD80500

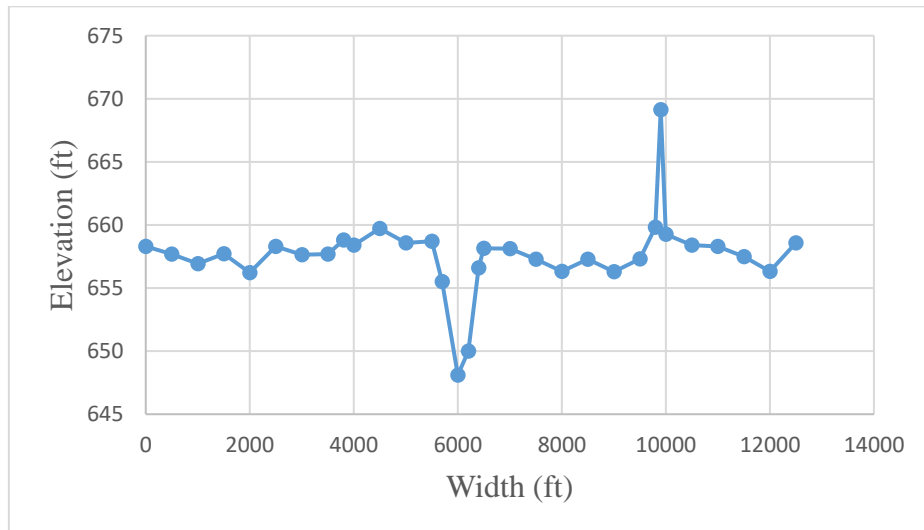


Figure D.26: Ravi Section at RD 83000

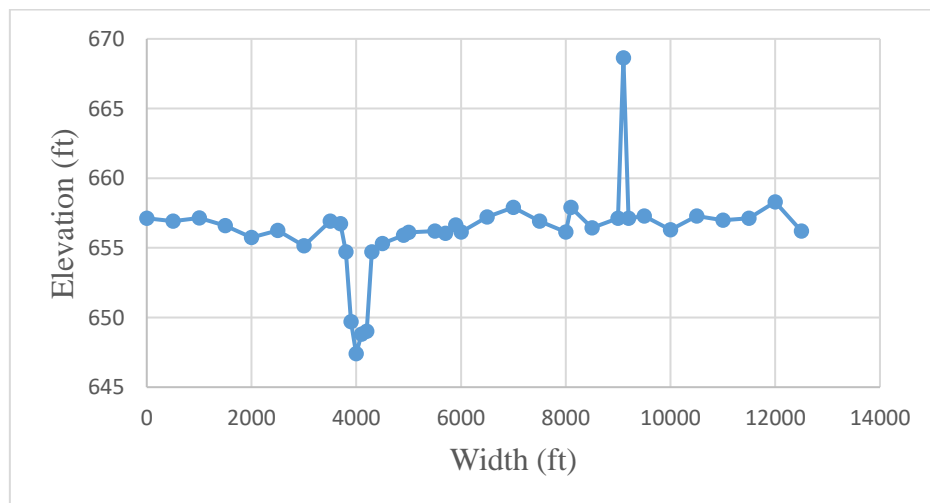


Figure D.27: Ravi Section at RD 86000

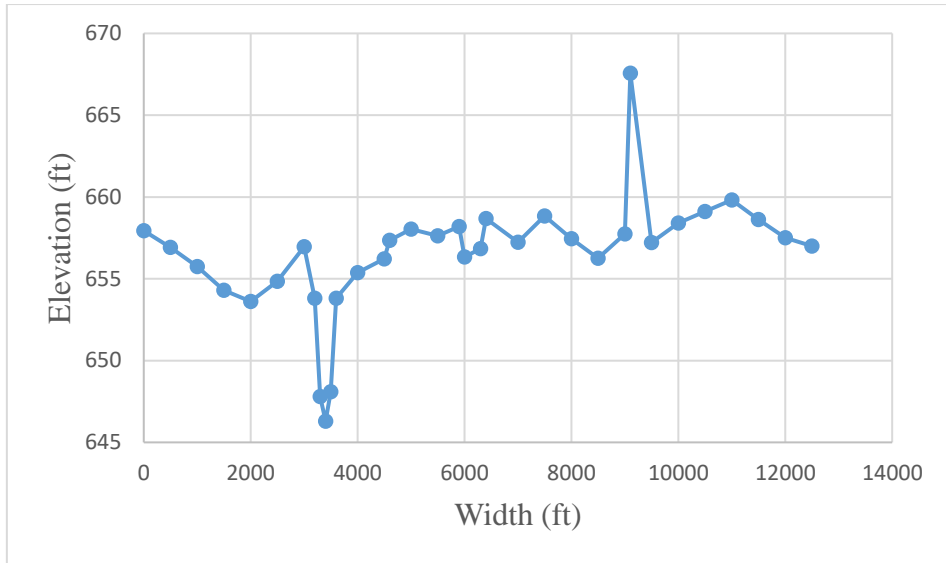


Figure D.28: Ravi Section at RD 89000

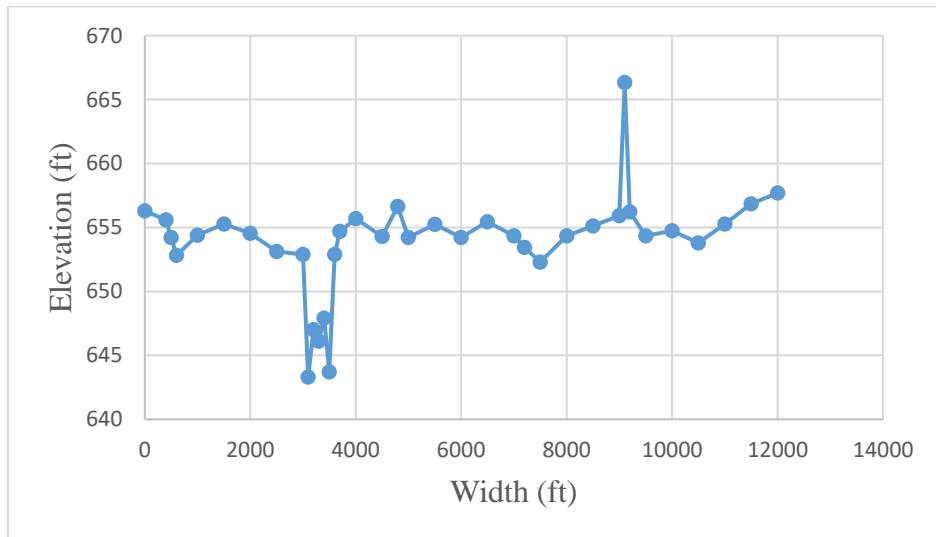


Figure D.29: Ravi Section at RD 92000

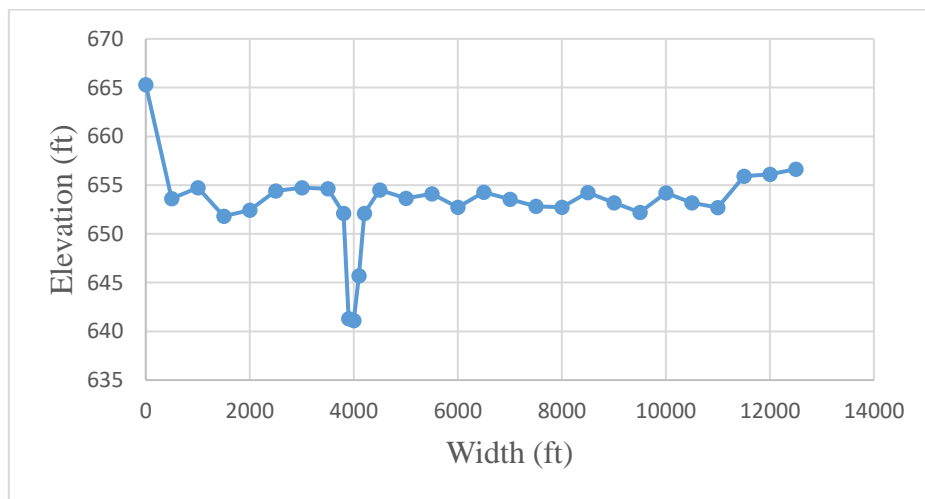


Figure D.30: Ravi Section at RD 95000

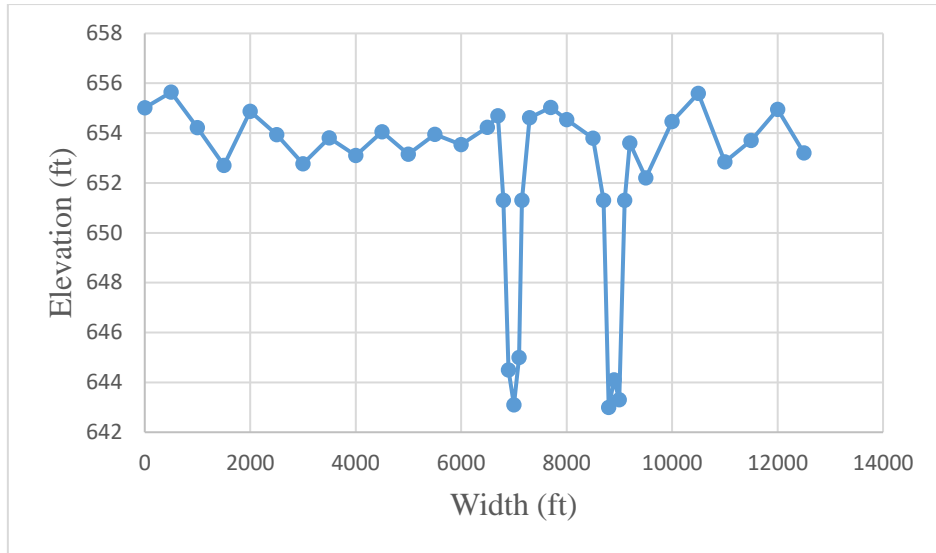


Figure D.31: Ravi Section at RD 98000

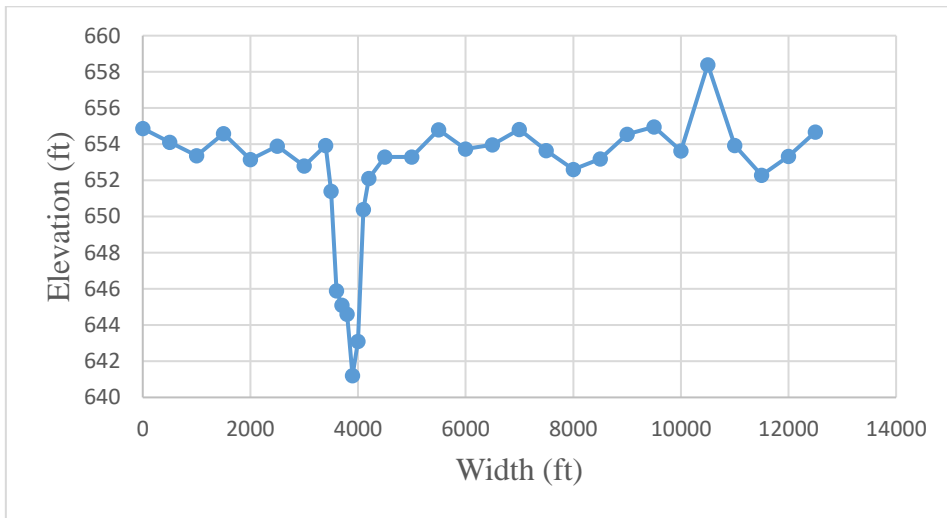


Figure D.32: Ravi Section at RD 111000

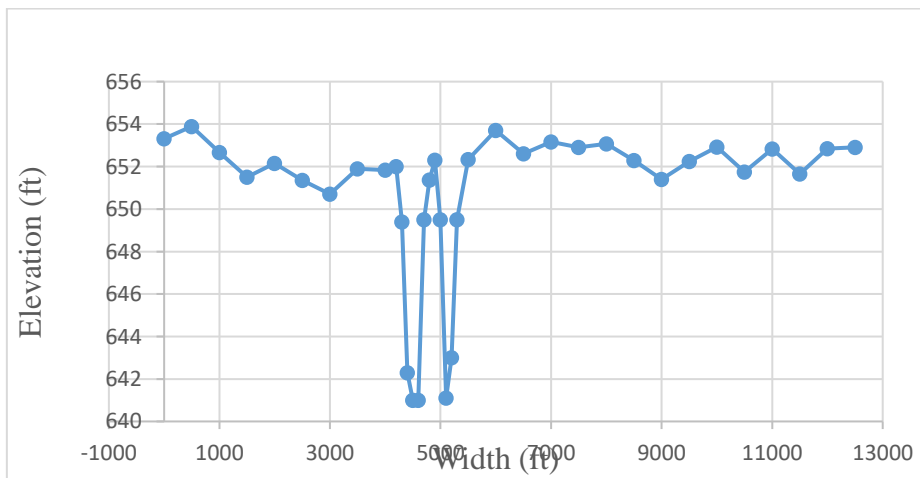


Figure D.33: Ravi Section at RD 114000

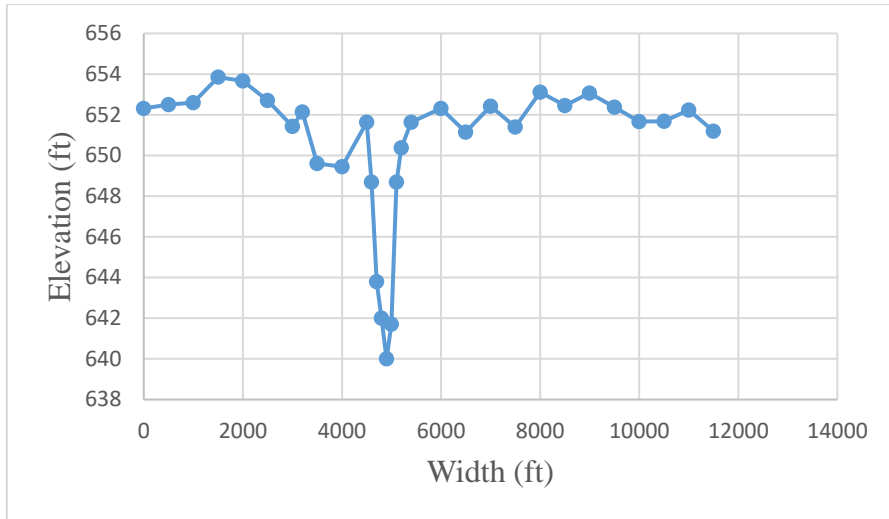


Figure D.34: Ravi Section at RD 117000

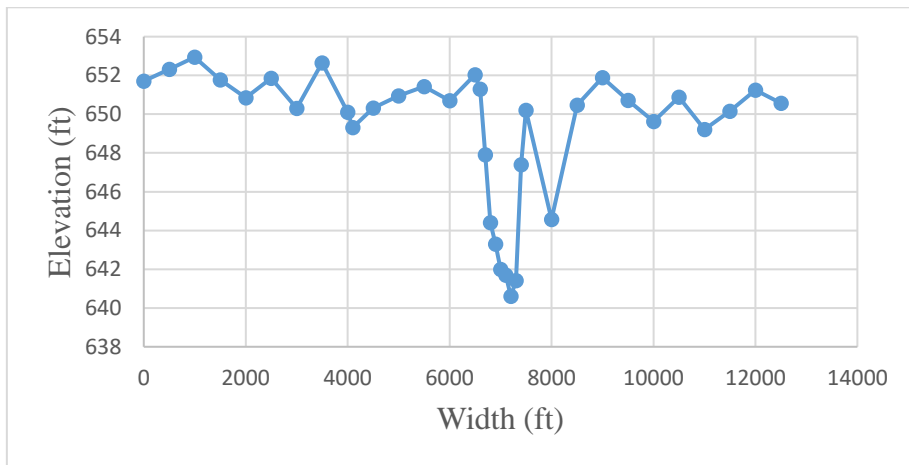


Figure D.35: Ravi Section at RD 123000

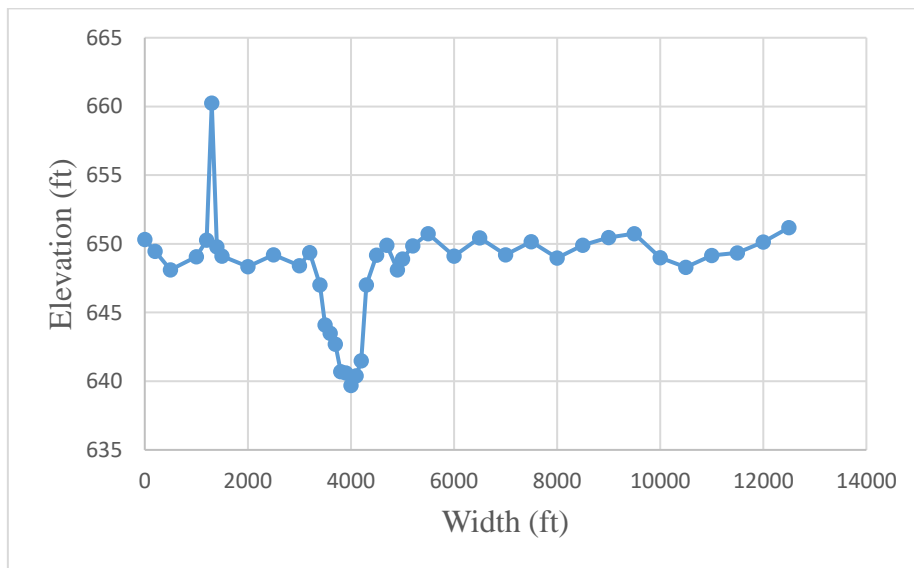


Figure D.36: Ravi Section at Rd 126000

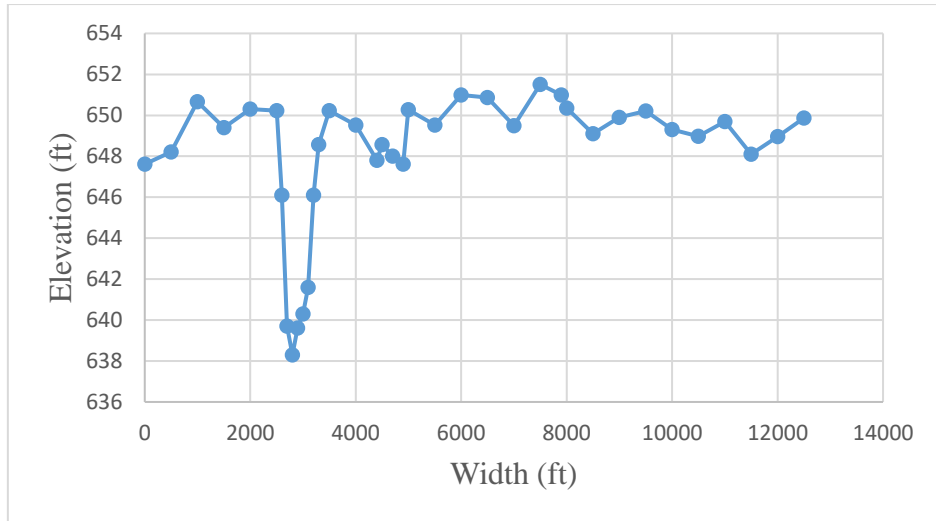


Figure D.37: Ravi Section at RD 129000

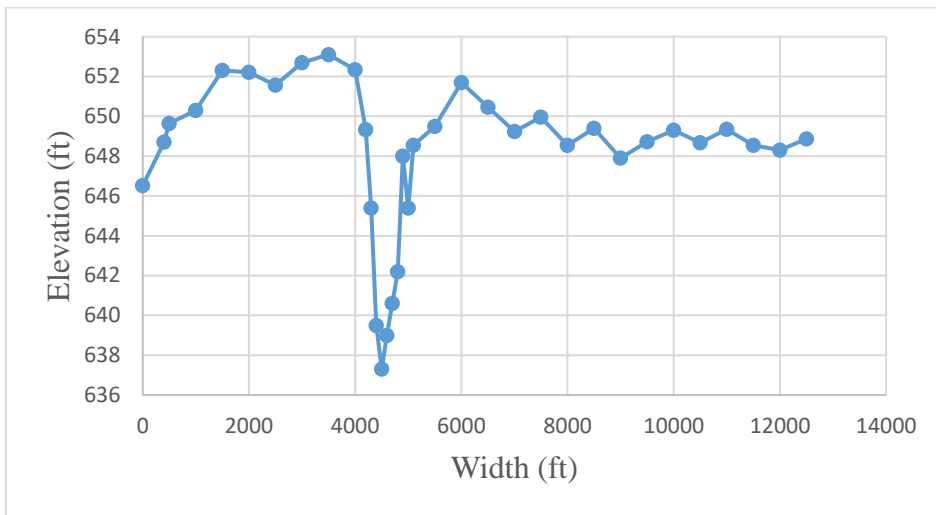


Figure D.38: Ravi Section at Rd 132000

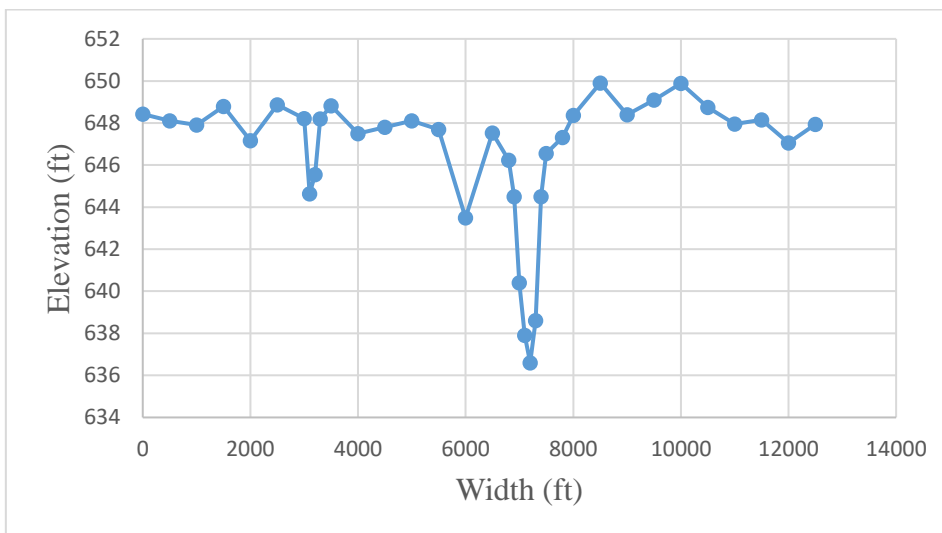


Figure D.39: Ravi Section at RD 135000

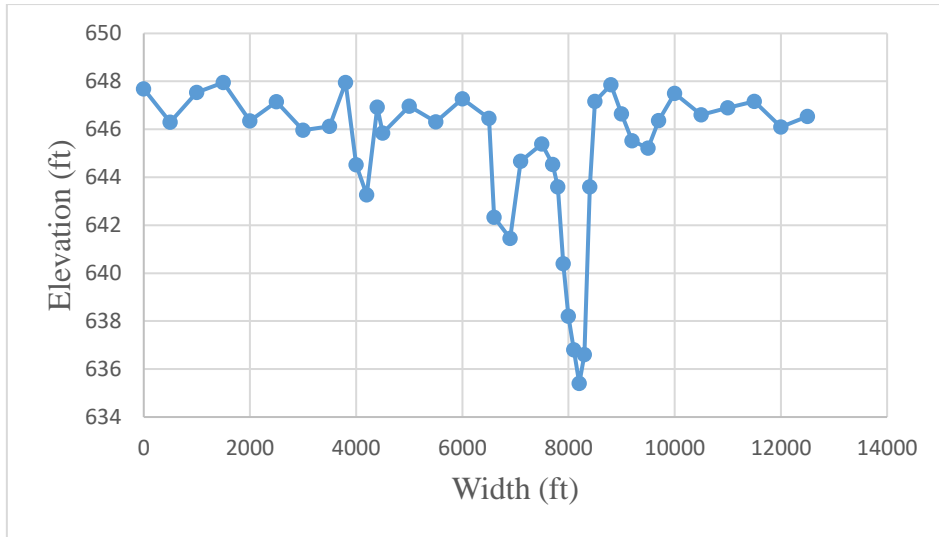


Figure D.40: Ravi Section at RD 138000

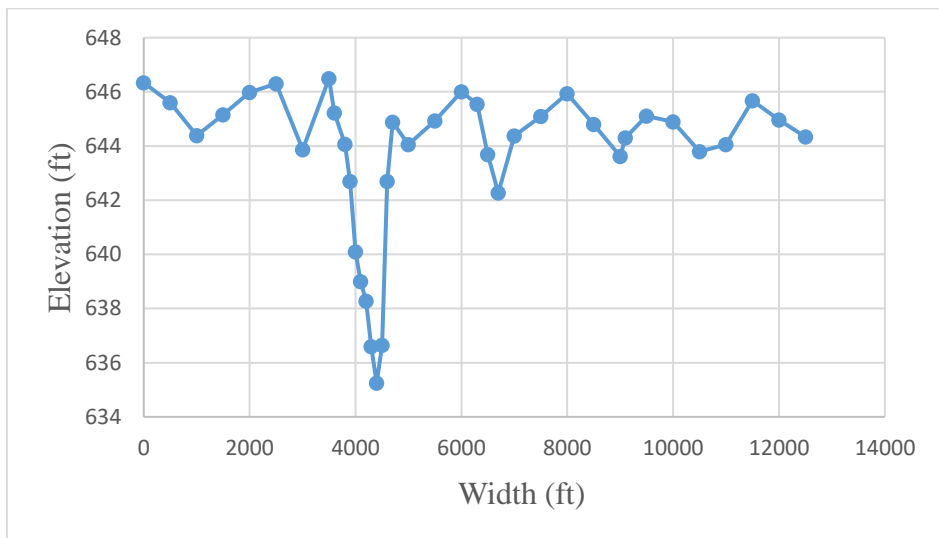


Figure D.41: Ravi Section at RD 141000

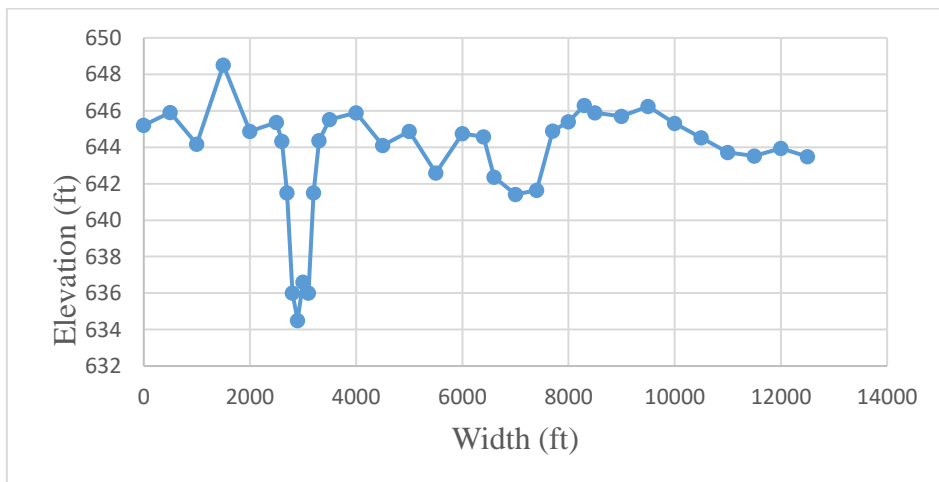


Figure D.42: Ravi Section at RD 144000

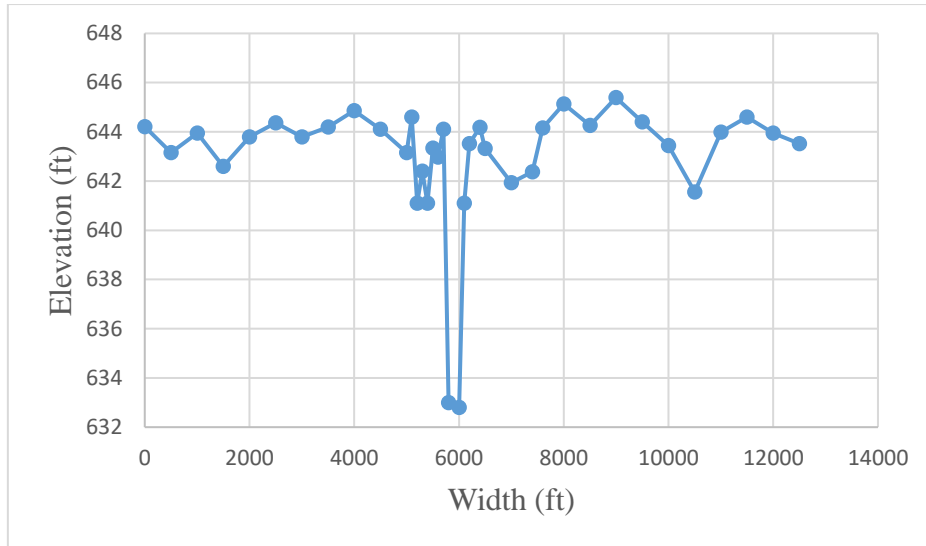


Figure D.43: Ravi Section at RD 147000

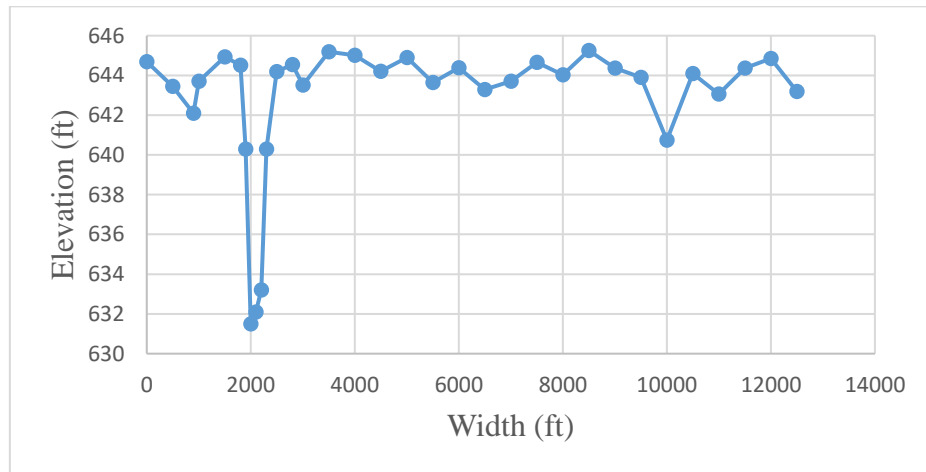


Figure D.44: Ravi Section at RD 154000

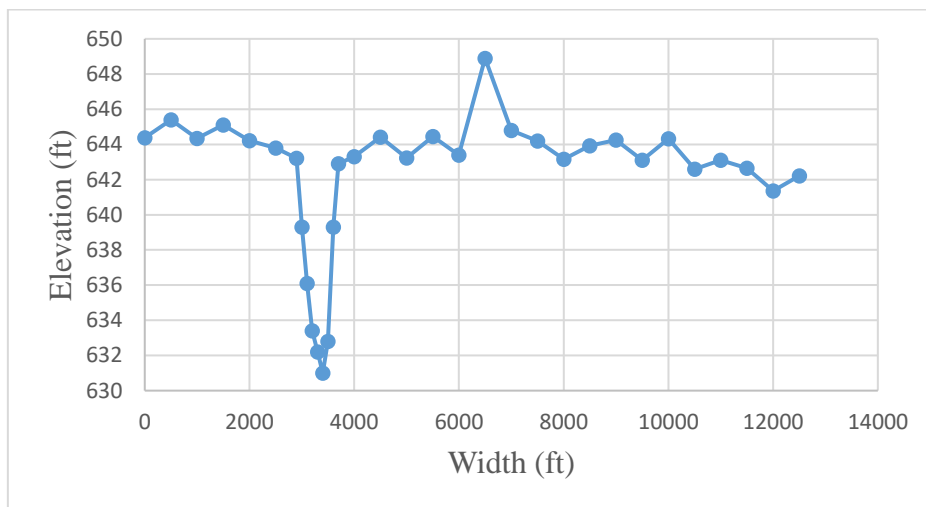


Figure D.45: Ravi Section at RD 156500

Annexure E

Observed Groundwater level maps of Lahore city

Simulated Groundwater level maps of Lahore city

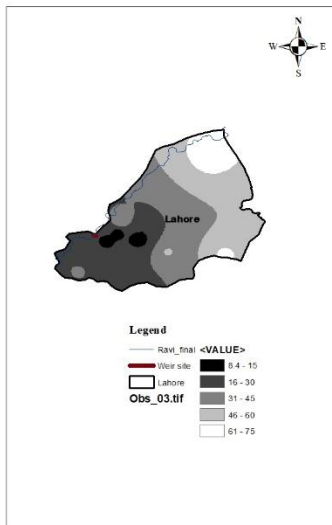


Figure E.1: Observed WL (ft.) map of Lahore city in year 2003

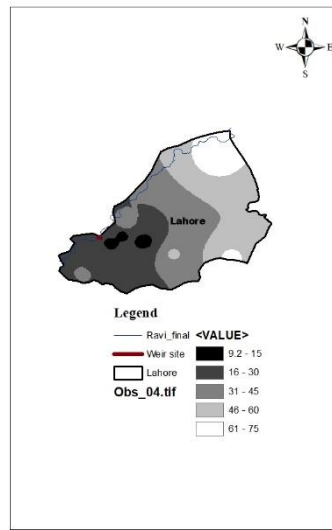


Figure E.2: Observed WL (ft.) map of Lahore city in year 2004

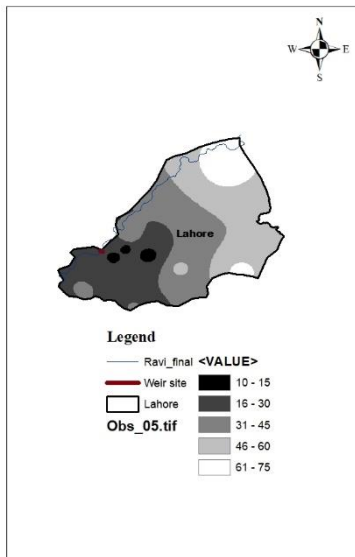


Figure E.3: Observed WL (ft.) map of Lahore city in year 2005

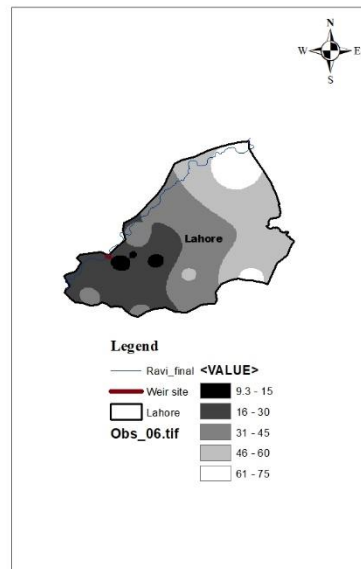


Figure E.4: Observed WL (ft.) map of Lahore city in year 2006

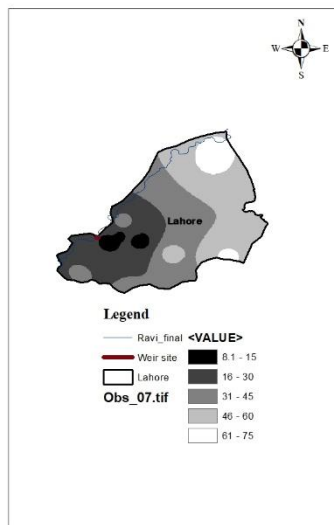


Figure E.5: Observed WL (ft.) map of Lahore city in year 2007

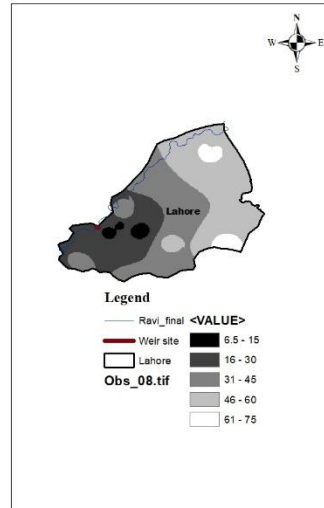


Figure E.6: Observed WL (ft.) map of Lahore city in year 2008

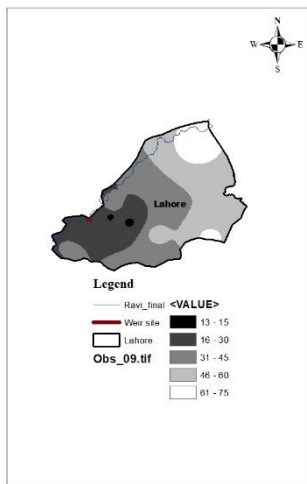


Figure E.7: Observed WL (ft.) map of Lahore city in year 2009

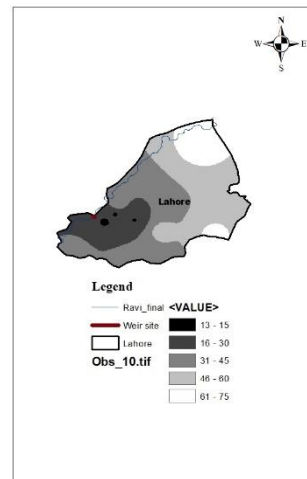


Figure E.8: Observed WL (ft.) map of Lahore city in year 2010

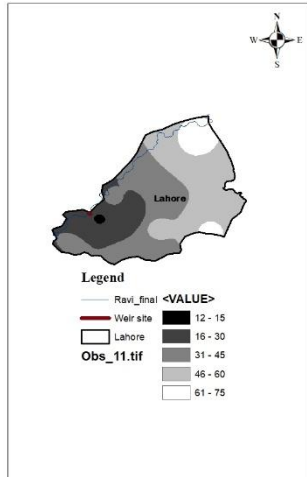


Figure E.9: Observed WL (ft.) map of Lahore city in year 2011

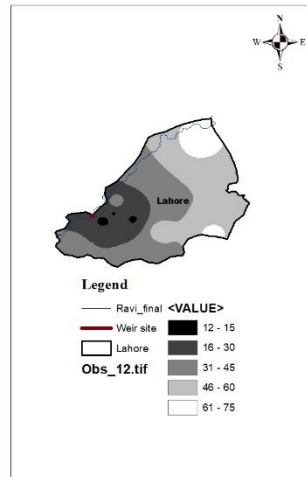


Figure E.10: Observed WL (ft.) map of Lahore city in year 2012

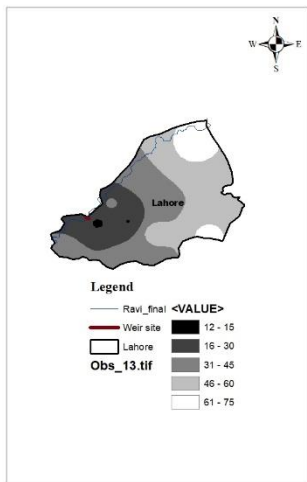


Figure E.11: Observed WL (ft.) map of Lahore city in year 2013

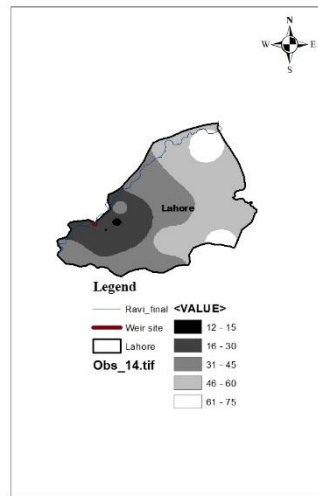


Figure E.12: Observed WL (ft.) map of Lahore city in year 2014

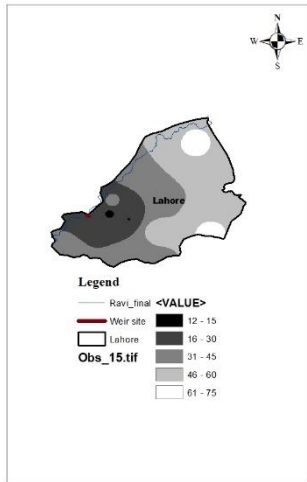


Figure E.13: Observed WL (ft.) map of Lahore city in year 2015

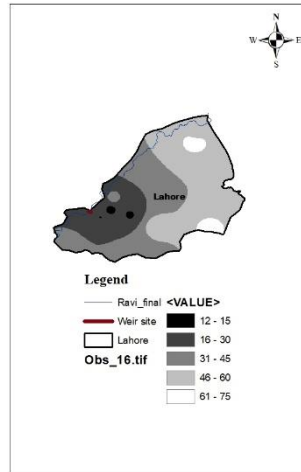


Figure E.14: Observed WL (ft.) map of Lahore city in year 2016

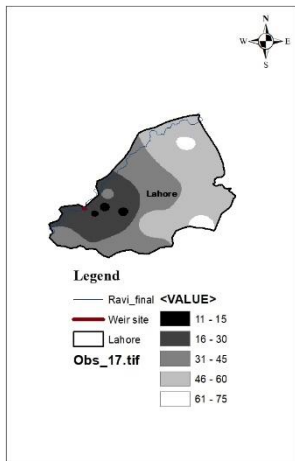


Figure E.15: Observed WL (ft.) map of Lahore city in year 2017

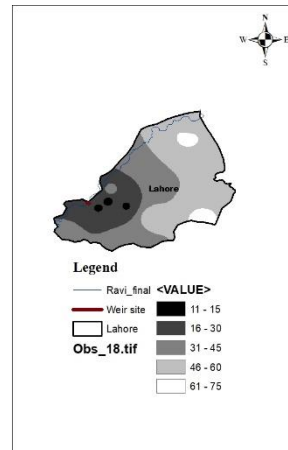


Figure E.16: Observed WL (ft.) map of Lahore city in year 2018

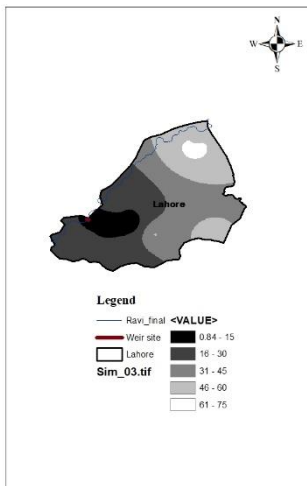


Figure E.17: Simulated WL (ft.) map of Lahore city in year 2003

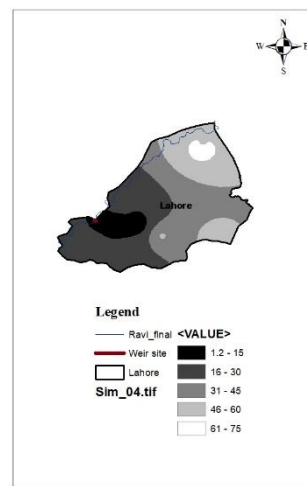


Figure E.18: Simulated WL (ft.) map of Lahore city in year 2004

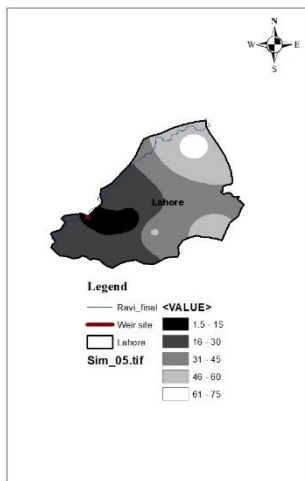


Figure E.19: Simulated WL (ft.) map of Lahore city in year 2005

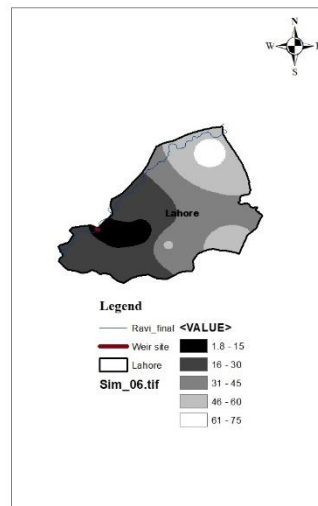


Figure E.20: Simulated WL (ft.) map of Lahore city in year 2006

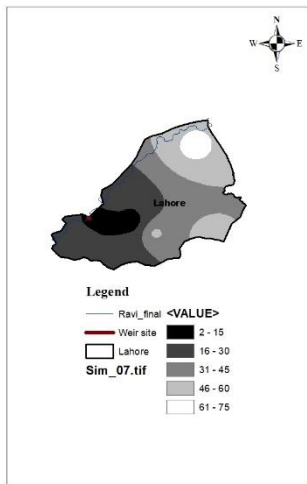


Figure E.21: Simulated WL (ft.) map of Lahore city in year 2007

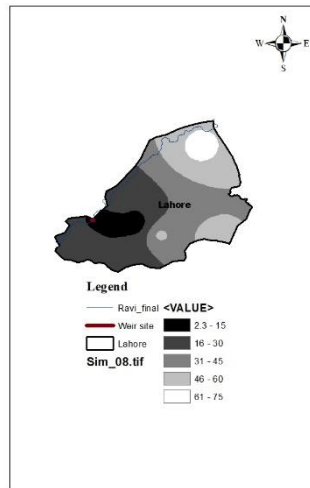


Figure E.22: Simulated WL (ft.) map of Lahore city in year 2008

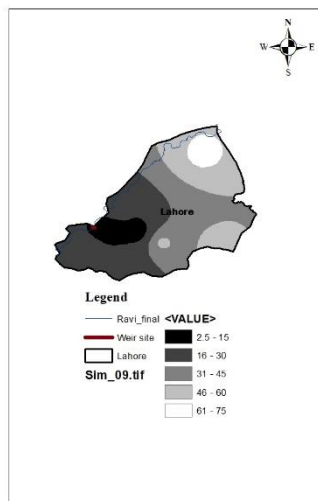


Figure E.23: Simulated WL (ft.) map of Lahore city in year 2009

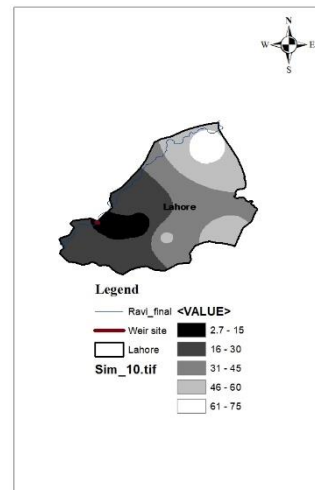


Figure E.24: Simulated WL (ft.) map of Lahore city in year 2010

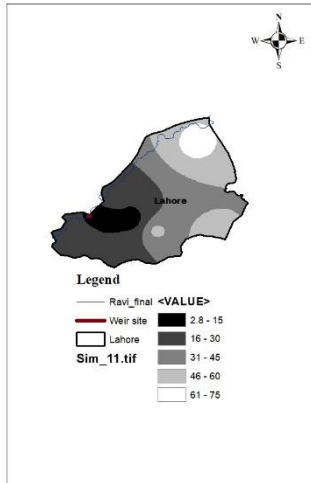


Figure E.25: Simulated WL (ft.) map of Lahore city in year 2011

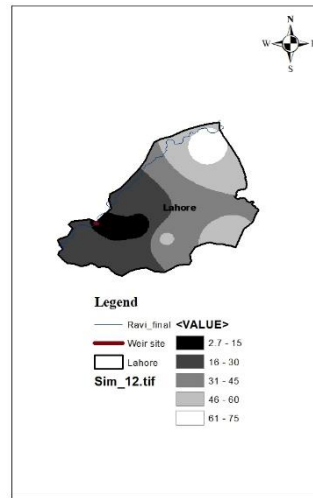


Figure E.26: Simulated WL (ft.) map of Lahore city in year 2012

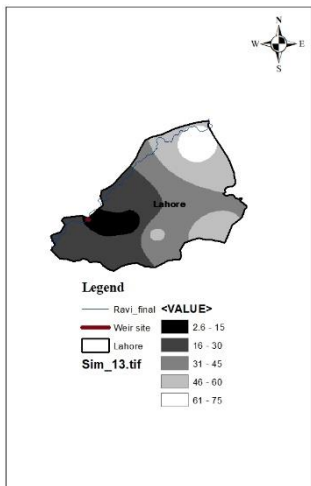


Figure E.27: Simulated WL (ft.) map of Lahore city in year 2013

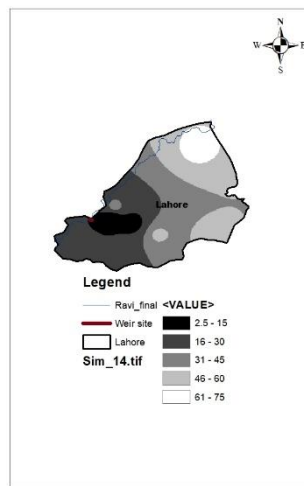


Figure E.28: Simulated WL (ft.) map of Lahore city in year 2014

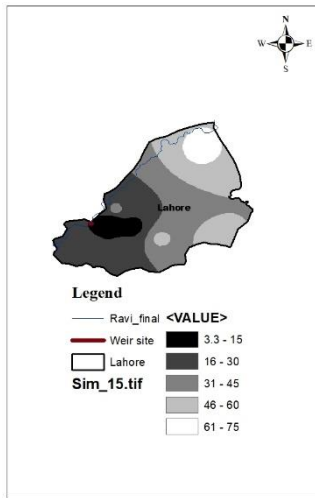


Figure E.29: Simulated WL (ft.) map of Lahore city in year 2015

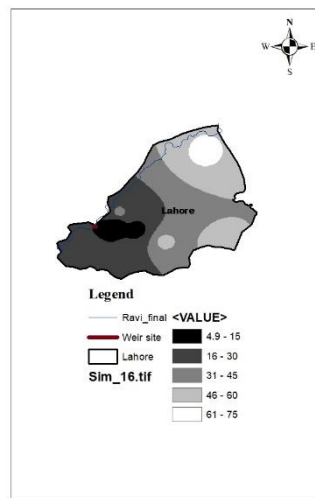


Figure E.30: Simulated WL (ft.) map of Lahore city in year 2016

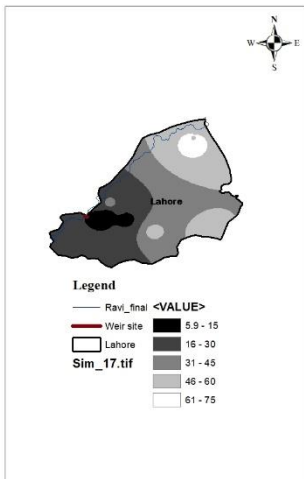


Figure E.31: Simulated WL (ft.) map of Lahore city in year 2017

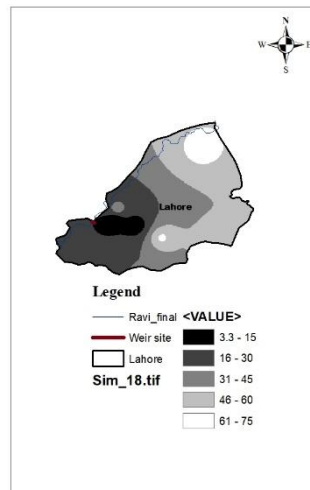


Figure E.32: Simulated WL (ft.) map of Lahore city in year 2018

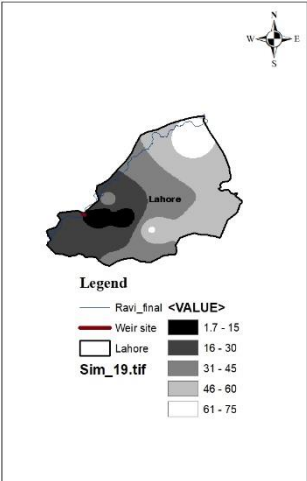


Figure E.33: Simulated WL (ft.) map of Lahore city in year 2019

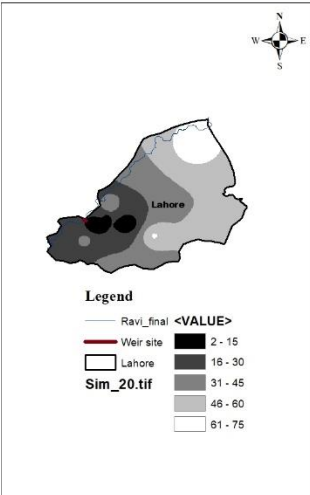


Figure E.34: Simulated WL (ft.) map of Lahore city in year 2020